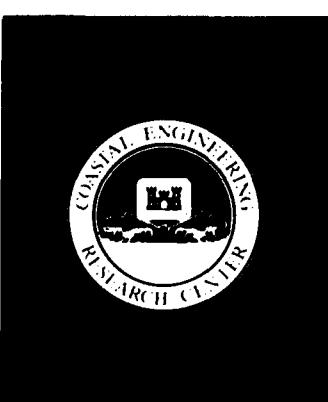
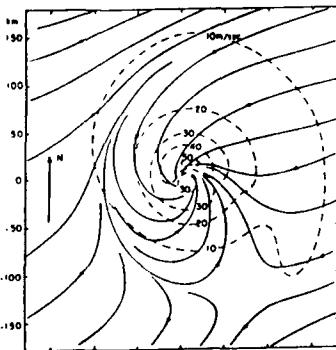
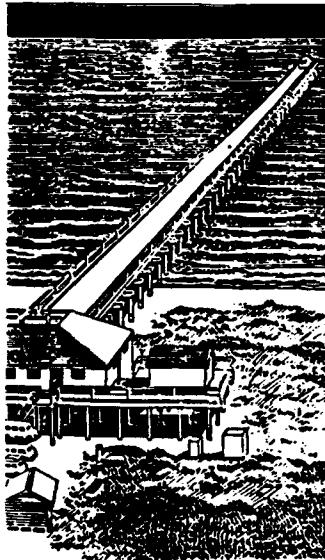


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of Engineers



CONTRACT REPORT CERC-92-1

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# UNIFIED PROGRAM FOR THE SPECIFICATION OF HURRICANE BOUNDARY LAYER WINDS OVER SURFACES OF SPECIFIED ROUGHNESS

by

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13. ABSTRACT (Maximum 200 words)

A method is developed to specify the surface stress and the wind speed and direction in the planetary boundary layer of a tropical cyclone from meteorological storm parameters available for historical hurricanes. The method is based upon a numerical primitive-equation model of the planetary boundary layer in a moving tropical cyclone. The complete time history of the evolution of the surface wind field is described from a series of characteristic wind field states calculated at discrete times in a storm's history by the steady-state model.

A surface drag formulation, based upon a contemporary similarity model (Arya 1977), coupled with a roughness parameter specification for a water surface consistent with Cardone's (1969) law, is incorporated into the numerical model and found to produce a consistent description of the integrated planetary boundary layer wind, the surface stress and its direction, and the wind speed and direction at anemometer level. The surface winds calculated in several recent hurricanes are found to be in excellent agreement with available, representative surface wind measurements made from offshore platforms and data buoys.

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Transformations based upon an equilibrium planetary-boundary-layer similarity model are developed to specify the surface wind over terrain of specified roughness, including lake surfaces, from the over-water wind-field solution. Calculated over-land and over-lake winds are compared to the limited measurements available for several recent storms. Agreement is generally good.

The method is incorporated in a computer program, which provides surface winds on a variable-resolution rectangular grid. This report includes program documentation and sample grid results for a test simulation performed on Hurricane Betsy (1965).

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## PREFACE

This report describes the methods incorporated in a computer program developed to provide hurricane surface wind fields. The wind fields can be used in wave and surge modeling activities. The report also serves to document the computer program delivered as part of the study.

The work described in this report was originally performed under Work Unit No. 12114, "Wave Information Studies," Coastal Field Data Collection Program. Publication of the report is funded by Work Unit No. 32683, "Wind Estimation for Coastal Modeling," Coastal Research Program. Both programs are sponsored by Headquarters, US Army Corps of Engineers (HQUSACE). Messrs. John H. Lockhart, Jr., and John G. Housley were the HQUSACE Technical Monitors. Ms. Carolyn M. Holmes of the Coastal Engineering Research Center (CERC), US Army Engineer Waterways Experiment Station (WES), was Program Manager. Drs. Jon M. Hubertz and Edward F. Thompson were the Principal Investigators of Work Unit Nos. 12114 and 32683, respectively.

This study was conducted under Contract No. DACW39-78-C-0100 by Oceanweather, Inc., Cos Cob, Connecticut, and provided to CERC on October 31, 1979. This report is the original contract report provided to CERC. It is being published as a CERC Contract Report at this time because CERC has used and continues to use the study results extensively to estimate wave growth in hurricanes. This report provides an important historical basis for present CERC practice. The hurricane wind model described in this report has recently been modified under Work Unit No. 32683 and included in CERC's Coastal Modeling System (Instruction Report CERC-91-1). Both work units are under the direct supervision of Dr. Martin C. Miller, Chief, Coastal Oceanography Branch, and Mr. H. Lee Butler, Chief, Research Division, and under the general supervision of Mr. Charles C. Calhoun, Jr., Assistant Director, CERC, and Dr. James R. Houston, Director, CERC.

The authors express their appreciation to the technical contract monitor at the time the work was performed, Dr. Robert W. Whalin, and to the late Dr. Charles E. Abel, both of WES, for their support, assistance and patience throughout the course of the work. The authors also acknowledge the role of the late Dr. John Wanstrath, who recognized the critical need for improved wind field models in hurricane surge modelling, and whose intense interest

**stimulated the initiation of this research study.**

**At the time of publication of this report, Director of WES was Dr. Robert W. Whalin. Commander and Deputy Director was COL Leonard G. Hassell, EN.**

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**CONVERSION FACTORS, NON-SI TO SI (METRIC)  
UNITS OF MEASUREMENT**

Non-SI units of measurement can be converted to SI (metric) units as follows:

<u>Multiply</u>	<u>By</u>	<u>To Obtain</u>
degrees (angle)	0.01745329	radians
feet	0.3048	metres
knots (international)	0.5144444	metres per second
miles (US nautical)	1.852	kilometres

A UNIFIED PROGRAM FOR THE SPECIFICATION OF HURRICANE BOUNDARY  
LAYER WINDS OVER SURFACES OF SPECIFIED ROUGHNESS

PART I: INTRODUCTION

Statement of Problem

1. Specification of the wind velocity and vector wind stress near the sea surface in tropical cyclones is required for the description of ocean surface phenomena (such as ocean currents, the storm surge, and surface gravity waves) related to such cyclones. Dynamical-numerical computer-based models to describe such phenomena continue to be developed: for example, the work of Jelesnianski (1970) and Wanstrath et al. (1976) on the open coast surge, Forristall (1974) on currents, Cardone et al. (1976) on waves, and Wanstrath (1978) on the surge, including coastal flooding.

2. In most studies very simple descriptions of the hurricane wind field have been used to drive very complicated ocean response models. This disparity has hindered further refinement and validation of ocean response models and has limited the defensibility of climatological series, design criteria, and like results based upon the application of such models. The lack of near-surface wind measurements in hurricanes serves as an excuse for applying simple wind models: in fact, the most widely applied empirical model, discussed below, of surface marine wind distribution in hurricanes is calibrated against wind measurements made in a lake during the passage of a single storm.

3. In recent years, however, great progress has been made in our understanding of the basic physical and dynamical characteristics of tropical cyclones, including the part of such storms relevant to this discussion: the planetary boundary layer (PBL). Further, a series of measurement programs, both public and private, employing offshore oil and gas production platforms, automated data buoys, and low-flying aircraft, have made available a wealth of data on wind structure in the PBL in

tropical cyclones.

4. Cardone et al. (1976), exploiting this recent progress, developed a method for specifying the surface wind field in hurricanes over the ocean by applying a dynamical-numerical model of the PBL in hurricanes. The method, requiring as input only a description of the surface pressure field and specification of storm motion and latitude, has been used to model the surface wind field in nearly every major hurricane to affect the Gulf of Mexico or the East coast of the United States in the past decade (Camille, 1969; Delia, 1973; Eloise, 1975; Belle, 1976). At least one representative series of wind measurements over water was available in each of those storms to validate the method for the intended environment (open water) and the intended parameter (time-averaged winds at a specific height). Those studies confirm that the model is able to give a convincing numerical representation of how friction, latitude, storm motion, and the shape and intensity of the sea-level pressure pattern in a severe storm interact to produce an asymmetrical vertically integrated flow in the PBL. The model, however, includes an empirically based scaling law to relate the integrated boundary layer wind to the effective 19.5 meter level wind. Further, the surface stress distribution in the numerical solution has not been validated.

5. The purpose of the study reported here is to generalize the method so that it can be applied to specify surface wind and wind stress in hurricanes in a self-consistent way not only over the open sea but over waters typical of the near-shore environment, over inland bodies of water (lakes), and over terrain of varying roughness in general (e.g. open marshland, dense forest, cities). Such a capability is required for the application of surge models which treat the open-coast storm tide and inland flooding of coastal areas (e.g. the Mississippi delta and Lake Pontchartrain area) or those applied to surge studies of large inland bodies of water (e.g. Lake Okeechobee).

6. The goal of this study was the development of an efficient algorithm, implemented as a computer program free from proprietary constraints, that can be used to specify hurricane-generated surface winds and wind stresses, in historical storms or hypothetical storms, from the

kind of meteorological information available for historical storms.

### Review of Prior Methods

7. Considered theoretically, the problem of surface wind specification in tropical cyclones is to solve the basic equations of hurricane-scale circulation, subject to appropriate initial and boundary conditions. As a part of the system of equations, consider the primitive equations of motion in cylindrical coordinates  $r$ ,  $\lambda$ ,  $z$ , with origin at the center of the cyclone, for  $u$  and  $v$ , the (horizontal) tangential and radial components of the wind:

$$\rho \left\{ \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial r} + \frac{v}{r} \frac{\partial u}{\partial \lambda} + w \frac{\partial u}{\partial z} - fv - \frac{v^2}{r} \right\}$$

$$= \frac{\partial p}{\partial r} + \frac{\partial \tau_{zr}}{\partial z} + \left\{ \frac{\partial \tau_{rr}}{\partial r} + \frac{1}{r} \frac{\partial \tau_{\lambda r}}{\partial \lambda} + \frac{\tau_{rr} - \tau_{r\lambda}}{r} \right\}$$

(1a)

$$\rho \left\{ \frac{\partial v}{\partial t} + u \frac{\partial v}{\partial r} + \frac{v}{r} \frac{\partial v}{\partial \lambda} + w \frac{\partial v}{\partial z} + fu + \frac{w}{r} \right\}$$

$$= \frac{1}{r} \frac{\partial p}{\partial \lambda} + \frac{\partial \tau_{z\lambda}}{\partial z} + \left\{ \frac{\partial \tau_{r\lambda}}{\partial r} + \frac{1}{r} \frac{\partial \tau_{\lambda\lambda}}{\partial \lambda} + \frac{2\tau_{r\lambda}}{r} \right\}$$

(1b)

where  $w$  is the vertical component of the wind and the  $\tau$ 's are the radial, tangential, and vertical eddy stresses of the tangential and radial velocity components.

8. Lack of knowledge how to solve equations (1), in particular lack of knowledge how to specify the eddy stresses, has until rather recently prevented a straightforward application of the primitive equations to the study of tropical cyclones. Practical demands, however, dictated the development of surface wind models as early as the late 1940's and the early 1950's. Even today the most widely applied method for wind specification derives from a procedure first published in a se-

ries of reports of the U.S. Weather Bureau, Hydrometeorological Section, during the 1950's. This method, hereafter referred to as the HP (Hydrometeorological-Parametric) method, may be considered a three-parameter method: it requires knowledge of only three quantities related to a storm's structure, to specify the entire areal distribution of surface wind.

9. The basic steps involved in the HP method are the following:

- a. Assume that the sea-level pressure distribution in a tropical cyclone is symmetric; specify pressure as a function of radius  $r$ ,

$$p(r) = p_0 + \Delta p e^{-R/r},$$

where  $p_0$  is the central pressure (at the eye),  $\Delta p$  is a storm pressure anomaly, and  $R$  is a scale radius related to the radius of maximum wind.

- b. Compute the profile of gradient wind as a function of radius: this step basically solves the  $u$  equation of motion for the steady-state, frictionless, locally horizontally homogenous solution:

$$\frac{u^2}{gr} + \frac{u}{fr} \frac{gr}{fr} = \frac{1}{\rho f} \frac{\partial p}{\partial r}.$$

- c. Attempt to compensate for neglect of all other terms by reducing the gradient wind to the equivalent wind over water at 10 meters, using a reduction of the form

$$u_{10}/u_{gr} = F(r/R)$$

where  $F$  varies from .865 at the radius of maximum wind to about .60 at the periphery of the storm.

- d. Add asymmetry to the storm by vectorially adding 50% of the storm's forward motion  $V_f$  to the right side of the storm and subtracting 50% from the left side.
- e. Partition the reduced winds into tangential and radial components by specifying an inflow angle that is a function of  $r/R$  alone; in effect, this step attempts to compensate for the neglect of all physical effects contained in the  $v$  equation of motion.

10. The function  $F$  was derived from pressure and wind measurements made in and around Lake Okeechobee, Florida, during the passage of two hurricanes in 1949 and 1950; in fact, the constant .865 was chosen

from measurements in that single storm deemed more representative.

11. An attempt to solve by graphical means a less simplified version of equations (1), proposed by Myers and Malkin (1961), was applied to the determination of the wind field in hurricane Helene, 1958, by Schauss (1962). The study of Myers and Malkin (1961) was the first to demonstrate the dynamical inconsistencies in the HP method, especially the fallacy that superposition of storm motion on a circularly symmetric wind field can produce the right/left asymmetry observed in hurricanes. Schauss (1962) represented the eddy stresses in terms of friction coefficients which in turn were estimated indirectly from sea-level pressure analyses and ships' wind observations in hurricane Helene. An important element of Schauss' study was the realization that the pressure field about tropical cyclones is not axisymmetric. Schauss varied  $\Delta p$  and  $R$  by quadrant; he found that he could reliably estimate the quadrantal variation from conventional coastal measurements and marine pressure data, even though Helene remained offshore, east of the East coast of the United States.

12. The method of Myers and Malkin and Schauss is not, to our knowledge, in use today; the HP method, however, is in widespread use in design studies, climatological studies, and real-time forecast systems. Variants of the HP method are described by Patterson (1972) and Bretschneider (1976); the latest version is documented in Memorandum HUR7-120, Hydrometeorology Branch, Office of Hydrology, NOAA.

#### Relevant Recent Research

13. Much of our current basic knowledge of the structure, dynamics, and energetics of tropical cyclones has been generated only within the past 10 or 15 years. This progress is a result of two factors: first, an extensive data base has been created as a result of the program of reconnaissance by NOAA research aircraft begun in the late 1950's; second, models of the tropical cyclone, based upon numerical integration of the primitive equations, have provided new insight into the basic dynamical and thermodynamic processes operative in tropical cyclones.

14. As a result of analysis of the reconnaissance data, the general distribution of pressure, wind, and temperature throughout the free atmosphere above the friction layer in the inner core of hurricanes has been revealed. Shea and Gray (1973), for example, composited all aircraft data obtained between 1957 and 1969, and derived the distribution of the mean tangential and radial components of actual winds at various levels including the 900 mb level, the level closest to the boundary layer. Their analysis suggested that at that level, the asymmetry in the tangential component exceeds the forward motion of the storm, and that the radial (inflow) component is not symmetrically distributed about the storm axis, as is assumed in the HP model.

15. In the area of numerical modelling, at least a dozen distinct prognostic models have been developed since 1964. Anthes (1974) summarized numerical models down to about 1972, several of which continue to be further developed today. Changes have been mainly in the use of increased horizontal and vertical resolution and in parametrizing the effects of cumulus convection.

16. Models of the prognostic type are designed to study the dynamics and energetics of tropical cyclones, to expose the mechanisms of hurricane formation from incipient tropical disturbances, to study the sensitivity of the tropical cyclone to boundary conditions at the lower boundary (in particular sea surface temperature), and to assess the prospective influence of various proposed schemes for artificially modifying such cyclones (e.g. cloud seeding near the eye wall). Such models typically can not be used directly to specify the distribution of surface wind or wind stress in hurricanes, because the models suffer from relatively low horizontal resolution, crude parametrization of the boundary layer, and time-dependent integration schemes: in fact, the models were designed not as diagnostic tools, but rather to describe the evolution of the circulation from an arbitrary set of initial conditions.

17. Recently there has been greater emphasis on the boundary layer of tropical cyclones in hurricane research: this is because the boundary layer is an important dynamical component of the total circulation in cyclones; additionally, much of the new knowledge gained within

the past decade about the structure of the surface and planetary boundary layers of the atmosphere over the sea can be profitably applied to the boundary layer in hurricanes. Elsberry et al. (1974) obtained realistic solutions for the temperature and moisture distribution within the boundary layer of an axisymmetric storm by applying the marine PBL model of Cardone (1969), originally developed for application to the extra-tropical atmosphere. More recently, Moss and Rosenthal (1975) estimated the vertical exchange rates of momentum, heat, and cloud mass in the boundary layer of hurricanes by combining the boundary layer model of Deardoff (1975) with the roughness parameter formulation of Cardone (1969). Anthes and Chang (1978) used a new parameterization of the planetary boundary layer in an axisymmetric numerical hurricane model to study the response of a hurricane boundary layer to changes of sea surface temperature.

18. In this study, a diagnostic model of the hurricane PBL, developed originally by Chow (1971) and applied later by Cardone et al. (1976), is modified to produce a consistent description of the vertically integrated wind in the PBL, the surface drag and the wind speed and direction at anemometer height in a moving hurricane with asymmetric horizontal wind distribution over water. Equilibrium PBL theory is used to extend the surface wind description to terrain of specified roughness. The wind model is incorporated in a computer program to provide a gridded temporal and spatial history of the surface wind for use in surge models.

## PART II: THE HURRICANE PBL MODEL

### Review of Chow's (1971) Vortex Model

19. As time dependent numerical hurricane models have been extended to three dimensions, it has become necessary to find efficient numerical schemes of solving the primitive equations on the high resolution grids required and to give more emphasis to the asymmetry of planetary boundary layer flow. The latter is important because frictionally induced convergence in the boundary layer (Ekman pumping) is an important mechanism in organizing moist convection and triggering the instability responsible for the development of tropical cyclones. Those requirements served as the principal motivation for Chow's study.

20. Chow's model concerned the planetary boundary layer (PBL) only and sought the solution for the wind field and horizontal convergence in the PBL of a moving tropical cyclone from the equations of motion. The pressure field in the boundary layer was prescribed and fixed, so that there would be no gravity waves excited in the numerical solution. This facilitated the use of a nested grid system, which allowed grid spacings as small as 5 km near the hurricane inner region without sacrifice of overall computational efficiency.

21. The model is based upon the equation of horizontal motion, vertically averaged through the depth of the PBL, written in coordinates fixed to the earth as

$$\frac{\hat{d}\vec{v}}{dt} + f|\vec{k} \times \vec{v}| = -\frac{1}{\rho} \nabla p + \nabla \cdot (K_H \vec{v}) - \frac{C_D}{h} |\vec{v}| \vec{v} \quad (2)$$

where

$$\frac{d}{dt} = \frac{\partial}{\partial t} + \vec{v} \cdot \vec{\nabla};$$

$\frac{\partial}{\partial t}$  is the time change local to the fixed coordinates;  $\vec{v}$  the two-dimensional del operator;  $\vec{v}$  the vertically averaged horizontal velocity;  $f$  the Coriolis parameter;  $|\vec{k}|$  the unit vector in the vertical direction;  $\rho$  the mean air density;  $p$  the pressure;  $K_H$  the horizontal eddy viscosity coefficient;  $C_D$  the drag coefficient;  $h$  the

depth of the planetary boundary layer. It is assumed that the vertical advection of momentum is small compared to the horizontal advection and can be neglected and that the shearing stress vanishes at the top of the PBL.

22. The pressure is prescribed as the sum of  $p_c$  and  $\bar{p}$ ;

$$p = p_c + \bar{p}$$

where  $p_c$ , not necessarily axisymmetric, is the pressure field representing the tropical cyclone and assumed to translate with the storm at a specified speed  $\vec{v}_c$ ; and  $\bar{p}$  is a large scale pressure field which may be specified by the corresponding constant geostrophic flow,  $\vec{v}_g$ , as

$$f|K \times \frac{\hat{\vec{v}}}{g} = -\frac{1}{\rho} \nabla \bar{p} \quad (3)$$

With this pressure specification, equation (2) may be written:

$$\frac{d\hat{\vec{v}}}{dt} + f|K \times (\hat{\vec{v}} - \hat{\vec{v}}_g) = -\frac{1}{\rho} \nabla p_c + \nabla \cdot (K_H \hat{\vec{v}}) - \frac{C_D}{h} |\hat{\vec{v}}| \hat{\vec{v}} \quad (4)$$

With respect to a moving Cartesian coordinate system ( $x, y$ ) whose origin is located at the moving low center of  $p_c$ , equation (4) is transformed into

$$\frac{d\hat{\vec{v}}}{dt} + f|K \times (\hat{\vec{v}} - \hat{\vec{v}}_g) = -\frac{1}{\rho} \nabla p_c + \nabla \cdot (K_H \nabla \cdot \hat{\vec{v}}) - \frac{C_D}{h} |\hat{\vec{v}} + \hat{\vec{v}}_c| (\hat{\vec{v}} + \hat{\vec{v}}_c) \quad (4)$$

where

$$\frac{d}{dt} = \left(\frac{\partial}{\partial t}\right)_c + \hat{\vec{v}} \cdot \nabla$$

$$\left(\frac{\partial}{\partial t}\right)_c = \frac{\partial}{\partial t} + \hat{\vec{v}}_c \cdot \nabla$$

$$\hat{\vec{v}} = \hat{\vec{v}} - \hat{\vec{v}}_c$$

$$\hat{\vec{v}}_g = \hat{\vec{v}}_g - \hat{\vec{v}}_c$$

$\vec{v}$  is now the horizontal wind relative to the low center;  $\vec{v}_g$  the effective geostrophic flow relative to the low center; and  $(\frac{\partial}{\partial t})_c$  the time change local to the moving coordinates.

23. Chow solves equation (5) in component form

$$\frac{\partial u}{\partial t} = fv - Au - Pu + Hu - Fu \quad (5a)$$

$$\frac{\partial v}{\partial t} = -fu - Av - Pv + Hv - Fv \quad (5b)$$

where

$$Au = u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y}$$

$$Av = u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y}$$

$$Pu = fv_g + \frac{1}{\rho} \frac{\partial P_c}{\partial x}$$

$$Pv = -fu_g + \frac{1}{\rho} \frac{\partial P_c}{\partial y}$$

$$Hu = \frac{\partial}{\partial x} (K_H \frac{\partial u}{\partial x}) + \frac{\partial}{\partial y} (K_H \frac{\partial u}{\partial y})$$

$$Hv = \frac{\partial}{\partial x} (K_H \frac{\partial v}{\partial x}) + \frac{\partial}{\partial y} (K_H \frac{\partial v}{\partial y})$$

$$Fu = \frac{C_D}{h} [(u + u_c)^2 + (v + v_c)^2]^{1/2} (u + u_c) \quad (6a)$$

$$Fv = \frac{C_D}{h} [(u + u_c)^2 + (v + v_c)^2]^{1/2} (v + v_c) \quad (6b)$$

24. The general formulation is completed with the specification of the form of  $C_D$ ,  $K_H$  and the boundary condition at the outermost boundary of the grid. Following Smagorinsky (1963)

$$K_H = 2\kappa^2 \left(\frac{\Delta}{2}\right)^2 |\text{Def}|$$

where  $|\text{Def}|$  is the total deformation,  $\Delta$  is the mesh size and  $\kappa$  is a non-dimensional constant ( $\kappa = .4$  is assumed). The drag coefficient was assumed to increase linearly with wind speed

$$C_D = (0.5 + 0.6 |\vec{v}|) \times 10^{-3} \quad (\vec{v} \text{ in m/sec}) \quad (7)$$

At the outermost boundary of the grid, the acceleration and the horizontal diffusion of momentum are neglected, implying a balance between Coriolis force, pressure gradient force and the surface frictional force

$$f|k \times (\vec{V} - \vec{V}_c) = - \frac{1}{\rho} \nabla p_c - \frac{C_D}{h} |\vec{V} + \vec{V}_c| (\vec{V} + \vec{V}_c)$$

25. The computational grid is a rectangular nested grid system consisting of five nests, within each of which the mesh is constant. Figure 1 shows the inner three nests in one quadrant of the grid system; if the mesh size of the innermost nest is say 5 km, the second through fifth mesh sizes are 10, 20, 40, and 80 km respectively, and the entire grid covers an area of  $(1600 \text{ km})^2$ .

26. The details of the finite difference formulation and of the computational scheme are given by Chow (1971) and will not be repeated in detail here. Basically, a combination of diagonal and ordinary upstream differencing is used for spatial derivatives in order to reduce computational errors in the calculation of the advection terms and at the intermesh boundaries. The computation starts with an initial guess field consisting of the gradient wind components (computed from  $p_c$ ). At each grid point, the equations (5) are integrated forward in time until the acceleration  $(\partial \vec{V} / \partial t)_c$  is tolerably small. For an innermost grid mesh of size 5 km, the time step is 60 sec. Chow found that 800

iterations (equivalent to 13 hours 20 min) were sufficient to achieve a steady state solution.

27. Chow studied the accuracy of the numerical scheme by obtaining numerical solutions for a frictionless ( $K_H = C_D = 0$ ) stationary ( $\vec{V}_c = 0$ ) vortex in gradient balance. The pressure field was axisymmetric and defined by the well known exponential pressure law

$$p_c = p_o + \Delta p e^{-R_p/r} \quad (8)$$

where  $p_o$  is storm central pressure,  $\Delta p$  is the storm pressure anomaly,  $R_p$  is the scale radius and  $r$  is the radial distance from the eye.

For the test storm ( $\Delta p = 50$  mb,  $R = 40$  km), the truncation error decreased with the spacing of the inner nest mesh size. For a spacing of 5 km, the numerical solution for the radial and tangential components is compared to the analytical gradient-wind solution, which of course contains only a tangential component, in Figure 2. It is seen that the truncation error appears only in the radial component and serves to introduce a small inward-directed radial component.

28. Figure 3 shows the integrated boundary layer wind solution found by Chow for the same test storm but with friction, storm motion (10 m/sec) and a steering gradient of 10 m/sec aligned with the motion. The pattern is remarkably realistic, at least qualitatively, and displays considerable asymmetry in both speed and direction. In several sensitivity experiments Chow deduced that, at least for the test storm, friction and storm motion effects combine non-linearly to produce a strongly asymmetric inflow pattern. Storm motion is primarily responsible for front-back asymmetry in wind speed, while the asymmetric pressure field in the steering-flow case is mainly responsible for the left-right asymmetry in wind speed. These characteristics of the hurricane PBL wind field were proposed much earlier by Myers and Malkin (1961) on the basis of graphical integration of a vector equation of motion similar to equation (2), without lateral friction, but including both tangential and normal vertical friction forces.

## A Consistent Surface Stress Parameterization

### The model of Cardone et al. (1976)

29. Cardone et al. (1976) adapted Chow's model in order to specify hurricane surface winds in historical storms. While Chow's model provided a good qualitative description of the PBL wind field, Cardone et al. found it necessary to modify the model in several ways in order to attain good quantitative agreement between the modelled winds in real hurricanes and the limited amount of measured wind data available.

30. The model calibration rested mostly on the wind record measured on an oil rig directly in the path of Camille, 1969; at the time, this wind trace was believed to be the only existing accurate wind record representative of the surface boundary layer over open sea in a well-documented extreme hurricane.

31. The first modification made was to generalize the storm input parameter specification scheme adopted by Chow. The 3-parameter scheme, requiring only  $\Delta p$ , R, and  $\vec{V}_c$ , was retained; as was the 4-parameter scheme, which adds an ambient geostrophic ( $\vec{V}_g$ ) pressure gradient directed normal to the storm track, a so-called steering flow. A 5-parameter scheme allowed the angle between the storm track and the ambient gradient to be varied. Finally the option to specify  $\Delta p$  and R in equation 8 by storm quadrant was added, providing up to an 11-parameter scheme. The motion and pressure parameters, along with storm latitude, completely defined the integrated boundary-layer wind solution for a given storm.

32. The calibration against Camille data, and to a lesser extent Carla wind direction information, involved two major modifications to the model. First, it was necessary to modify the boundary layer depth formulation. Chow had assumed a constant depth of 1 km. In the modified version, the depth was allowed to increase toward the storm center from a minimum depth of 1 km at the storm periphery to nearly 2 km in the vicinity of the eye wall of intense storms. This modification primarily affected the directional properties of the integrated boundary-layer wind solution.

33. Second, since the model was needed to drive a wave prediction model which required wind data at 20 meter elevation, the wind speed calibration incorporated a scaling law to reduce the integrated boundary-layer wind to 20 meters. The law was found to be dependent on wind speed, with low winds reduced only slightly and winds of 50 m/sec reduced by about 60%.

34. The modified vortex model has been applied to the specification of the time history of surface winds in many historical storms (Ward et al., 1979) and in many recent storms which have affected the U.S. coast (Cardone and Ross, 1977; Ross and Cardone, 1977; Cardone and Ross, 1978). Since most hurricanes are not strictly in steady state (though the calibration storm Camille was remarkably so), a wind time history in a given storm is interpolated from several characteristic states computed from the vortex model, on the implicit assumption that there is a rapid mutual adjustment between the wind field and pressure field in hurricanes. The simulated storms all included some direct PBL wind data measured from offshore platforms, data buoys or aircraft. The model appears to provide a good surface wind description for a fairly wide range of storm types.

#### Shortcomings

35. Recent theoretical and field studies of the structure of the PBL in hurricanes have revealed several shortcomings in the modified model described above. First, with regard to the depth of the PBL, the evidence now suggests a much lower height than 1-2 km as assumed above. Moss (1978) studied the PBL turbulence structure from aircraft measurements for a peripheral portion of Eloise, 1975; the data support a PBL height of about 650 meters.

36. Moss and Rosenthal (1978) estimated PBL variables in two intense storms by applying Deardorff's (1972) PBL parameterization scheme to bulk data to compute the surface exchange coefficients. Cardone's (1969) relation for the roughness length was used and found to provide drag coefficients in reasonable agreement with previous estimates. The calculation included the estimation of the PBL depth under the assumption that the top of the PBL coincides with the cloud base, which was

taken as the lifting condensation level of surface air. Figure 4 shows the PBL depth calculated for the storms studied. A depth of 600-700 meters characterized both storms except near the eye wall where the depth lowers to about 500 meters. Finally, Chang (1977) added a PBL parameterization of contemporary formulation to a time dependent multilevel primitive equation model and derived the radial distribution of the PBL height in both steady and unsteady cyclones. Figure 5 shows that in the steady-state hurricane the depth varies between 380 and 450 m with the PBL height lifted slightly near the eye wall. The PBL height was not found to exceed these heights significantly in the several unsteady cases studied.

37. Another inconsistency in the model results from the retention of the drag law, equation (7), used by Chow. That law is actually intended for use with marine wind data at standard height (10 m). Since the drag coefficient decreases with increasing height in the PBL, the use of (7) with vertically integrated winds implies overestimation of the surface drag. In addition, new evidence (Figure 6) for the behavior of the 10 m drag coefficient,  $C_{10}$ , at sea reported by Garratt (1977) supports a Charnock-type roughness law

$$z_o = a \frac{u_*^2}{g} \quad (9)$$

where  $z_o$  is the roughness parameter,  $u_*$  is friction velocity and  $g$  is the gravitational constant. The Charnock constant  $a$  proposed was .0144 though the value is sensitive to the value assumed for the Kármán constant in the PBL model applied.

38. The scaling law adopted by Cardone et al. (1976) apparently compensates for the shortcomings noted above, at least insofar as the specification of the 20 meter wind speed is concerned. However, the computed surface wind is not consistent with the surface drag, which itself is likely to be incorrect. The integrated boundary-layer wind therefore is also likely to be in error in a given storm and is not suitable for the extension of the model solution to surfaces of arbitrarily specified roughness as required in this study.

39. In the next section, a revised PBL parameterization is adopted and shown to provide an accurate and consistent PBL wind and stress representation within the vortex model.

The revised PBL parameterization

40. A new framework for parameterization of the fluxes of momentum, heat and moisture in the PBL has been developed within the past decade, beginning mainly with the work of Blackadar and Tennekes (1968) and Zilitinkevich (1969). The parametric relations result from the matching of mean profiles of wind, temperature, and moisture predicted by surface and outer-layer similarity theories for a PBL in which the flow is assumed to be horizontally homogeneous and quasi-stationary.

41. A particularly convenient form of the parameterization, first proposed by Deardorff (1972), expresses the PBL fluxes in terms of layer-averaged mean PBL properties. Deardorff's parameterization, combined with Cardone's roughness parameterization, was found by Moss and Rosenthal (1977) to provide reasonable results for hurricanes. The parameterization adapted here is taken from Arya's (1977) update of Deardorff's scheme.

42. The general form of the parametric relations may be written

$$\frac{ku}{u_*} = - (\ln \hat{z}_o + A_m) \quad (10a)$$

$$\frac{kv}{u_*} = - B_m \text{ sign } f \quad (10b)$$

$$k(\theta_v - \theta_o)/\theta_* = - (\ln \hat{z}_o + C_m) \quad (10c)$$

$$k(q - q_o)/q_* = - (\ln \hat{z}_o + D_m) \quad (10d)$$

where  $u$  and  $v$  are the vertically integrated (as in equation 5) horizontal wind components (in the direction of the surface shear and perpendicular to it, respectively),  $\theta_v$  and  $q$  are the mean layer virtual potential temperature and specific humidity, respectively,  $\hat{z}_o$  is the roughness parameter normalized by the PBL height ( $z_o/h$ ),  $k$  is von Kármán's constant,  $\theta_*$  is a potential temperature scale expressed in

terms of the heat flux  $H$ ,  $q_*$  is a specific humidity scale involving the moisture flux, and  $A_m$ ,  $B_m$ ,  $C_m$ ,  $D_m$  are universal functions of dimensionless similarity parameters.

43. The Monin-Obukov length  $L$  may be expressed in terms of  $\theta_*$  and  $u_*$  since

$$L = \frac{-U_*^3 \theta v \rho c}{kgH} = \frac{-U_*^2 \theta v}{g \theta_*} \quad (11)$$

44. There exist two competing theories for the form of universal functions. In one theory, known as Rossby-number similarity theory, the boundary layer height is uniquely determined by  $u_*/f$  and  $L$ . For the near neutral hurricane PBL, that theory predicts the PBL to increase as the ratio  $u_*/f$  increases toward the center of storms, in apparent contradiction to observation.

45. In the generalized theory, the depth of the PBL,  $h$ , is specified as an independent variable. Arya (1977) presents updated expressions for the similarity functions in terms of this generalized theory as follows:

$$A_m = \ln(-\frac{h}{L}) + \ln \frac{fh}{u_*} + 1.5$$

$$B_m = 1.8 \frac{fh}{u_*} e^{-2h/L} \quad \frac{h}{L} \leq -2 \quad (12)$$

(unstable)

$$C_m = \ln(-h/L) + 3.7$$

$$A_m = -.96(h/L) + 2.5$$

$$B_m = .80(h/L) + 1.1 \quad \frac{h}{L} \geq +2 \quad (13)$$

$$C_m = -2.0(h/L) + 4.7 \quad (\text{stable})$$

For near-neutral conditions,  $-2 < h/L < 2$ ,  $A_m$ ,  $B_m$ ,  $C_m$  are assumed to be given by linear interpolation between the above computed values at  $h/L = \pm 2$ .

46. In terms of the similarity relations (10), the drag coeffi-

cient with respect to the integrated PBL wind is

$$C_d = \frac{k^2}{[(\ln z_o + A_m)^2 + B_m^2]} \quad (14)$$

while the angle  $\beta$ , between the surface wind and the integrated PBL wind is

$$\beta = \tan^{-1}(v/u) \quad (15)$$

47. To incorporate the similarity theory in the hurricane model, two cases were considered:

- a. Land. In a PBL over land, the following parameters are prescribed:  $f$ ,  $z_o$ ,  $h$ ,  $\theta_v - \theta_o$ . The parametric relations then define the following functions

$$C_d = F_1(\vec{v})_{f, z_o, h, \theta_v - \theta_o} \quad (16)$$

$$\beta = F_2(\vec{v})_{f, z_o, h, \theta_v - \theta_o} \quad (17)$$

- b. Water. Over water, the roughness parameter is not known but can be expressed in terms of  $u_*$  through a Charnock-type relation (equation 9) or the form proposed by Cardone (1969). The parametric relations can then be solved for

$$C_d = F_3(\vec{v})_{f, h, \theta_v - \theta_o} \quad (18)$$

$$\beta = F_4(\vec{v})_{f, h, \theta_v - \theta_o} \quad (19)$$

48. Arya (1977) gives solutions graphically only for the condition  $fh/u_* = 1$ , in which case  $C_d$  and  $\beta$  can be expressed, for a given latitude, in terms of  $z_o$  and a bulk Richardson number. In this form, the theory is an updated version of Deardorff's scheme, which did not include dependence of  $A_m$ ,  $B_m$ ,  $C_m$ ,  $D_m$  on  $fh/u_*$ . In general, however,  $fh/u_*$  or  $z_o$  are not known a priori; equations (10-15) are solved by iteration starting from an initial guess on  $u_*$ .

49. In order to avoid the prohibitive computational expense of solving equations 10-15 iteratively at each grid point within the numerical vortex model, the assumptions are made that over water the air-sea

temperature difference and boundary layer height can be considered to be invariant over the domain of the storm. In view of the preceding discussion on  $h$ , this appears to be a reasonable approximation. Except for hurricanes crossing major ocean-current boundaries, the assumption of horizontal homogeneity of  $\theta_v - \theta_o$  also seems reasonable, especially for Gulf of Mexico and lower U.S. East Coast hurricanes.

Little is known about the characteristics of the PBL in hurricanes over land. However, given the high level of turbulent mixing there in hurricanes, it is reasonable to assume that an adiabatic lapse rate is established and that at least in the near-coastal zone, the boundary layer depth is close to that assumed for the over-water case.

50. Given the above conditions, the parametric relations  $F_1$ ,  $F_2$ ,  $F_3$ ,  $F_4$  can be found once for a given storm by iteration, and expressed in terms of tables. In practice, the upwind and crosswind drag coefficients, the ratio  $u_n/|\vec{v}|$  and the angle,  $\beta$ , between the surface wind and the integrated wind are computed for  $|\vec{v}| = 0.8(.8)80$  m/s and tabled. Values for intermediate wind speeds are found by linear interpolation.

#### Test results for the general parameterization

51. The behavior of the over water drag coefficient for typical hurricane conditions ( $\theta_v - \theta_o = -2^\circ\text{C}$ ,  $h = 650$  m,  $f = 10^{-4}$ ) according to the general parameterization is shown in Figure 7. The over water cases are computed for two separate values of the Charnock constant: .0144 and .035. The former value was recommended by Garratt (1977). However, in this model,  $k = .35$ , to be consistent with the Arya formulation, and  $a = .035$  provides a better fit to the 10 meter drag coefficient measurements analyzed by Garratt. The drag coefficient for the land case is for a roughness parameter of .08 m, a boundary layer depth of 650 m and neutral stability. For comparison, Figure 7 also includes the form adopted by Chow.

52. While it is reasonable to expect the new theory to yield drag coefficients lower than equation (7), the magnitude of the decrease and the form of the wind dependence seen was surprising. The revised parameterization was tested in the numerical vortex model with the Camille

inputs used in the study of Cardone et al. (1976). The revised solution was referred to 20 meters by solving

$$v_{20} = \frac{u_*}{k} \ln \left( \frac{20}{z_0} \right)$$

where  $u_*$  and  $z_0$  are determined by the numerical model, since for typical hurricane conditions, stability effects can be neglected in the surface layer.

53. Figure 8 compares the new solution and that derived by Cardone et al. Clearly, the general parameterization is underestimating the surface stress and therefore the surface wind. The difficulty with the general Arya formulation was traced to the dominating influence of the scale-height ratio  $fh/u_*$  on the similarity variables. This parameter was added by Arya mainly to describe the relative influence of the Coriolis force in an Ekman-type layer over a wide range of latitudes. In a hurricane, however, the parameter varies over two orders of magnitude because of the variation of  $u_*$  over an extreme range (typically 0-200 cm/sec) over a short distance. Considering the highly curved nature and spatial inhomogeneity of the flow in the hurricane PBL, that parameter is apparently irrelevant. Its influence was controlled therefore by restricting the solution to  $fh/u_* = 1$ .

54. The results of restricting the scale-height ratio to unity were dramatic. The revised dependence of drag coefficient on wind speed is shown in Figure 7 and the revised 20-meter wind profiles in Camille, for two boundary layer heights ( $h = 650, 325$  m) are shown in Figure 8. The comparisons of modelled and measured wind speeds at Rig 50 in Camille for three PBL heights are shown in Figure 9. In the range of  $h = 325-650$  m, the new theoretical calculation fits the measurements as well as the calibrated solution of Cardone et al. (1976).

#### Properties of the over-water solution

55. The final PBL parameterization chosen for the over-water case consists of the Arya (1977) parameterization, with the scale height ratio restricted ( $fh/u_* = 1$ ) and with the crosswind drag term retained both in the calculation of the surface stress in the numerical solu-

tion and in the derivation of surface-layer wind direction from the integrated boundary-layer wind direction. The Charnock constant,  $a$ , is assignable as is the value of  $k$ . However, since a value of  $k$  of .35 is consistent with Arya's model, an  $a$  of .035 provides a better fit to Garratt's (Figure 6) drag coefficient data than the value suggested by Garratt.

56. The properties of the new over-water hurricane wind model solution are most directly seen in the relatively simple case of a stationary, symmetric vortex. The solution may then be compared more readily to the gradient wind. This is done in Figure 10 for a stationary symmetric vortex with the scale radius ( $R_p = 12$  n.mi)\* and pressure anomaly ( $\Delta p = 105$  mb) of Camille. Note that the comparisons are only made at and outside the radius of maximum wind, since inside that radius, truncation errors remain fairly large for a small storm such as Camille.

57. The vortex model predicts integrated PBL winds which are supergradient within a radius of about  $3 \times R_p$ . The 20-meter winds, however, vary from 75 to 85% of the gradient wind. The surface inflow reaches a maximum of  $25^\circ$  at a radius of about  $5 \times R_p$  and decreases sharply at the eye wall. These features are entirely consistent with those deduced by Shea and Gray (1973) from composited low level aircraft data in hurricanes. Comparisons of modelled and measured winds over water in several recent storms will be presented in a later section.

#### Numerical Experiments over Mixed Terrain

58. The similarity PBL theory as modified above may be easily applied to the specification of  $C_D$  and  $\beta$  in the hurricane PBL over land. For a neutral PBL, the revised theory predicts that  $C_D$  and  $\beta$  are functions only of  $-z_o$ . For  $h$  of 500 m, some typical values of  $C_D$  ( $\beta$ ) are  $1.78 \times 10^{-3}$  ( $13.6^\circ$ ) for  $z_o = .04$  m and  $2.54 \times 10^{-3}$  ( $16.3^\circ$ ) for  $z_o = .16$  m.

59. The revised theory was tested in two ways. First, the numerical model was solved for the steady state wind field in a stationary

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\* A table of factors for converting non-SI units of measurement to SI (metric) units is presented on page 4.

symmetric intense hurricane situated entirely over homogeneous terrain. While this case is unrealistic, it was intended to provide an indication of the sensitivity of the numerical solution to a greatly increased drag. Second, the solution was sought for the more realistic case of an intense hurricane situated at the coast with terrain of homogeneous roughness under the northern half of the storm and open sea beneath the southern half. The storm was stationary since only in such a case is the numerical model, formed in a moving coordinate system, rigorous.

60. Solutions typical of both cases are shown in Table 1 where they are compared to the open-sea solution discussed above. The cases were run with symmetric stationary Camille inputs. The all-land case returned a symmetric solution as expected. The case shown is for  $z_0 = .08$  m, which is typical of open, level, countryside with low vegetation. Within about 200 km of the eye, the integrated PBL wind is found to be larger than the open-sea case, though surface winds are slightly lower. The inflow angle is about 50% greater than the open-sea solution. In effect, the increased surface drag results in a more intense vortex for a given pressure field, reflecting the well known dual role of friction in tropical cyclones. In nature, decreased evaporation is largely responsible for hurricane decay over land, and rapid decrease in PBL winds.

61. Two typical examples of the mixed land/sea solution are shown in Table 1. The solutions in these cases were quite asymmetrical. In general, the solution over-land matched quite closely the solutions found in strictly over land cases for comparable  $z_0$ . However, over-water, the solution departed significantly from the reference over-water case. In particular, for all values of land roughness attempted, the PBL winds over water downwind of land increased to values above the all-water case, thus causing the formation of a distinct wind maximum in the left rear quadrant (with respect to north) of the storm.

62. In Table 1, the vertically integrated winds over sea in the mixed land/sea solution are taken along a radial extending south from the storm center. For a land  $z_0$  of .04 m the maximum wind over water

Table 1

Comparison of over water, over land, and mixed terrain radial wind speed profiles in vortex model numerical solutions for a stationary, symmetric storm with pressure profile parameters  $\Delta p = 105$  mb,  $R_p = 12$  n.mi., and modified ( $f_b/U_* = 1$ ) Arya (1977) similarity PBL parameterization.

R	$V_g$	Over water		$z_o = .08m$		$z_o = .04m$		$z_o = .32m$	
		$\overrightarrow{ V }$	$\overrightarrow{v}_{20}$	$\overrightarrow{ V }$	$\overrightarrow{v}_{20}$	$\overrightarrow{ V }_{\text{sea}}$	$\overrightarrow{ V }_{\text{land}}$	$\overrightarrow{ V }_{\text{sea}}$	$\overrightarrow{ V }_{\text{land}}$
800	6.9	7.2	7.2	7.3	5.1	7.3	7.3	7.1	7.3
400	20.6	20.2	19.7	18.6	13.0	20.3	19.0	20.3	17.5
200	42.3	41.2	34.9	37.9	26.5	41.5	38.7	41.8	35.5
100	66.5	67.0	53.5	64.5	45.1	69.0	64.7	70.4	61.0
50	90.7	94.2	72.7	97.2	67.9	100.2	91.7	104.9	88.6
40	97.6	102.4	78.3	108.1	75.5	110.0	99.3	116.9	95.1
30	105.2	108.4	82.4	116.9	81.7	116.7	104.6	129.4	98.9
20	111.2	112.5	85.2	117.0	81.8	116.9	109.6	133.9	103.5

## Explanation of table:

- R radial distance from hurricane center (km)
- $V_g$  magnitude of the gradient wind
- $\overrightarrow{|V|}$  integrated boundary layer wind speed
- $\overrightarrow{v}_{20}$  20 meter wind speed
- $\overrightarrow{|V|}_{\text{sea}}$  in mixed terrain solution, the over-sea values are taken from radial over sea extending out from eye (at coast) normal to coastline.
- $\overrightarrow{|V|}_{\text{land}}$  as above except radial extended over land

All wind speeds are expressed in knots.

is increased about 5% over the all-water case, and about 15% for a land  $z_o$  of .32 m. Even greater increases were found in the left rear quadrant. A similar experiment for actual Camille inputs (not shown) led to comparable results, showing an unrealistic speed maximum in the left rear quadrant, whereas the nominal over-water Camille solution displayed a maximum in the right hand quadrant.

63. It is difficult to explain the precise cause of the anomalous model behavior in the mixed-terrain case. Test cases were run with modified over-land drag laws derived from Rossby-number similarity theory but no change in the basic structure of the solution was noted. The solution was verified to be steady-state in trial integrations carried out for 1600 rather than 800 iterations of the vortex model. Indeed, the behavior might be attributed to the basic model formulation, which forces a steady state solution unrealistically. However, in an experiment with a three-level, three-dimensional nested-grid numerical model, reported by Moss and Jones (1978), behavior similar to that found here has been seen in a simulation of an intense hurricane approaching the coastline. In their integration, it was noted that as the hurricane approached the coastline, even though the sea-level pressure in the storm began to rise, a transient wind-speed maximum greater than found in the all-water control case was established over water on the landward side of the vortex and extending to the left side of the circulation (looking down the track toward land).

64. Moss and Jones attribute the anomalous behavior of the wind field near landfall to changes in the pressure field induced by greatly increased inflow in the part of the wind field over land. In our simulation, however, the anomalous behavior is present even though the pressure field is fixed. More likely, the behavior is caused by the simplified treatment of the PBL vertical structure and the neglect of boundary-layer-adjustment processes which occur on vertical scales of the order of the depth of the PBL and on horizontal scales of the order of the grid spacing.

65. Basic knowledge of the behavior of the PBL flow across discontinuities in roughness is in a remarkably inchoate state.

The theories fall into two classes: (1) those which treat surface boundary layers only (e.g. Elliott, 1958); (2) those which treat the Ekman layer as a whole. Nearly all the available field data are restricted to the former case: there is considerable evidence that upon crossing a change in surface roughness, the wind profile is modified from below as a new (internal) boundary layer grows in thickness at a rate which depends upon stability and roughness. As a crude rule, the new boundary layer is established to a height given by one tenth the distance to the upwind change in roughness. If we were to extend this rule to the hurricane PBL, whose height is typically 500 m, we would expect the flow to adjust within 5 km. However, the more general Ekman-layer adjustment theories (e.g. Taylor, 1969) suggest a more complicated process, in which the wind-speed profile near the surface adjusts rapidly as in the surface-layer theories, but in which the surface stress, the turbulence structure and the wind direction take much longer (by nearly an order of magnitude) to attain equilibrium with the new surface. The theory and scant available data (Jensen, 1978) suggest also that the process is not symmetrical with respect to roughness transitions with the adjustment from rough to smooth conditions taking place at a slower rate than from smooth to rough. It appears therefore that the basic process of PBL adjustment in hurricanes needs to be better understood before the simulation of PBL flow across discontinuities within numerical hurricane models can be improved.

#### Equilibrium Theory Terrain Transformations

##### Empirical Evidence

66. The numerical experiments described above indicated that the numerical model could not of itself provide reliable wind fields over terrain of arbitrary roughness. This raised three questions:

- (a) are fetch effects near roughness discontinuities important enough to be accounted for empirically in the specification of surface winds in hurricanes?
- (b) can over-land PBL winds be prescribed from the over water numerical solution?

(c) are winds over lakes such as Pontchartrain and Okeechobee different from winds over open water, apart from fetch effects, other factors being equal?

67. Extensive empirical studies of the effect of fetch on winds over water downwind of the coastline have been reported by Richards et al. (1966) and Phillips and Irbe (1976), who processed large data sets obtained around and in the Great Lakes. Both studies employed the same analytical method. Extensive series of simultaneous surface wind measurements from coastal land stations and downwind ships and buoys were paired and stratified by wind speed, fetch, and stability. Both studies employed the same fetch and stability classes. The stability was parametrized by the temperature difference (land air minus lake water).

68. The problem of concern here is the adjustment of the surface wind over a shallow inland lake of the dimensions of lakes Pontchartrain and Okeechobee (fetch < 40 n.mi.) in hurricane conditions. The latter are characterized by winds > 8 m/sec and near-neutral stability. The data of Phillips and Irbe, and Richards et al., in those wind-speed and stability categories are plotted in Figure 11. Clearly, up to a fetch of at least 40 n. mi., no significant trend with fetch is evident in the wind-speed ratio (over-land : over-water). The ratio tends to be slightly larger for the more recent study, because the over-water winds in the recent study were measured at 3 meter height while in the study of Richards et al. the anemometer height was 10 meters. Since the first fetch class covered the range 0-5 n.mi., the implication is that the wind speed over the lake in the indicated classes and at heights up to 10 m attained equilibrium with the lake surface within the first few miles of the coast. This behavior is consistent with the height-to-fetch ratio of 1:10 noted above for surface layer adjustment.

69. It should be noted that, in both of the studies cited, a dependence of the wind ratio on fetch is proposed. The dependence however appears mainly in data from low wind speed (< 8 m/sec) classes and very stable or unstable stability classes. The fetch dependence therefore probably is caused by an adjustment of the turbulence in the PBL due to changes in stability rather than in surface roughness and is therefore

not likely to be important in the hurricane environment.

70. An interesting corroborative piece of evidence has been provided recently by remote sensing data. The Seasat mission has proved that the marine surface wind speed as might be measured at 10 m height from say a data buoy, can be measured by a radar scatterometer to an accuracy of 1 m/s or better. The radar backscatter cross-section of the sea surface ( $\sigma_0$ ) therefore appears to be mainly dependent on surface stress and wind speed (Jones et al. 1978). Ross and Jones (1978) reported several aircraft experiments in which a scatterometer was flown directly downwind off the U.S. East Coast in moderately strong offshore wind conditions. The  $\sigma_0$  measurements were plotted versus fetch (Figure 12), beginning within 1 km of the shoreline. There was found to be no fetch dependence at least to 40 km offshore, which suggests that the surface stress, the friction velocity, and the surface wind speed adjusted quickly to the surface roughness.

71. The above studies all suggest strongly that the adjustment scale for near-surface wind speed in strong-wind, near-neutral conditions is a few kilometers at most, which is a small distance compared to the dimensions of lakes Okeechobee and Pontchartrain. A dependence of wind speed on fetch may therefore be omitted. The situation with regard to wind direction is not as clear. There are no comparable field data for the adjustment of PBL wind direction across a roughness discontinuity. The available theories suggest a much longer adjustment scale. In the adopted scheme, no fetch-dependent adjustment for wind direction is incorporated; however, the case is allowed whereby the wind speed adjusts to the lake roughness immediately while throughout the lake, the wind direction is computed in accordance with the roughness of the terrain surrounding the lake (subroutine BREEZE/SPECIAL). This is correct if the adjustment scale for wind direction in the PBL is larger than the lake dimension. Field measurements are required to test this hypothesis.

72. Apart from fetch considerations, the surface winds over a large lake in hurricanes might be different from equivalent over-ocean winds because of differences in the surface roughness between a lake and

open sea. That is, the drag coefficient over a lake might be different than the drag coefficient over open sea, particularly for uniformly shallow lakes in which the surface wave structure can be expected to differ significantly for a given wind from deep-water surface waves. Unfortunately, the correct drag formulation for a shallow lake is much less certain than for the sea. Whitaker, Reid and Vastano (1973) have inferred a drag law for Lake Okeechobee which is quite different from the form proposed by Garratt. The scheme adopted below therefore includes the allowance for the specification of over-lake winds relative to an altered lake specification of surface roughness.

#### The Transformation Model

73. In this section a simple model is proposed to derive the surface wind and stress over terrain of arbitrary roughness from the numerical wind-field solution computed exclusively from the revised over-water drag formulation. The approach is to employ equilibrium PBL theory to relate the over-water integrated PBL wind to the flow at the top of the PBL and then to employ a consistent equilibrium model to compute the surface wind stress from the wind at the top of the PBL for terrain (including lakes) of arbitrarily specified roughness. The model proposed assumes that the PBL over land or inland lakes in a hurricane is neutrally stratified.

74. The transformations are derived quite simply from consideration of the alternate forms of the similarity PBL theory adapted in this study. To parameterize the surface drag in the numerical vortex model we applied equations 10a and 10b which relate the stress to the integrated PBL wind. Alternatively (Arya, 1977), the surface drag may be referenced to the wind at the top of the PBL  $u_h$ ,  $v_h$ :

$$\frac{ku}{u_*} = - (\ln \hat{z}_0 + A) \quad (20a)$$

$$\frac{kv}{u_*} = - B \operatorname{sign} f \quad (20b)$$

or to the surface geostrophic wind components

$$\frac{ku}{u_*} = - (\ln \hat{z}_o + A_o)$$

$$\frac{kv}{u_*} = - B_o \operatorname{sign} f$$

As noted by Arya,  $A_o$  and  $B_o$  may be expected to differ from  $A$  and  $B$  due to the presence of baroclinicity and also in very low latitudes where winds are strongly geostrophic. In hurricanes, baroclinicity in the PBL may be ignored, but the flow at the top of the PBL is more nearly in gradient balance. The effects of curvature on  $A$ ,  $B$  have not been studied, but the success achieved with the similarity PBL theory in the over water case suggests that such effects are not large. In this model, differences between  $A$ ,  $B$  and  $A_o$ ,  $B_o$  are ignored.

75. The relationship between  $A_o$ ,  $B_o$  and  $A_m$ ,  $B_m$  is given by Arya (1977) as derived from the equations of mean motion for a barotropic atmosphere in which the momentum flux is assumed to vanish at  $z = h$  :

$$A_m = A_o \quad (21a)$$

$$B_m = B_o - k (fh/u_*)^{-1} \quad (21b)$$

For the restricted case of  $fh/u_* = 1$ , which we have adopted only for the purposes of attaining a workable parameterization, it can be shown simply from equations 10, 20 and 21 that the flow at the top of the PBL,  $z = h$ , is related to the vertically integrated flow through

$$u_h = u \quad (22a)$$

$$v_h = v - u_* \quad (22b)$$

In the coordinate system adopted (see Figure 13),  $v_m$  is negative,  $u_*$  is always positive so the wind speed at the top of the PBL is always larger than and turned clockwise (in the Northern Hemisphere) with respect to the mean layer wind  $\vec{V}$ .

76. Given the wind at the level  $h$ , the consistent similarity theory defined by equations 20a and 20b may be solved for the surface stress and surface layer wind in a neutral PBL over terrain of specified roughness  $z_0$ . For a neutrally stratified PBL,  $A_0$  and  $B_0$  are reduced simply to constants (1.39,  $1.95 + k$ , respectively). If  $z_0$  is a constant, as say over a homogeneous land surface,  $u_*$  may be obtained directly from equations 20a and 20b. However since over a lake, the roughness probably depends on  $u_*$ ,  $z_0$  may in general be prescribed in terms of  $u_*$  using the general form proposed by Cardone (1969).

$$z_0 = C_1 u_*^{-1} + C_2 u_*^2 + C_3 \quad (23)$$

where  $C_1$ ,  $C_2$ ,  $C_3$  are constants to be chosen to impose a desired drag law. (For example, for a Charnock law,  $C_1 = C_3 = 0$ ,  $C_2 = a/g$ ; for land,  $C_1 = C_2 = 0$ ,  $C_3$  is the roughness parameter for the terrain type).

77. The implementation of the above model in the specification of hurricane surface wind fields over terrain of arbitrary roughness or lakes is coded as subroutine UPDOWN, which allows for the calculation of transformations for up to six terrain categories, for each of which roughness constants  $C_1$ ,  $C_2$ ,  $C_3$  have been specified. The procedure followed is:

- a. Given the integrated boundary layer wind  $u$ ,  $v$  and the conditions of a given hurricane over water ( $h$ ,  $\theta_a - \theta_0$ ,  $f$ ,  $k$ ,  $a$ ) compute  $u_*$  from the revised over water similarity parameterization.
- b. From equations 22a, 22b, compute the wind speed and direction at level  $h$ , the top of the PBL.
- c. For each terrain roughness, specified in terms of equation 23, use the neutral similarity model (20a, 20b) to compute the friction velocity appropriate to the terrain roughness,  $u_{*t}$ , the ratio  $u_{*t}/|\vec{v}|$ , and the angle between the surface stress and the integrated wind. The ratio and the turning angle are computed for  $|\vec{v}| = 0.8(.8)80.0$  m/sec, for each roughness category and stored for use in the specification of surface winds in a given simulation over a grid covering different terrain types.

78. The overall behavior of the transformations is exemplified in Figure 14, which shows the ratio of surface wind speeds at 20 meters (over-land + over-sea) and the difference between the over-land and over-sea inflow angle for two terrain roughnesses: .04m, .32m. For comparison, Figure 14 shows the results for the wind-speed ratio derived from numerical mixed-terrain and over-water solutions for a symmetric stationary vortex (radius and pressure drop as in Camille). To arrive at the indicated quantities, the surface wind speed and direction along a radial extending north of the eye over land in the mixed terrain case was referenced to the (symmetrical) solution along the radial for open ocean. It should be recalled that in the mixed-terrain solution, the wind field over land looked quite reasonable. Apparently, that solution can be retrieved quite simply from the over-water solution using the equilibrium model described above. It is also interesting to note that the form of the dependence shown in Figure 14 conforms quite closely to the empirical wind-speed ratio derived from measurements in hurricanes in and around Lake Okeechobee (U.S. Weather Bureau, Hydrometeorological Section, 1954).

#### Results for Test Storms

79. In this section, the results of calculations of the entire history of the surface wind field in selected hurricanes are checked at measurement locations at which representative surface wind measurements are available. Indeed, the storms were chosen because over-water wind measurements and, in some cases, representative over-land wind measurements were available, and because extensive analyses of storm pressure and track characteristics had been performed in previous studies. Results are presented for hurricanes Camille, 1969; Betsy, 1965; Delia, 1973; Belle, 1976; Anita, 1977; and the Lake Okeechobee (LO) storm. of 1949. Comparisons against over water winds are made in Camille, Delia, Belle, and Anita. A limited evaluation of over-land and over-lake winds in Camille and the LO storm is made. Finally, sample results for Betsy are presented without comparison against measurements, as no represen-

tative measured winds over land or lake are available in that storm.

#### Camille

80. Hurricane Camille played a crucial role in the calibration of the method developed by Cardone et al. (1976), because at that time, the wind trace in Camille from Rig 50 was deemed to be the only extant representative over-water wind trace in an intense hurricane. Indeed, the empirical law developed by Cardone et al. to scale integrated boundary layer winds to standard anemometer height was based largely on that wind trace. In the present scheme, the assignable model parameters are physical quantities related to fundamental properties of the planetary boundary layer model (PBL height, stability, roughness parameter formulation). Also, the method provides the wind speed as might be measured at any height in the constant-stress surface boundary layer, which in hurricanes may extend to a height of at least 50 meters.

81. The comparison of measured and modelled wind speed at Rig 50 at measurement height in Camille is shown in Figure 15. The agreement is at least as good as that achieved by Cardone et al. (1976). Results for two boundary-layer heights are shown. The numerical solutions differ little, except near the eye, where the lower height ( $h = 500$  m) matches the peak measured wind better. This may reflect the fact that near the eye of intense hurricanes, the PBL height lowers slightly. Since the solutions differ little outside the eye, it is perhaps prudent to use a PBL height of 500 m in simulations of strong storms in the Gulf of Mexico.

82. To generate modelled surface-wind time histories at land sites requires knowledge of the roughness parameter representative of the terrain. Typical  $z_0$  values for various terrain types have been given by several workers (e.g. Figure 16 after ESDU 72026, 1972). The roughness parameter is very sensitive to the terrain type within a few kilometers of a given measurement site, and therefore often varies significantly with wind direction at a site. For the sites at which winds have been measured in the selected hurricanes, the roughness parameter has not been determined experimentally. We therefore depict modelled time histories covering a reasonable range of roughness parameters at

land sites. Also, with the exception of the lakefront comparison in Camille, comparisons are restricted to only those measurement times at which the anemometer was recorded on strip chart, from which 30-minute-average wind speed and direction could be extracted.

83. Figure 17 compares measured and modelled winds at Keesler Air Force Base, Biloxi, Mississippi. The anemometer at Keesler is mounted 16 feet above the runway surrounded by fairly level terrain but with the coast less than a block away to the southeast. The anemometer measured winds up to the time of eye landfall at 0000 CDT. Measured wind directions support an over land trajectory up to about 2300 CDT and an over water trajectory thereafter. The modelled wind history for  $z_0$  of .16 m, compares favorably with the measurements up to 2300 CDT; after which the measurements agree better with an equivalent over-water history (also shown in Figure 17). Modelled wind direction is generally within  $\pm 20^\circ$  of that measured.

84. The anemometer at Burwood CGS, Southwest Pass, is in a complex environment, with at least some influence of land to be expected, especially for wind directions between northwest and northeast. The wind comparisons in Figures 18 and 19 show better agreement in wind direction for an assumed over-water exposure. For wind speed, however, over-land histories agree better, but the effect of a shift in wind direction to an off-water direction (1800-2000 CDT) can clearly be seen in the measured wind speed.

85. Since fetch effects are not built into the present model, the Burwood and Keesler comparisons present worst-case examples of limitations of the present scheme. Nevertheless, the results appear accurate enough for specification of over-land winds in the coastal zone.

86. An important function of the present model is to specify winds over Lake Pontchartrain in hurricanes. The scheme adopted includes the provision for the specification of winds over the lake different from winds over open water in two ways. First, surface wind speeds may be computed relative to a  $z_0$  specified for lakes which differs from the Charnock law assumed over water. Second, the program accommodates the condition, discussed earlier, whereby the surface layer

wind speed and stress are assumed to adjust to the lake roughness on length scales small compared to the lake width, but the PBL wind direction is governed by the roughness of that terrain upwind of the lake.

87. There is virtually no data on the roughness properties of Lake Pontchartrain. Whitaker, Reid and Vastano (1973) have deduced a drag law for Lake Okeechobee which differs substantially in level and wind dependence from equation 9. Their form was fit to equation 23, and was used to provide test wind histories at measurement sites around and in the lakes in the test simulations. (The form of equation 23 does not provide a particularly accurate fit to the odd form proposed by Whitaker et al., but is within about 20% over the range 10-50 m/s.)

88. The model was used to specify surface wind speed at the location and measurement height of Lakefront Airport, New Orleans, in Camille, for both water and lake roughness laws (Figure 20). The wind direction was computed for the water  $z_o$  and for the special case described above, assuming an upwind terrain roughness of 1 m. Observed 1-minute hourly surface winds from the airport station are also plotted. Quantitative comparison of wind speeds is not warranted because the measurement is poorly averaged and the measurement site may not be representative of over-lake conditions. The reported wind directions should be more representative; the comparisons there suggest that there may indeed be a larger inflow angle for winds over Lake Pontchartrain in hurricanes than returned by the nominal over-water or over-lake transformation. Higher quality measurements in the lake in storm conditions are required to verify this possibility.

89. Figure 20 also shows the wind speed over downtown New Orleans, at 85 feet, calculated with a  $z_o$  of 1 m. Time-averaged measured winds are not available for comparison.

#### Delia

90. Forristall et al. (1977), using the method of Cardone et al. (1976), ran a simulation of hurricane Delia in order to produce a wind field as accurate as possible, permitting the simulation of the surface wind and current at Buccaneer tower, offshore of Galveston, Texas. Thus

the storm parameters and track were adjusted, within their range of uncertainty, to produce best agreement between measured and modelled winds at the tower. Those storm parameters and track were used, without alteration, as input data to the present scheme and the wind field was computed. The boundary layer height was specified as 500 m.

91. The comparison of measured and modelled wind speed and direction in this storm is shown in Figures 21 and 22. The agreement in wind speed and direction is generally excellent, except for about an 10-20° excess of inflow in the modelled wind directions. Slight alteration of the input parameters and track of this poorly organized highly erratic storm would probably have provided even better agreement.

#### Belle

92. The wind and wave fields in hurricane Belle have been modelled by Cardone and Ross (1979), using both the methods of Cardone et al. and simpler parametric schemes. Belle moved rapidly up along the east coast. As it did so, the central pressure rose sharply and the eye diameter increased. This storm, therefore, provides a critical test of the present method, which simulates such time changes in terms of a series of steady-state numerical solutions.

93. The center of Belle passed directly over two NOAA data buoys, one EB15 located offshore South Carolina; the other, EB41, located east of New Jersey. As for the Delia simulation, the storm input parameters and storm track determined by Cardone and Ross (1979) was used without alteration to drive the present model. Four steady-state solutions were utilized to attempt to accomodate the rapid changes in storm intensity, shape and speed.

94. The modelled and measured time histories are compared at sensor height at EB15 and EB41, in Figures 23, and 24. Agreement is generally excellent. The departure between modelled and measured wind direction early in the EB41 history is related to the presence of a frontal trough of low pressure which was located off the New Jersey coast and which distorted the pattern of isobars in the forward quadrants of the storm.

#### Lake Okeechobee Storm of 1949

95. The Lake Okeechobee storm of 1949 has been the subject of much past study. It is particularly important because winds were measured over the lake with calibrated anemometers mounted at 10 meters height. The wind data had been reduced to 10-minute averages by the Hydrometeorological Section, U.S. Weather Bureau. For the comparisons shown in this study, the data were reduced further to 30-minute averages, to be more consistent with the implied averaging interval of modelled winds.

96. Storm input parameters were specified as objectively as possible from the data published on this storm. While there is considerable storm data from the lake stations, there remains some uncertainty in the precise storm track; particularly east of the Florida coast and as the storm recurved northwest of the lake. There is also some uncertainty in the filling rate. In our simulation, no filling is applied until just after the eye of the hurricane has crossed the lake. Storm pressure input parameters ( $p_0 = 954$  mb,  $R = 22$  n.mi.) prior to the filling stage were taken directly from Graham and Hudson (1960), while the steering flow was estimated from historical Northern Hemisphere Surface Analyses. The four-parameter ( $\Delta p_c$ ,  $R$ ,  $V_c$ ,  $V_g$ ) pressure initialization scheme was adopted since the storm appeared to translate in the direction of the steering flow. There did not appear to be sufficient data in this storm to estimate  $\Delta p$  and  $R$  by quadrant. Storm track was taken from published track charts, though it is not clear whether the track relates to the pressure center or the center of surface wind circulation.

97. Surface winds were measured reliably at two locations within the lake, stations 14 and 16, shown in Figure 25. The same figure shows the path of the storm schematically. Surface winds were computed for two roughness categories: (1) the over-water Charnock law; (2) the roughness specification, equation 23, with constants consistent with the Lake Okeechobee drag law of Whitaker et al. (1973).

98. Measured and modelled winds at the lake stations are shown in Figures 27 and 28. The histories cover the period from about 8 hours before the occurrence of maximum wind on the lake, at which time the

storm center was between Nassau, Bahamas and West Palm Beach, Florida, and the 8-10 hour period after the occurrence of maximum winds, when the storm was filling rapidly and recurving northward over central Florida.

99. The wind speed comparisons show that in general the lake roughness histories compare better with the measurements than the over-water specification. There is one serious discrepancy between modelled and measured wind speeds: this occurs over a two-hour period just after the occurrence of maximum winds, when the modelled winds show a significant drop before rejoining the measurement history. The measurements also show a drop but to a much lesser degree. The double maximum suggests that the stations "entered" the eye briefly on the western side of the hurricane and that this effect was accentuated in the simulation perhaps because the offset between the wind circulation center and the pressure center was larger in the numerical model than actually occurred. Another possibility is that the pressure field in this storm was simply more complex than could be described by three parameters.

100. The modelled wind directions agree well with measurements at Station 14 in the forward quadrant of the storm and at Station 16 in the rear quadrant of the storm; otherwise systematic departures of  $20-40^{\circ}$  are evident. The sense of the discrepancy is that there is less inflow than modelled in the forward quadrants and more inflow than modelled in the rear quadrants, over the lake. Frictional effects associated with the presence of roughness boundaries are not as likely to be the cause of systematic effects of this nature as are differences between the actual large scale pressure field in this storm and the one simply modelled.

#### Betsy

101. Hurricane Betsy served to test the complete history program, including the specification of surface wind fields over a rectangular high resolution grid, each hour, throughout a 24-hour history run. Input data and sample output fields are included in Appendix C.

102. As a part of the test, surface wind histories were calculated at several locations around Lake Pontchartrain. Figure 28 displays sample wind histories for Lake Pontchartrain at Lakefront (open water

roughness), Lake Pontchartrain at Lakefront (Whitaker et al. lake roughness), New Orleans Moisant Airport ( $z_o = .16$  m) and the U.S. Weather Bureau Office city office, New Orleans ( $z_o = 1$  m). No attempt is made to compare these calculated histories to measurements.

Anita

103. Hurricane Anita was one of the most intense hurricanes of historical record to cross the Gulf of Mexico. The storm formed in the east central Gulf of Mexico on August 28, 1977 and moved west-southwestward into Mexico on September 02, 1977, sparing populated areas from its fury. Anita passed about 50 n.mi. north of NOAA buoy EB04 on the 30th as a poorly organized but deepening tropical storm and about 10 n.mi. south of EB71 early on September 01. As the storm moved past EB71, it was much better organized and undergoing explosive development. The general path of the storm and the time history of central pressure are shown in Figure 29.

104. Four steady-state solutions were generated for Anita, corresponding to the storms' parameters at the times indicated in Figure 29. The radius to maximum wind in Anita contracted from 30 n.mi at the time of the first solution to 15 n.mi. at the last solution. The four solutions were used to interpolate winds in space to the buoy locations over a 48 hour period.

105. The modelled wind series at the locations of the buoys are compared to the winds measured at 20 meter height in Figures 30 and 31. At EB04, the agreement is surprisingly good considering the poorly organized nature of the storm during the period shown. Agreement is generally very good also at EB71. As suggested by the Belle test, the steady-state model provides reasonably good simulations even when applied to storms undergoing rapid changes in intensity and structure.

### PART III: THE COMPUTER PROGRAM

#### Program Description

. The program task which produces tropical storm wind histories at specified locations is divided into two main programs. The first, SNAP, produces snapshot wind fields on a nested grid and writes them onto an output data file from which the second program, HIST, obtains nested-grid wind fields for each hour of the storm's history, using linear interpolation if necessary. HIST then gets and prints the wind history at each measurement station specified, and, if requested, writes the wind history for a different grid onto an output data file.

. The programs were written in Fortran V and were run on a UNIVAC 1108 computer. Each snapshot takes approximately 6 minutes of computer time, and execution of program HIST usually takes less than 3 minutes. Some mass storage is required, the amount varying with the number of snapshots, the number of interpolations between snapshots, the length of the history, and whether or not winds are to be interpolated to an output grid; 250,000 words is sufficient for most storms. Substitution of tapes for mass storage is possible, but efficiency would be decreased.

- . Program SNAP is composed of:

MAIN

- Together with its subprograms, produces one or more snapshot wind fields on a nested grid according to card input specifications. It prints the computed winds and writes them onto an output file, and it optionally prints corresponding pressure fields and initial guess winds. MAIN itself reads all input cards, calls subroutine CCROSS which sets up tables, calls BLOWUQ which controls computation of the winds, and writes the final snapshot wind fields onto the output file.

SUBROUTINE AANGEL -

- Converts grid components (UN , VN) of integrated wind to speed (VTN) and direction (ANG) for all points of the 21 × 21 × 5 wind grid. In addition, if

the switch variable I20 ≠ 0 , AANGEL reduces the speed (VTN) to a height of 19.5 m; TWIST, the necessary correction to ANG, is computed by interpolating the array TURN;  $u_*$  is computed by interpolating the array UXV containing  $u_*/V_m$ ; anemometer wind is computed from  $u_*$  (called UXX in the code) by the usual logarithmic profile.

Arguments: input, typing implicit

I20: Flag, if non-zero winds are reduced to 19.5 meters.

- SUBROUTINE ABCC - Computes Arya's  $A_m$ ,  $B_m$ ,  $C_m$  (called AM, BM, CM in the FORTRAN program) and UV, the square of the integrated wind speed, all as functions of the friction velocity  $u_*$ . This is passed in common from CCROSS at location UX(K123), where K123 = 1, 2, or 3 at various stages of the iteration. The code is slightly more general than needed in the hurricane model, catering for unstable, neutral, and stable wind profiles (indexed by the sign of HL). In computing  $A_m$  and  $B_m$ , the ratio  $f_h/u_*$  is taken equal to unity.
- SUBROUTINE BLOWUQ - Controls computation and printing of all output on the nested grid. It is called once for each snapshot wind field.
- SUBROUTINE CCROSS - Computes the upwind and crosswind drag coefficients, the ratio  $u_*/V_m$ , and the angle between surface wind and integrated wind, for all  $V_m = 0.8(0.8)80.0$ . Values for intermediate  $V_m$  are linearly interpolated when required (lines 72-83 of COMQUT; lines 27-37 of AANGEL). The computation implements Arya's theory. The numerical method is an initial guess at  $u_*$ , followed by an iterative series of corrections by inverse interpolation. The iterations proper, and the computation of Arya's  $A_m$ ,  $B_m$ ,  $C_m$ , take place in subroutine ABCC.
- SUBROUTINE COMQUT - Solves equations which determine the final wind fields on the nested grid. For each snapshot, COMQUT is called as many times as specified in input NAME3. At each calling a grid level is specified, and computation is done on that level

and all finer-meshed levels only, so that wind computation on the innermost nest only is computed NM times.

Arguments: input, typing implicit

LEVEL: Input  $1 \leq LEVEL \leq 5$ .

Grid distance is doubled at each increase of LEVEL. Computation is done on all grid levels  $\leq LEVEL$ .

SUBROUTINE GRAD

- Computes the radial and tangential gradients  $\partial P / \partial r$  and  $r^{-1} \partial P / \partial \theta$  of an exponential pressure field and converts them to rectangular gradients  $\partial P / \partial x$  and  $\partial P / \partial y$ .

Mathematical Method:

1. Compute polar coordinates  $(r, \cos \theta, \sin \theta)$  of the  $21 \times 21 \times 5$  grid points.
2. Convert direction of track to radians.
3. Convert forward speed to meters per second.
4. Compute x- and y- components of forward speed. The method here divides into two cases: circularly symmetric pressure field ( $JA(6) = 0$ ) and quadrantal pressure field ( $JA(6) \neq 0$ ).

A: Circularly symmetric pressure field.

The governing equation

$$P = P_0 + \Delta P \exp(-R/r) \quad (*)$$

yields on differentiation

$$\partial P / \partial r = \Delta P R r^{-2} \exp(-R/r).$$

5. Convert R to kilometers.
6. Compute P.
7. Compute  $\partial P / \partial r$ ,  $\partial P / \partial x$ ,  $\partial P / \partial y$ .

B: Quadrantal pressure field.

$\Delta P$ , prescribed in four quadrants, is expanded as the trigonometric polynomial

$$a_0 + a_1 \cos \theta + a_2 \sin \theta + [a_3 \cos 2\theta]$$

and then smoothed by removing the

bracketed term.  $R$  is similarly expanded and smoothed. Substituting these trigonometric polynomials in (\*) yields

$$P = P_0 + (a_0 + a_1 \cos \theta \\ + a_2 \sin \theta) \exp (-[b_0 + b_1 \cos \theta \\ + b_2 \sin \theta]/r)$$

and the radial and tangential gradients

$$\frac{\partial P}{\partial r} = (a_0 + a_1 \cos \theta \\ + a_2 \sin \theta)(b_0 + b_1 \cos \theta \\ + b_2 \sin \theta) r^{-2} \times \exp (-[b_0 \\ + b_1 \cos \theta + b_2 \sin \theta]/r)$$

$$r^{-1} \frac{\partial P}{\partial \theta} = [r^{-1}(-a_1 \sin \theta \\ + a_2 \cos \theta) + r^2(b_1 \sin \theta \\ - b_2 \cos \theta)] \times \exp (-[b_0 \\ + b_1 \cos \theta + b_2 \sin \theta]/r)$$

8. Convert  $R$  to kilometers in each quadrant.
9. Compute  $\Delta P$  to millibars in each quadrant.
10. Compute the coefficients in trigonometric polynomials.
11. Compute  $\frac{\partial P}{\partial r}$  and  $r^{-1} \frac{\partial P}{\partial \theta}$ .
12. Compute  $\frac{\partial P}{\partial x}$  and  $\frac{\partial P}{\partial y}$ .

Arguments: none

Variables (not in common):

AD             $\frac{\partial P}{\partial r}$ ,  $r^{-1} \frac{\partial P}{\partial \theta}$ ,  
 $\frac{\partial P}{\partial x}$ ,  $\frac{\partial P}{\partial y}$

AE,AF,AG,AH    Temporary storage

AP, CR, CD    Coefficients in trigonometric polynomials

AT,CT        Direction cosines of motion of storm

BA,BC        JA, floated and converted to metric units

BP            Temporary storage

DEG          Radian measure of 1 deg.

IC, ID, IE, Do-loop indexes  
IG, IH

LU              Logical unit number of  
                  standard print unit  
S45              Circular sine of 45°

**SUBROUTINE OUTBY1** - Sets outer boundary winds for the next time level. It is called once for each cycle of wind computation and grid level if the grid level is not the outermost computed at that time.

Arguments: typing implicit

NEST: Input - grid level at which boundary to be set.

**SUBROUTINE OUTBY2** - Sets outer boundary winds for the same time level. It is called once for each cycle of wind computation for the coarsest meshed grid level being computed at that time unless that grid level is the outermost in the entire grid.

Arguments: typing implicit

NEST: Input - grid level at which boundary to be set.

**SUBROUTINE OUTFLO** - Operates on the entire field of  $21 \times 21 \times 5$  wind vectors, rotating every vector clockwise (in the northern hemisphere) by  $8^\circ$ . Extensive numerical experiment with earlier versions of the hurricane wind model has indicated that the finite difference scheme used leads to excessive inflow; the  $8^\circ$  rotation approximately removes that bias.

**SUBROUTINE OUTQUT** - Controls printing of winds on nested grid.

Arguments: input, typing implicit

I20: Passed on to subroutine AANGEL.  
If I20  $\neq$  0, wind speed will be adjusted to 19.5 meters.

NAME: 4-character name of storm  
IDENT: 4-character identification of type of field (INIT or SNAP)  
NSEQ: Snapshot sequence number.

**SUBROUTINE PXYM** - Receives the pressure gradients computed in GRAD, divides them by the density, and rearranges them in the order demanded by BLOWUQ. It then computes the gradient wind from the radial pres-

sure gradient to provide initial values for COMQUT.

Mathematical Method:

The gradient wind is given by Hess (1959, p. 183) as

$$C = -\frac{1}{2}fr + [(\frac{1}{2}fr)^2 + \rho^{-1}r \frac{\partial p}{\partial r}]^{\frac{1}{2}} (*)$$

For large  $r$ , (\*) expresses  $C$  as the difference of two nearly equal terms, and so it is advisable to compute the equivalent expression

$$C = \frac{\rho^{-1}r \frac{\partial p}{\partial r}}{\frac{1}{2}fr + [(\frac{1}{2}fr)^2 + \rho^{-1}r \frac{\partial p}{\partial r}]^{\frac{1}{2}}}$$

where  $f$  is the Coriolis acceleration and  $\rho$  is the density of the atmosphere.

1. Coriolis, computed in SNAP, is passed in  $C_1$ . Note that the latitude is taken as a constant throughout the cyclone; this approximation is inappropriate between  $\pm 5^\circ$  of latitude.
2. Divide  $\frac{\partial p}{\partial x}$  and  $\frac{\partial p}{\partial y}$  by  $\rho$ .  $1.15 \times 10^{-3}$  is the density;  $1 \times 10^{-4}$  is a conversion factor from mb/km to newtons/m<sup>3</sup>.
3. Compute gradient wind. BC contains  $r$  in km, and the factor 1000 converts  $r$  to meters.
4. Resolve gradient wind into  $x$ - and  $y$ - components.
5. If steering flow is used, add  $(W_c \times F)$  to the vector  $\rho^{-1}\Delta P$ .

Arguments: none

Variables (not in common):

AG,AH,AI	Temporary storage
I	Index variable for X
J, MJ	Index variable for Y
NEST	Index variable for nest (counted from inside out)
FADE	Attenuation factor for steering flow

- SUBROUTINE SHORE** - Sets the land/sea table for the nested grid. It is currently set to 'sea' throughout.
- SUBROUTINE TVEL** - Prints the contents of arrays VTN and ANG on the grid level indicated as well as on the next coarser grid level.  
Arguments: Input, typing implicit
- VTN: The top number in each pair of numbers printed is taken from this array. It is printed with the decimal moved one place to the right.
- ANG: The bottom number in each pair of numbers printed is taken from this array.
- NBASE: 4-character heading information
- IDENT: 4-character heading information
- KSEQ: Header information, 4-digit integer number.
- LV: Finer-meshed (lower numbered) grid level to be printed at this call to TVEL.
- I20: Flag, controlling printing of the legends "reduced" and "not reduced".

Note heading:

[NBASE] [IDENT] [KSEQ] LEVEL LV+1..

- Program HIST is composed of:

MAIN

- Produces winds at specified measurement stations for each hour of a storm's history. If requested, it will also interpolate winds to a grid. After card input has been read and tables have been set up, it writes all unique nested-grid winds (i.e. all snapshots plus all interpolated fields) on a temporary file in the order needed. It next loops through the list of stations and, using the temporary file just written, finds and prints the winds for each station throughout the storm's history. Then, if requested, it interpolates winds to another grid for each hour of the storm's history, prints the winds on the grid for any hours indicated, and writes the fields onto an output file.

SUBROUTINE ABCCC

- (called by UXV) Operates exactly the same algorithm as ABCC (called by CCROSS). The only difference in coding is that ABBCCC references the common block /C57/, defined in program HIST.

SUBROUTINE BREEZE

- Given LA0, L00, ROT, LA1, L01, DX, STHT, and LANSEA in common block D1, BREEZE determines the wind at point LA1, L01 at height STHT on the nested grid wind field in array XX in common block D2 and returns the wind data in W1, TH1, D, AL, and UST of common block D1.

Mathematical Method:

Let  $a$ ,  $b$ ,  $c$  be the sides of a spherical triangle, and  $\alpha$ ,  $\beta$ ,  $\gamma$  the angles opposite; define  $s = \frac{1}{2}(a + b + c)$ . Then

$$\begin{aligned} \text{hav } c &= \text{hav } (a - b) \\ &\quad + \sin a \sin b \text{ hav } \gamma; \end{aligned}$$

$$\tan^2 \frac{1}{2}\alpha = [\sin(s - b) \sin(s - c)] / [\sin(s - a) \sin s]_{(*)}$$

1. Compute distance and bearing. If  $c$  is very small,  $(s - a)$  and  $(s - b)$  are nearly zero, and eq. (\*) is unsuitable for computation; the bearing can then be computed without sen-

sible error by solving a plane triangle.

2. Reduce bearing to rectangular grid.
3. Reduce distance to kilometers.
4. Compute rectangular coordinates.
5. Search for the smallest rectangular grid in whose interior the point lies. If point lies without fifth nest, no wind has been computed; set wind to zero.
6. Interpolate components of wind, using bivariate linear interpolation:

$$\phi(x + f_1 \Delta x, y + f_2 \Delta y)$$

$$= (1 - f_1)(1 - f_2) \phi(x, y)$$

$$+ (1 - f_1) f_2 \phi(x, y + \Delta y)$$

$$+ f_1(1 - f_2) \phi(x + \Delta x, y)$$

$$+ f_1 f_2 \phi(x + \Delta x, y + \Delta y)$$

7. Compute wind speed and reduce to anemometer height. The factor 3600/1852 converts from m/sec to knots; the MIN function assures that the anemometer wind is never greater than the integrated wind.
8. Compute wind direction and reduce to true south.

Arguments: none

Variables: All variables except temporary storage are annotated in the program listing.

SUBROUTINE INVJD

- Is the inverse of the function JULIAN. Given the Julian date, it returns the month, day, and year.

Arguments:

J: Input-Julian date, type integer

M: Output-Month, type integer

D: Output-Day, type integer

Y: Output-Year, type integer

FUNCTION JULIAN

- Returns the Julian date.

Arguments:

MO: Month, type integer  
DA: Day, type integer  
YR: Year, type integer

SUBROUTINE PRLAKE - Is a grid-dependent subroutine and must be changed or replaced if the output grid is changed. It prints winds for each grid point, with speed on top, direction in the middle, and terrain code on the bottom.

Arguments: Input, typing implicit

NBASE: 4-character name of storm;  
KHR: Integer sequence number of hour of storm;  
ISTART: First hour of storm, corresponds to KHR = 1 . Format is YYMMDDHH;  
IZONE: 3-character time zone of ISTART;  
WIND: As on output file 20;  
LSTAB: Terrain code table for output grid;  
MAXI: Number of longitudes in output grid;  
MAXJ: Number of latitudes in output grid.

Note: The grid currently used uses unequally spaced meridians and parallels; printed map is distorted.

SUBROUTINE RDGRID - Is a grid-dependent subroutine and must be replaced or altered if the output grid is changed. Its function is to read latitude, longitude, and terrain code for each grid point and store them in arrays ZLA, ZLØ, and LSTAB respectively. It is called only if winds are to be interpolated to an output grid.

Arguments: typing implicit

ZLA: Output - latitudes in radians ordered south to north.  
ZLØ: Output - west longitudes in radians, ordered west to east.  
LSTAB: Output - terrain codes for all grid points; the first subscript increases eastward, the second increases northward.  
MAXI: Input - number of longitudes in grid.

MAXJ : Input - number of latitudes  
in grid

SUBROUTINE UPDOWN - Computes the ratio  $u_*/V_m$  and the angle between surface wind and integrated wind, all for  $V_m = 0.8(0.8)80.0$ , for terrains other than open ocean. The program consists of two parts: "UP" (lines 8-19) and "DOWN" (the rest of the code). UP computes UM and VTOP (components of wind at the top of the boundary layer), VW2 (squared wind speed at top), and TARN (tan of angle between integrated wind and wind at top). The computation in UP uses quantities computed in UXXV (open ocean) and is consistent with Arya's theory. The assumption is now made that wind at the top of the boundary layer in a hurricane does not "see" the terrain below, so that the surface wind over any terrain can be computed by working UP and then DOWN. DOWN follows a logic similar to UXXV : Arya's  $A_o$  and  $B_o$  are constants (neutral wind profile); the roughness length is computed as

$$Z_o^2 = AZ/u_* + BZu_* + CZ.$$

SUBROUTINE UXXV - Computes the ratio  $u_*/V_m$ , the angle between surface wind and integrated wind and the cosine and sine of this angle, all for  $V_m = 0.8(0.8)80.0$ . Values for intermediate  $V_m$  are linearly interpolated when required (line 73-82 of BREEZE). The computation implements Arya's theory. The numerical method is an initial guess at  $u_*$ , followed by an iterative series of corrections by inverse interpolation. The iteration proper, and the computation of Arya's  $A_m$ ,  $B_m$ ,  $C_m$ , take place in subroutine ABCCC. UXXV works the part of the algorithm of CCROSS (called from SNAP) pertaining to sea, i.e. LS = 2. The computations in UXXV are valid for open ocean only; all other terrains are treated in subroutine UPDOWN.

Remarks

. In all arrays dimensioned  $21 \times 21 \times N$ , where  $N$  is a multiple of 5, the 1st dimension increases eastward, the 2nd dimension increases northward, and the 3rd dimension, grid nest, increases with grid spacing. Grid spacing doubles with each increasing nest level.

. Logical unit numbers for the card reader and printer are 5 and 6, respectively, on the system under which these programs were run. These unit numbers are set by a DATA statement into variables LR and LP in programs MIST and SNAP, and the printer unit is set into LU in subroutines GRAD and TVEL .

. Equivalences between snapshot input parameters and array JA in COMMON block C2 of program SNAP:

<u>JA</u>	<u>NAME</u>	<u>TYPE</u>
1	ITRACK	I
2	EYELAT	R
3	EYLONG	R
4	DIREC	R
5	SPEED	R
6	IQUAD	I
7	EYPRES	R
8	RADIUS(1)	R
9	RADIUS(2)	R
10	RADIUS(3)	R
11	RADIUS(4)	R
12	PFAR(1)	R
13	PFAR(2)	R
14	PFAR(3)	R
15	PFAR(4)	R

**Explanation of program organization charts:**

Rectangular box - program element

Cut-off corner - punched card image

Diamond - printer

Barrel - mass storage

Rounded ends - program stop

Table 2

Files

<u>Description</u>	<u>Unit Number</u>	<u>Size</u>	<u>Program SNAP</u>	<u>Program HIST</u>	<u>Save</u>
Card Reader	5		Input	Input	
Printer	6		Output	Output	
Snapshot Wind Fields on Nested Grid	13	4436 Words each Record (Snapshot)	Output	Input	/
Ordered Unique Wind Fields on Nested Grid	10	4411 Words each Record (Wind Field)		Work	
Hourly Wind Fields on Output Grid	20	Grid-dependent, 3855 Words each Record (Hour) for Test Grid		Output	/

Records in files 10, 13, and 20 are written with Fortran unformatted WRITE statements. There are file marks after the last data records in output files 13 and 20.

Table 3  
Card Input

Description and order within programs of card input groups

<u>Program</u>	<u>Seq. Number</u>	<u>Name*(If Namelist)</u>	<u>Number/Remarks</u>	<u>Description</u>
SNAP	1	NAME1	1	Processing control
SNAP	2	NAME2	1	Parameters in roughness law (usually constant)
SNAP	3	NAME3	1 for each wind snapshot	Parameters describing wind field
HIST	1	NAME4	1	Storm identification, also grid parameters if winds are to be inter- polated to an output grid
HIST	2	NAME5	1	Terrain coefficients and number of types of terrain other than open ocean
HIST (RDGRID)	3		Only if grid conversion	Longitudes and latitudes of grid points. Card count, format and ordering of data must agree with subroutine RDGRID.
HIST (RDGRID)	4		Only if grid conversion	Code for type of terrain at each grid point. Card count, format, and ordering of data must agree with subroutine RDGRID.

\* Namelist is not a construction recognized by the current FORTRAN standard (ANSI X3.9-1978).  
Appendix A contains an account of Namelist as used in this program.

Table 3 (concluded)

<u>Program</u>	<u>Seq. Number</u>	<u>Name ( If Namelist )</u>	<u>Number/Remarks</u>	<u>Description</u>
HIST	5(or 3)	One card for each station where wind history is needed, terminated by end-of-file		Station location and height at which measurements taken.
HIST	6(or 4)	One card for each hour of storm history, terminated by end-of-file		Location of eye of storm, snapshot identification

Table 4

Program SNAP namelist input

<u>Namelist Name</u>	<u>Variable Name</u>	<u>Size in Words</u>	<u>Units</u>	<u>Type</u>	<u>Default Value</u>	<u>Description</u>
NAME1	IB	1		I		Switch variable: 0 suppresses printing of pressures and initial winds.
NZ		1		I		Number of snapshot wind fields to compute
NAME2	DTH	2	°K	R	0,-2.	1: Air-land temperature difference 2: Air-sea temperature difference
	HH	1	m	R	650.	Boundary layer height over water
	ZOLAND	1		R	.08	Roughness length over land
	GARR	1		R	.0144	Charnock's constant
	PTH	1	°K	R	300.	Potential temperature
	K35	1		R	.35	Karman's constant

Table 4 (continued)

Name	Variable Name	Size in Words	Units	Type	Default Value	Description
NAME3	SGW	1	m/sec	R		Surface geostrophic wind of ambient flow
AN1	AN1	1	deg	R		Direction of SGW, counterclockwise from snapshot x-axis
NAME	NAME	1		H		4-character name of storm
EYELAT	EYELAT	1	deg	R		North latitude of eye of storm
EYLONG	EYLONG	1	deg	R		West longitude of eye of storm; for reference only
DIREC	DIREC	1		R		Direction of track of storm, clockwise from north. See ITTRACK for units.
SPEED	SPEED	1	kn	R		Forward speed of storm
EYPRES	EYPRES	1	mb	R		Pressure at eye of storm
RADIUS	RADIUS	4	nm	R		Exponential pressure profile scale radius in 4 quadrants. If the pressure field is circularly symmetric (IQUAD = 0) input is required for the first quadrant only
PFAR	PFAR	4	mb	R		Far field pressure in 4 quadrants. If the pressure field is circularly symmetric (IQUAD = 0) input is required for first quadrant only.
NH	NH	1		I	800	Number of times to cycle wind computation in innermost grid nest
DX	DX	1	km	R	5.	Grid spacing of innermost nest
STR2	STR2	1	km	R	0.	Distance from axis to $\frac{1}{2}$ magnitude of SGW

(continued)

Table 4 (concluded)

<u>Namelist Name</u>	<u>Variable Name</u>	<u>Size in Words</u>	<u>Units</u>	<u>Type</u>	<u>Default Value</u>	<u>Description</u>
NAME3	ITRACK	1		I	0	If 0, DIREC in degrees; if 1, DIREC in points of 11.25 degrees
	IQUAD	1		I	0	Indicator for quadrants of pressure field: 0, circularly symmetric pressure field; 1, 1st quadrant is right front; 2, 1st quadrant if forward

**Table 5**  
**Program HIST Namelist Input**

<u>Namelist Name</u>	<u>Variable Name</u>	<u>Size in Words</u>	<u>Type</u>	<u>Description</u>
NAME4	NBASE	1	H	4-character name of storm - must be same as 1st SNAP namelist NAME3, item NAME
	ISTART	1	I	Starting time of storm, format YYMMDDHH
	IZONE	1	H	3-character time zone of ISTART
	ICNVRT	1	I	Flag: if non-zero, winds will be interpolated to an output grid, and grid data must be input
	NPRT	1	I	Interval in hours at which to print winds on output grid. If zero, winds will be printed only for hours so flagged in the history table
NAME5	LAKE	1	I	Number of types of terrain in addition to open ocean. $0 \leq LAKE \leq 5$
	ZCOEFF	15	R	3 coefficients in formula relating $Z_o$ to $U^*$ for all terrains except open ocean. Dimensioned (3,5).

Table 6  
Program HIST Fixed Format Card Input

- 1) Station location and data: One card for each measurement station, format (5I4,F6.1,1X,I3)
  1. Degrees of north latitude
  2. Minutes of latitude
  3. Degrees of west longitude
  4. Minutes of longitude
  5. Terrain code, same codes as grid terrain code table
  6. Station height in meters
  7. Station number (for identification only)
- 2) History: One card for each hour, format (6I4,F8.4,2I4)
  1. Degrees of north latitude of eye of storm
  2. Minutes of latitude of eye
  3. Degrees of west longitude of eye
  4. Minutes of longitude of eye
  5. Sequence number of 1st snapshot wind field to be used for this hour
  6. Sequence number of 2nd snapshot wind field to be used for this hour (blank if no interpolation this hour)
  7. Interpolation distance between 1st and 2nd snapshots (blank if (6) is blank)
  8. Clockwise rotation of snapshot in degrees
  9. Flag: non-zero if output grid wind field is to be printed

Table 7

Input Cards with Data Defining Program HIST Test Output Grid

- 1) 5 cards with 62 west longitudes in the form DDMM and progressing from west to east. Each card, except the last has 15 longitudes and a sequence number in format 16I5. The last card is blank filled between the last data field and the sequence number.
- 2) 3 cards with 31 latitudes progressing from south to north. Form and format are the same as those of the longitude cards.
- 3) 31 cards of terrain code. Each card is for one latitude, and cards are ordered from south to north. The format is 2I3, for degrees and minutes of latitude, then 2X,62I1, where the 62I1's are one-digit numeric terrain codes for each longitude and progress from west to east. As used in the test grid for Lake Pontchartrain, the codes are as follows -

1: open ocean  
2 lake  
3: marsh  
4: plains  
5: woods  
6: cities

Table 8

COMMON BlocksProgram SNAP

Block Name	Variable Name	Size in Words	Units	Source	Disposition	Description
C1	NAME	1		Input NAME3	Snapshot wind data file	4-character name of storm
	NSNAP	1				Sequence number of wind snapshot
	DX	1	km	Input NAME3	Snapshot wind data file	Grid spacing of innermost nest
	DT	1	sec			Time increment for computation of winds in innermost nest
	F	1				Coriolis force
	SGW	1	m/sec	Input NAME3	Snapshot wind data file	Surface geostrophic wind of ambient flow
	AN1	1	deg	Input NAME3	Snapshot wind data file	Direction of SGW counterclockwise from snapshot x-axis
	UC	1	m/sec			X-component of velocity of storm movement
	VC	1	m/sec			Y-component of velocity of storm movement
	UG	1	m/sec			X-component of surface geostrophic wind
	VG	1	m/sec			Y-component of surface geostrophic wind
	CS	1	m/sec			Speed of storm movement
	NM	1		Input NAME3		Number of times to cycle wind computation in innermost grid nest
	IB	1		Input NAME1		Flag: if zero, do not print pressure field or initial wind

Table 8 (continued)

<u>Block Name</u>	<u>Variable Name</u>	<u>Size in Words</u>	<u>Units</u>	<u>Source</u>	<u>Disposition</u>	<u>Description</u>
[C1]	ST12	1	km	Input NAME3	Snapshot wind data file	Distance from axis to $\frac{1}{2}$ magnitude of SGW
C2	JA	15		Input NAME3	Snapshot wind data file	See program SNAP output data file record description, page . Also note equivalence list in remarks, page .
	AB	6615			Work array, 1st third, $\partial p/\partial x$ ; 2nd third, $\partial p/\partial y$ ; last third, $\partial p/\partial r$	
	AC	3087			Dimensioned 21 × 7 × 21. Location in grid defined by 1st and 3rd subscripts. If 2nd subscript, N, is 1, value is cosine of angle of grid point; if 2, value is sine of angle of grid point; if 3-7, value is radius in meters of point in nest N-2.	
66	C3	U	2205	m/sec	Work array, x-component of boundary layer wind at previous time level	
	V	2205	m/sec		Work array, y-component of boundary layer wind at previous time level	
	UN	2205	m/sec	Snapshot wind data file	X-component of boundary layer wind	
	VN	2205	m/sec	Snapshot wind data file	Y-component of boundary layer wind	
	PX	2205			Work array, $\partial p/\partial x$	
	PY	2205			Work array $\partial p/\partial y$	
	VTN	2205	kn		Work array, holds wind speeds to be printed	
	ANG	2205	deg		Work array, holds wind directions to be printed	

(Continued)

Table 8 (continued)

<u>Block Name</u>	<u>Variable Name</u>	<u>Size in Words</u>	<u>Units</u>	<u>Source</u>	<u>Disposition</u>	<u>Description</u>
[C3]	LW	2205				Land/sea table: 1 for land, 2 for sea
C4	CDR	400				Work array, dimensioned 100 x 2 x 2.
						Drag coefficients:
						CDR(I,1,1) is upwind component of drag coefficients over land, when integrated wind speed is (.8*I) m/sec.
						CDR(I,2,1) is crosswind component over land
						CDR(I,1,2) is upwind component over ocean
						CDR(I,2,2) is crosswind component over ocean
UXV	200					Work array, dimensioned 100 x 2.
						UXV(I,1) is $\frac{U_x}{V_m}$ over land, when
						$V_m = (.8*I) \text{ m/sec}$
						UXV(I,2) is the same over ocean
TURN	200		rad			Work array, dimensioned 100 x 2.
						TURN(I,1) is the angle between surface wind and integrated wind, over land, when integrated wind speed is (.8*I) m/sec.
						TURN(I,2) is the same over ocean
C5	FLAT	1	°K	Input NAME2	Snapshot wind data file	Coriolis force
	PTH	1	°K	Input NAME2	Potential temperature	
DTH	2	°K	Input NAME2	Snapshot wind data file	(1) Air-land temperature difference (2) Air-sea temperature difference	

(continued)

Table 8 (continued)

<u>Block Name</u>	<u>Variable Name</u>	<u>Size in Words</u>	<u>Units</u>	<u>Source</u>	<u>Disposition</u>	<u>Description</u>
[C5]	HH	1	m	Input NAME2	Snapshot wind data file	Boundary layer height over water
	ZOLAND	1		Input NAME2		Roughness length over land
	LS					Index variable
	VV	100	m/sec			Vertically integrated wind speeds. VV(I) at point I = .8m/sec x I
	UX	3	m/sec			Latest 3 values of $U_*$ in an iterative loop
	UV	3	$m^2/sec^2$			Latest 3 values of integrated wind speed corresponding to $U_*$
	DUV	3	$m^2/sec^2$			Interpolated value of wind speed squared minus desired value squared
	K35	1		Input NAME2	Snapshot wind data file	Karman's constant, type real
	K2	1				$K35^2$ , type real
	G	1				Acceleration of gravity = $9.806 m/sec^2$
	GA	1	$sec^2/m$			Charnock's constant divided by G
	DEN	1				Used in stability length computation
	VV2	1	$m^2/sec^2$			$VV(I)^2$ for current I
	HL	1				HH/stability length
	K123	1				Index variable
	Z0	1	m			Roughness length

(continued)

Table 8 (continued)

<u>Block</u>	<u>Variable Name</u>	<u>Size in Words</u>	<u>Units</u>	<u>Source</u>	<u>Disposition</u>	<u>Description</u>
[C5]	ZLOG	1				Log ( Z0/HH )
	AM	1				Constant in Arya's logarithmic scale law
	BH	1				Constant in Arya's logarithmic scale law
	CM	1				Constant in Arya's logarithmic scale law
	FF	1	sec <sup>-1</sup>			Coriolis parameter. Retained for consistency with other versions of CCROSS; not used in this program.
					<u>Program HIST</u>	
	C57	PTH	1	oK	Snapshot wind data file	Potential temperature
	DTH	1	oK	Snapshot wind data file	Air-sea temperature difference	
	HH	1	m	Snapshot wind data file	Boundary layer height over water	
	ZCOEFF	15		Input NAMES		3 coefficients in formula relating $Z_o$ to $U_k$ for all terrains except open ocean, where Garratt's formula is used
	LAKE	1		Input NAMES		Number of terrains other than open ocean. Integer, 0-5
	VW	100	m/sec			Vertically integrated wind speeds at point I: $(VW(I) = I \times .8 \text{ meters/sec}$
	UX	3	m/sec			Latest 3 values of $U_k$ in an iterative loop

(continued)

Table 8 (continued)

<u>Block Name</u>	<u>Variable Name</u>	<u>Size in Words</u>	<u>Units</u>	<u>Source</u>	<u>Disposition</u>	<u>Description</u>
[C57]	UV	3	$m^2/sec^2$			Latest 3 values of square of integrated wind speed corresponding to $U_*$
DUV		3	$m^2/sec^2$			Difference between squares of interpolated value of wind speed and desired value
K35				Snapshot wind data file		Karman's constant, type real
K2					$K35^2$ , type real	
G	1		$m/sec^2$			Acceleration of gravity: 9.806 constant
GA	1		$sec^2/m$			Charnock's constant divided by G
DEN						Temporary variable used in stability length computation
VV2			$m^2/sec^2$		$VV(I)^2$ for current I	
HL						HH/stability length
K123						Index variable
20						Roughness length
ZLOG						$\log(z_0/HH)$
AM						Correction terms in Arya's logarithmic scaling law
BM						Correction terms in Arya's logarithmic scaling law
CM						Correction terms in Arya's logarithmic scaling law
UXV	600				(continued)	$U_*/VV$ for each I, terrain type

Table 8 (continued)

<u>Block Name</u>	<u>Variable Name</u>	<u>Size in Words</u>	<u>Units</u>	<u>Source</u>	<u>Disposition</u>	<u>Description</u>
[C57]	TURN	600	rad			Turning angle between surface wind and integrated wind for each I, terrain type
	COST	100				Cosine of TURN for current terrain
	SINT	100				Sine of TURN for current terrain
D1	LA0	1	rad			Latitude of storm eye at current hour, type real
	LO0	1	rad			West longitude of storm eye at current hour, type real
	ROT	1	deg			Clockwise angle to rotate snapshot wind field
	LA1	1	rad			Latitude of 1 point at which wind is wanted, type real
	LO1	1	rad			West longitude of point of which wind is wanted, type real
	DX	1	km	Snapshot wind data file		Grid spacing of innermost nest
	STHT	1	m			Height of measurement at current location
	LANSEA	1				Indicated sea or type of terrain for current location
	W1	1	kn			Wind speed at current location
	TH1	1	deg			Meteorological wind direction
	D	1	km			Distance between locations LA0, LO0 and LA1, LO1

(continued)

Table 8 (concluded)

<u>Block</u>	<u>Variable</u>	<u>Size in</u>	<u>Units</u>	<u>Source</u>	<u>Disposition</u>	<u>Description</u>
	<u>Name</u>	<u>Name</u>	<u>Words</u>			
[D1]	AL	1	deg			Bearing of point LA1,L01, from point LA0,LO0, clockwise from north
	UST	1	m/sec			Friction velocity at current location
D2	XX	4410		Snapshot wind		Winds on nested grid at current hour
D3	NSMAP1	100		data file		1st wind snapshot sequence number
	NSMAP2	100				2nd wind snapshot sequence number (0 if none)
	PCT	100				Interpolation distance between snapshots NSMAP1 and NSMAP2
	IPI	100				Flag: non-zero if listing of winds on output wanted
	NHT	1				Number of hours in storm history
	INTVN	1				Not used
	INTVI	1	hours			Interval at which to print winds on output grid
LGRID*	ILAT	31				List of output grid latitudes, south to north, in the format DDDMM
	ILONG	62				List of output grid west longitudes, west to east, in the format DDDMM

\*Not in main program, in subroutines RDGRID and PRLAKE only.

Table 9

Description of Program HIST arrays not in COMMON

Station data arrays - all dimensioned 100

MAD	Location, from card input - degrees of north latitude
MAM	Location, from card input - minutes of latitude
MMOD	Location, from card input - degrees of west longitude
MOM	Location, from card input - minutes of longitude
LLAKE	Terrain code, from card input
STAHT	Height in meters, from card input
KSTA	Station number (identification), from card input
YLA	Latitude in radians
YLO	Longitude in radians

History table arrays - all dimensioned 100

NAD	Location of eye of storm, from card input - degrees of north latitude
NAM	Location of eye of storm, from card input - minutes of latitude
NOD	Location of eye of storm, from card input - degrees of west longitude
NOM	Location of eye of storm, from card input - minutes of longitude
IROT	Grid rotation angles, from card input
KDATE	Date in form YYMMDD
KTIME	Hour of KDATE
JSEQ	Sequence numbers of nested grid winds on work file

Output grid data arrays as used with test grid

ZLA	List of latitudes of grid points, south to north, in radians
ZLO	List of west longitudes of grid points, west to east, in radians

**Table 9 (concluded)**

**ZANG** Deviation between true north and grid north for each grid point - zero throughout test grid  
**LSTAB** List of terrain codes of grid points.

**Other arrays**

**XY** Dimensioned  $21 \times 21 \times 10$ . Contains 2nd snapshot wind field when needed

Table 10

Program Stops

<u>Program</u>	<u>Stop Number</u>	<u>Interpretation</u>
SNAP	999*	Normal completion of run
HIST	999*	Normal completion of run
HIST	5	Error in reading station or history input card
HIST	146	Error in reading mass storage work file of wind fields on nested grid
HIST	444	Snapshot interpolation distance out of range ( < 0 or > 1 ).
HIST	515	Input card and snapshot data file storm identifications are different
HIST	516	Too many station input cards ( > 100 )
HIST	517	Too many history input cards ( > 100 )
HIST (RDGRID)	21	Error in reading grid longitude card
HIST (RDGRID)	23	Error in reading grid latitude card
HIST (RDGRID)	51	Error in reading grid terrain code card

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\* If running a FORTRAN that requires STOP number to be octal, substitute 777.

Table 11  
Output Data File Record Description

Snapshot wind field record from program SNAP

Variable Name	(Dimension)	Size	Accumulated Word Count	Units	Applicable Format*	Description
UN	(21,21,5)	2205	2205	m/sec	F4.1	X-component of boundary layer wind
VN	(21,21,5)	2205	4410	m/sec	F4.1	Y-component of boundary layer wind
NAME		1	4411		A4	4-character name of storm
DX		1	4412	km	F3.0	Grid spacing of innermost nest
JA	(15)	15	4427		I1	1: Indicator for coding of (4) If 0, (4) in degrees; if 1, (4) in points of 11.25 degrees.
				deg	F5.1	2: North latitude of eye of storm
				deg	F6.1	3: West longitude of eye of storm
			see JA(1)		F4.0	4: Direction of track of storm, clockwise from north
				kn	F5.1	5: Forward speed of storm
					I1	6: Indicator for quadrants of pressure field

\* Format appropriate for printing this variable or array.

Table 11 (continued)

<u>Variable Name</u>	<u>(Dimension)</u>	<u>Size</u>	<u>Accumulated Word Count</u>	<u>Units</u>	<u>Applicable Format</u>	<u>Description</u>
<b>[JA]</b>						
						0 - circularly symmetric pressure field 1 - 1st quadrant is right front 2 - 1st quadrant is forward
mb		F6.1				7: Pressure at eye of storm
nm		F4.0				8-11: Exponential pressure profile scale radius in four quadrants. If JA(6) is zero, JA(9) - JA(11) may not contain valid data
mb		F5.0				12-15: Far field pressure in four quadrants. If JA(6) is zero, JA(13) - JA(15) may not contain valid data
SGW	1	4428		m/sec	F3.0	Surface geostrophic wind of ambient flow
AN1	1	4429		deg	F4.0	Direction of SGW counterclockwise from snapshot x-axis
ST12	1	4430		km	F5.1	Distance from axis to $\frac{1}{2}$ magnitude of SGW
DTH	2	4432		°K		(1) Air-land temperature difference (2) Air-sea temperature difference
HH	1	4433		m		Boundary layer height over water
GARR	1	4434				Charnock's constant
PTH	1	4435		°K		Potential temperature
K35	1	4436				Karmen's constant, type real
						(continued)

Table 11 (concluded)

Test output grid wind field record from program HIST

<u>Variable Name</u>	<u>(Dimension)</u>	<u>Size</u>	<u>Accumulated Word Count</u>	<u>Units</u>	<u>Applicable Format</u>	<u>Description</u>
NBASE		1	1		A4	4-character name of storm
KHR		1	2		I3	Sequence number of hour of storm
ISTART		1	3		I8	Starting time of storm in format YYMMDDHH (Time at which KHR=1)
IZONE		1	4		A3	Time zone of ISTART
IMAX		1	5		I2	Longitudinal dimension of output grid
JMAX		1	6		I2	Latitudinal dimension of output grid
GRIDHT		1	7	m	F6.1	Height to which wind speeds are scaled
NAD		1	8		I3	Location of eye of storm - degrees of north latitude
NAM		1	9		I2	Location of eye of storm - minutes of latitude
NOD		1	10		I4	Location of eye of storm - degrees of west longitude
NOM		1	11		I2	Location of eye of storm - minutes of longitude
WIND	(2,MAXI,MAXJ) (3844 for test grid)	11+2×MAXI×MAXJ				Wind at specified height on wave grid. If 1st subscript is 1, value is wind speed in knots. If 1st subscript is 2, value is direction in degrees toward which wind blows counterclockwise from north

Printed Output

. Program SNAP initially prints card input data and finally prints an end-of-job message. For each snapshot it prints card input data pertinent to that snapshot, a pressure field and an initial guess wind field if requested, and a final snapshot wind field on the nested grid. Snapshot wind speeds are printed in tenths of knots, directions are meteorological, and west is at the top of the page. Each snapshot is printed both with wind speeds as computed and with wind speeds scaled to 19.5 meters.

. Program HIST prints card input data, the wind history at each requested station, wind fields on the output grid for hours requested, and an end-of-job message. Output grid wind speeds are in knots and directions are meteorological. Terrain types are indicated by blank for type 1, '\*' for 2, '=' for 3, '-' for 4, '+' for 5 and '\$' for 6.

Table 12

Program Changes Needed for a New Output Grid

- 1) Subroutine RDGRID
- 2) Subroutine PRLAKE
- 3) COMMON block LGRID (used in RDGRID and PRLAKE only)
- 4) In program Hist:
  - a) Parameters MAXI, MAXJ, and IGRDHT, where MAXI is the longitudinal dimension of the grid, MAXJ is the latitudinal dimension, and IGRDHT is its height in tenths of meters.
  - b) ZLA, ZLO become two-dimensional if rows and columns do not fall on latitude, longitude lines, and the settings of LA1 and LO1 in the DO 111 and DO 110 loops will be affected
  - c) Array ZANG, the angle between true meridian and grid meridian, may become non-zero.

#### PART IV: CONCLUSIONS AND RECOMMENDATIONS

. A method is developed to specify the surface stress and the wind speed and direction in the planetary boundary layer of a tropical cyclone from meteorological storm parameters available for historical hurricanes. The method is based upon a numerical primitive-equation model of the planetary boundary layer in a moving tropical cyclone. The complete time history of the evolution of the surface wind field is described from a series of characteristic wind field states calculated at discrete times in a storm's history by the steady-state model.

. A surface drag formulation, based upon a contemporary similarity model (Arya, 1977) coupled with a roughness parameter specification for a water surface consistent with Cardone's (1969) law, is incorporated into the numerical model. As a result, the model was found to produce a consistent description of the integrated planetary boundary layer wind, the magnitude and direction of the surface stress, and the wind speed and direction at anemometer level, without recourse to arbitrary, empirical calibration schemes. The surface winds calculated in several recent hurricanes are found to be in excellent agreement with available, representative surface wind measurements made from offshore platforms and data buoys.

. Transformations based upon an equilibrium planetary-boundary-layer similarity model are developed to specify the surface wind over terrain of specified roughness, including lake surfaces, from the over-water wind-field solution. Calculated over-land and over-lake winds are compared to the limited measurements available for several recent storms. Agreement is generally good.

. The principal limitation of the model is the neglect of fetch effects in the adjustment of the PBL across roughness discontinuities. Adjustments in near-surface wind speed, however, are believed to occur sufficiently rapidly that accuracy over homogeneous terrain and lakes the size of Pontchartrain and Okeechobee should not be significantly limited. The adjustment scale for wind direction, however, might be much larger.

. The principal obstacle to further development and evaluation of the method developed here is the lack of high-quality measurements of surface winds over the lakes of interest in well documented storms. As part of an intensive field program conducted in Lake Pontchartrain by the U.S. Army Corps of Engineers during the past year, wind data was apparently collected at several points in the lake during the passage of two hurricanes (Bob and Frederick, 1979). Even though only the peripheral parts of those storms were sampled, it is strongly recommended that those measurements be carefully processed and that the method developed in this study be applied to those storms.

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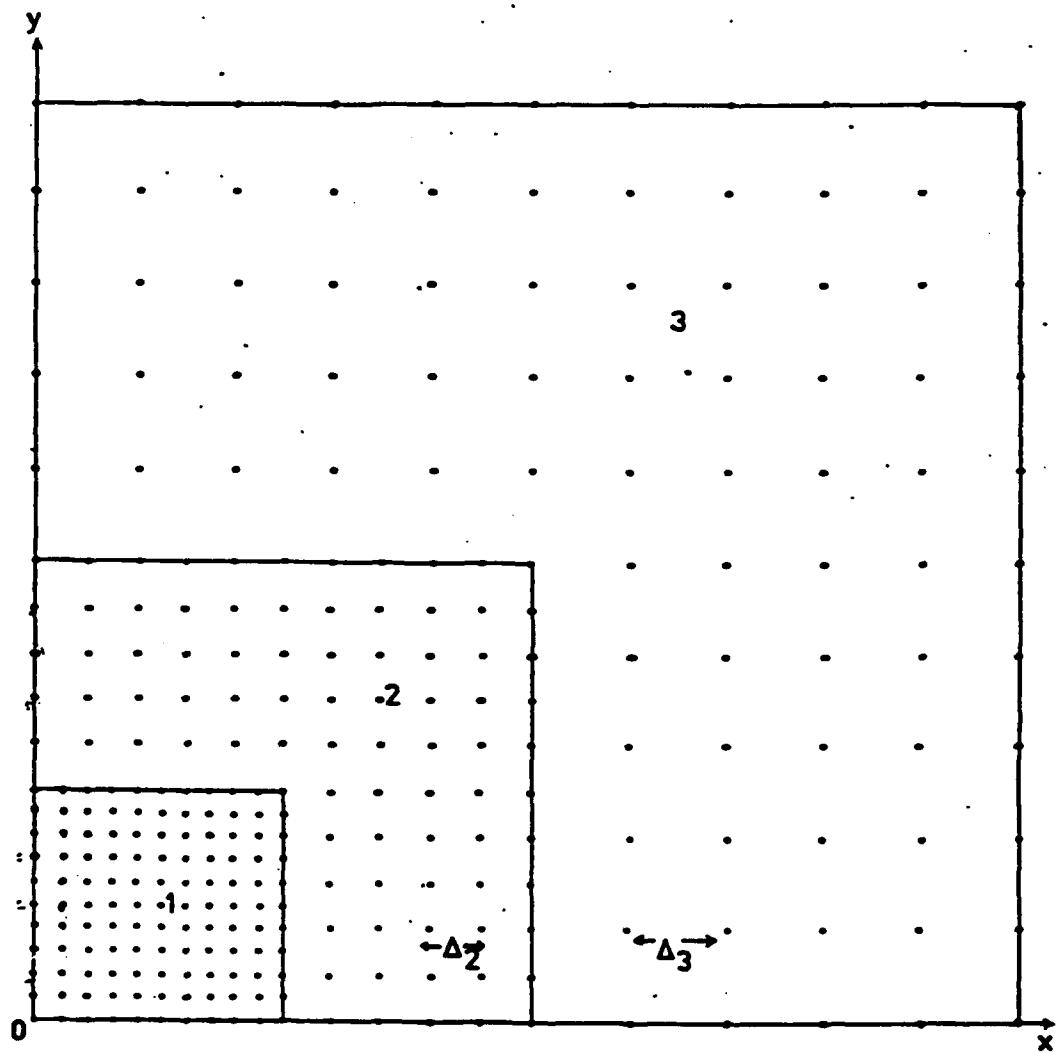


Figure 1. Grid points of the inner three grids in one quadrant of the nested grid system. The center of the grid system is indicated by 0 (from Chow 1971)

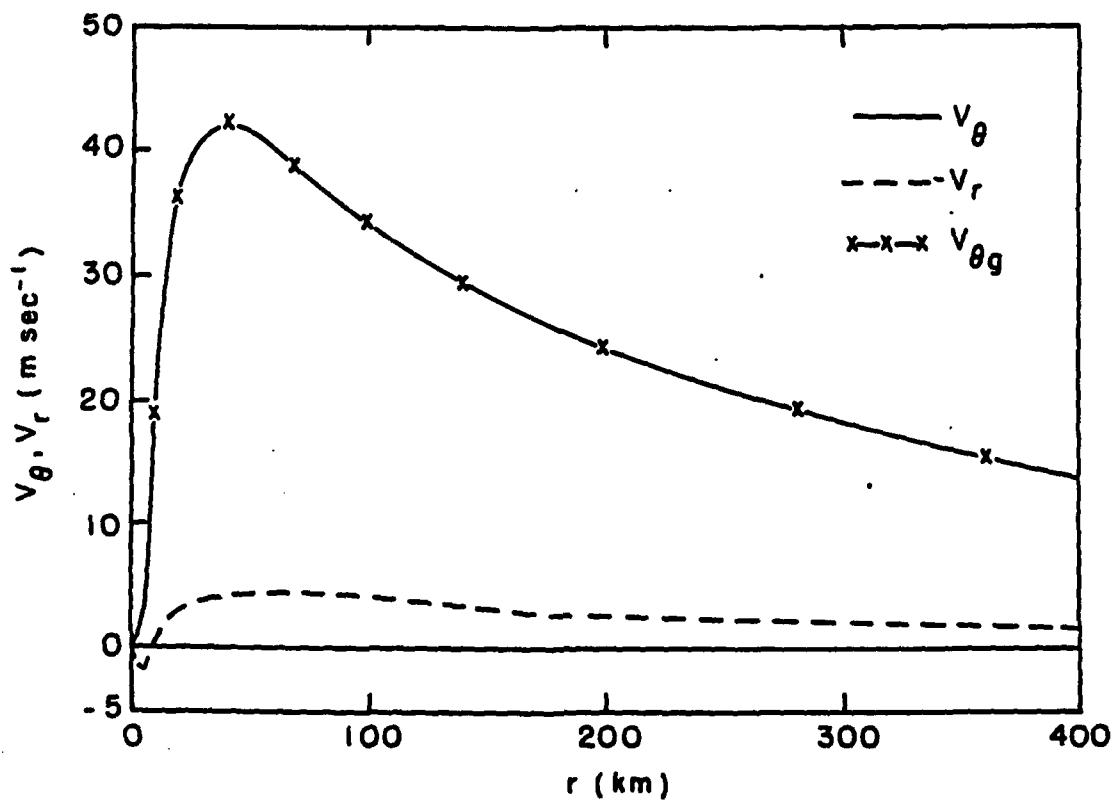
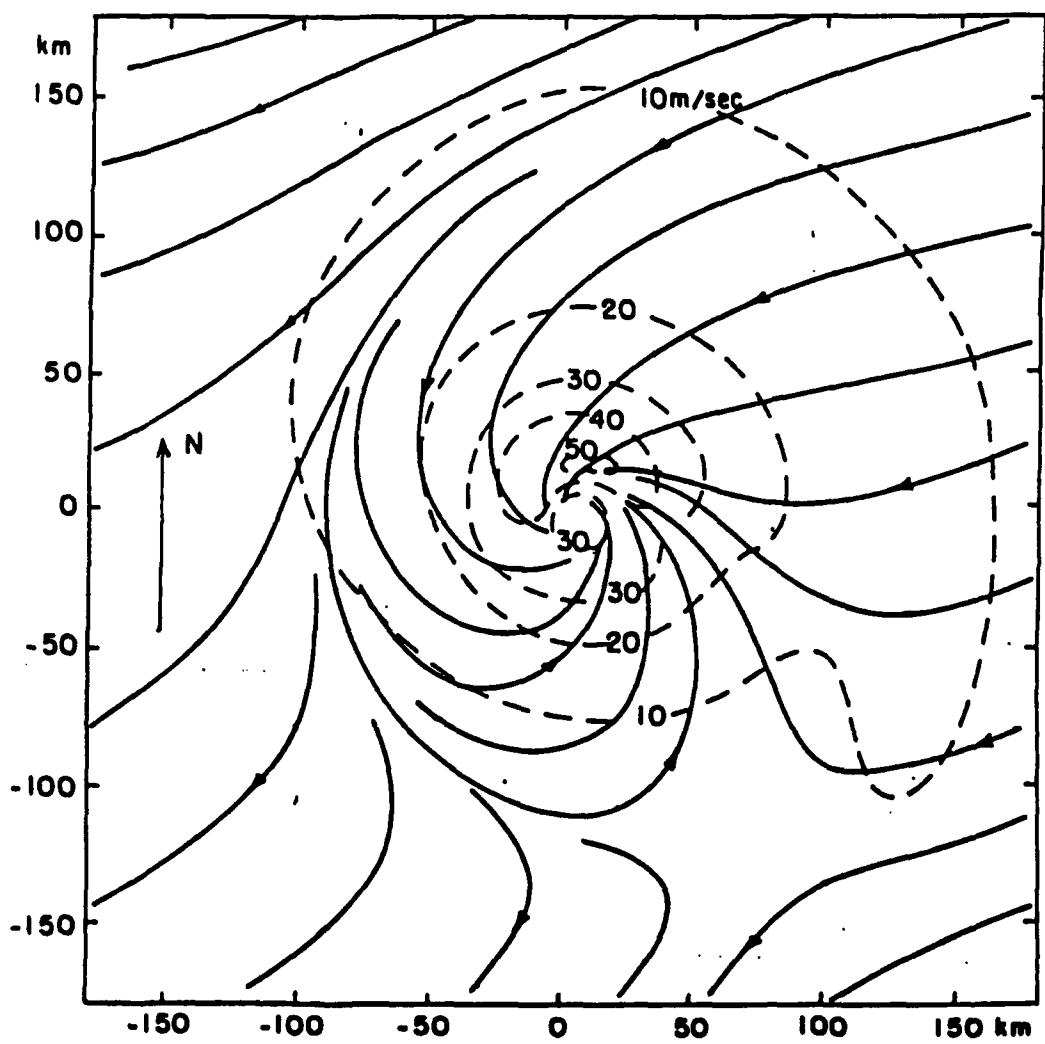


Figure 2. Radial distribution of tangential velocity ( $V_\theta$ ) and radial velocity ( $V_r$ ) for a frictionless, symmetrical, stationary storm given by  $\Delta p = 50$  mb and  $R = 40$  km, computed from Chow's numerical model. Analytical (gradient wind solution) solution ( $V_{\theta g}$ ) for specified pressure field is shown (from Chow 1971)



**Figure 3.** Streamlines (solid lines) and isotachs (dashed lines) of the steady-state solution for the vertically integrated boundary layer wind in a mature tropical cyclone moving westward at 10 m/sec, in a westerly steering flow (10 m/sec), from the model of Chow (1971)

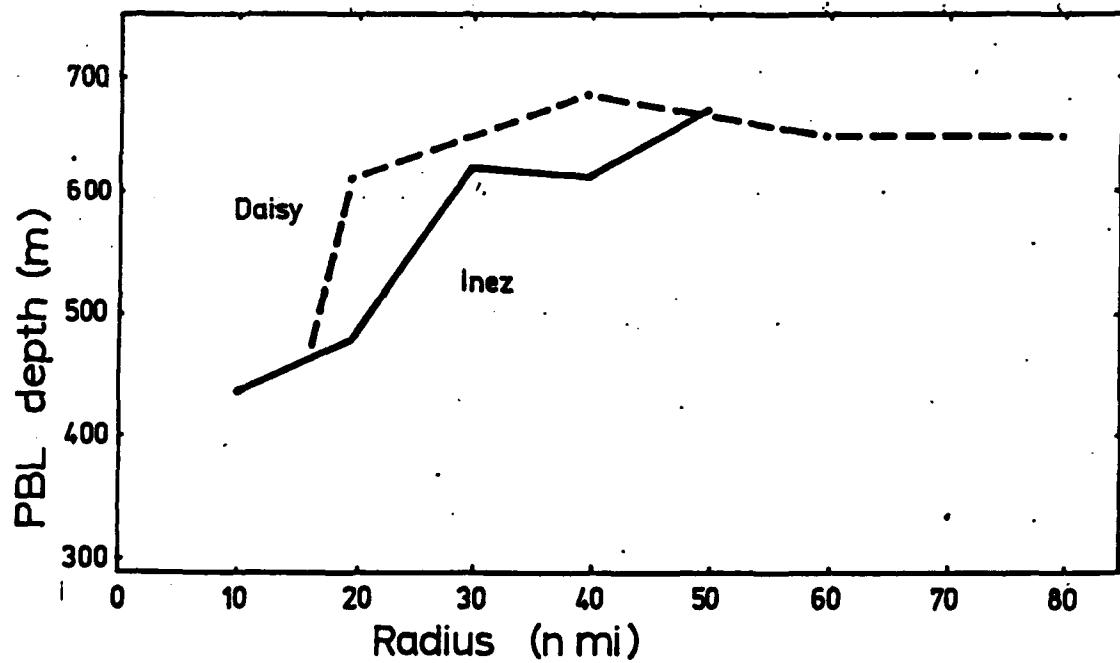


Figure 4. Computed depth of the planetary boundary layer versus radial distance from the eye for Hurricanes Daisy (1958) and Inez (1966)  
(from Moss and Rosenthal 1975)

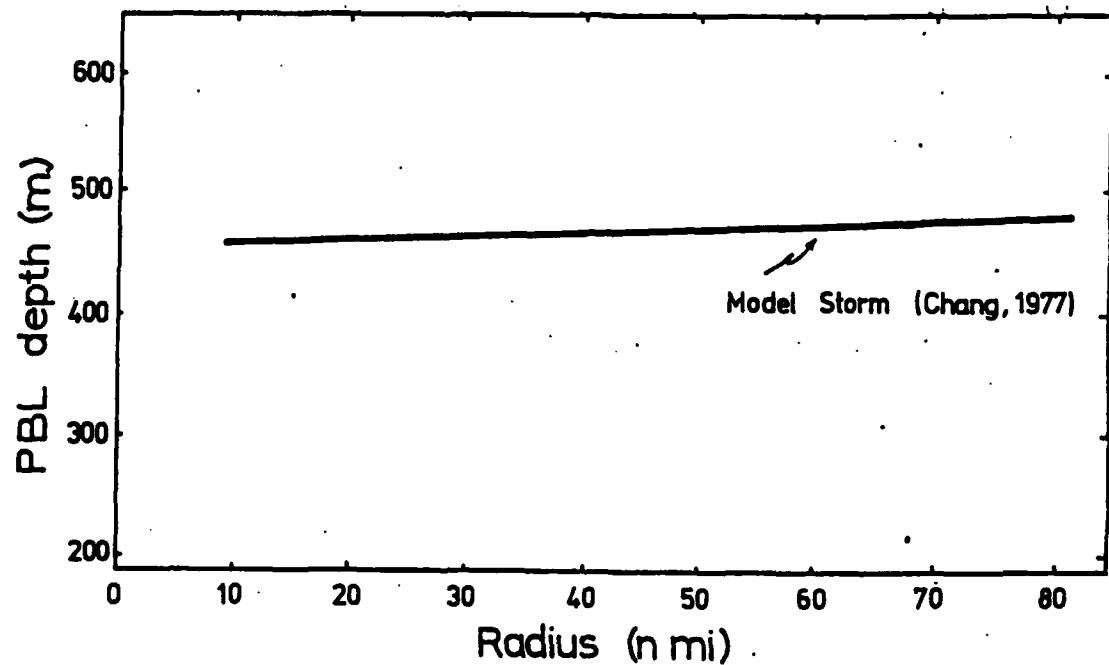


Figure 5. Computed depth of the planetary boundary layer in a mature, steady-state tropical cyclone (from Chang 1977)

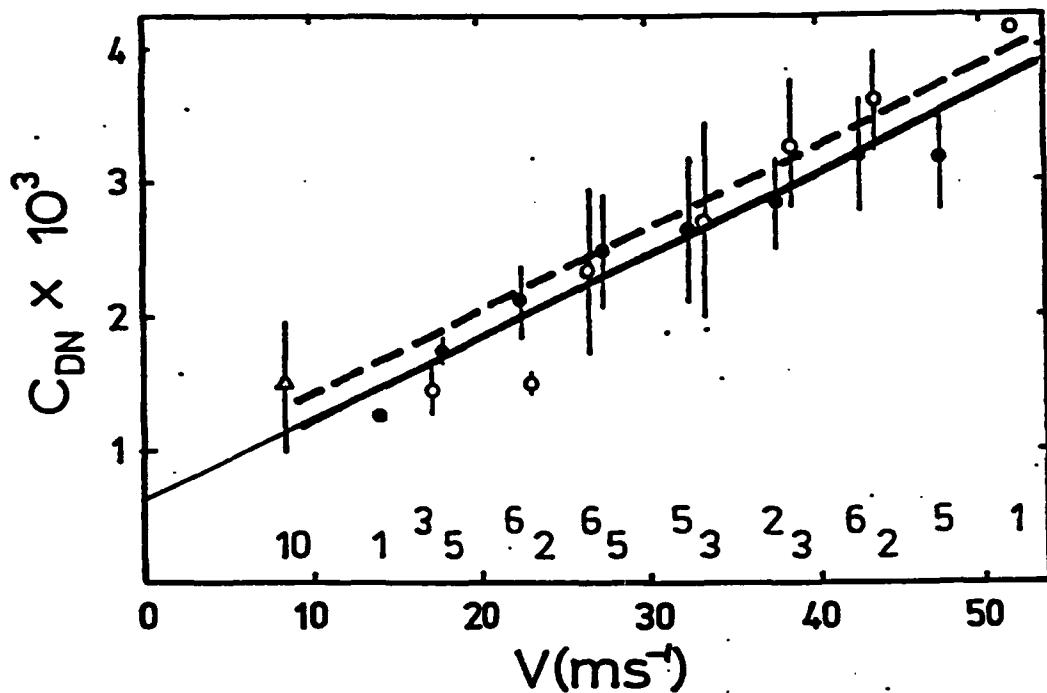


Figure 6. Garratt's (1977) collection of mean values of the drag coefficient as a function of wind speed at the 10-m height for 5-m/sec intervals, based on individual data from hurricane studies (○), wind flume experiments (●), and vorticity/mass budget analysis (Δ). Vertical bars refer to the standard deviation of individual data for each mean, with the number of data used in each mean shown below each mean value immediately above the abscissa scale. The dashed curve represents the variation of the 10-m neutral drag coefficient  $C_{DN}$  with wind speed based on

$$z_0 \approx a \frac{u_*^2}{g}$$

with  $a = 0.0144$  and a value of the barrier constant  $k$  of 0.41. The solid curve represents the variation with  $a = 0.035$  and  $k = 0.35$

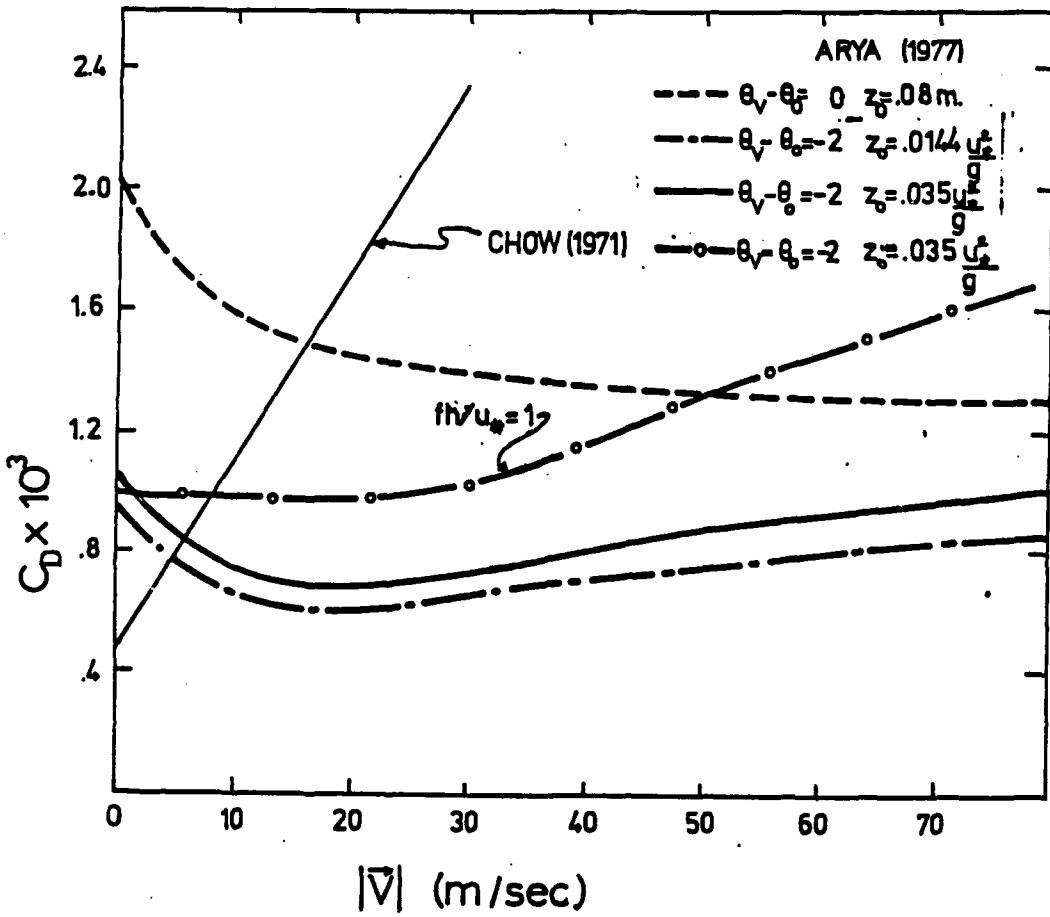


Figure 7. Drag coefficient with respect to the vertically integrated planetary boundary layer wind versus integrated boundary layer wind from Arya's model for alternate air-sea temperature and roughness parameter specifications and/or the case of restricted scale/height ratio. The form used by Chow (1971) is shown for comparison

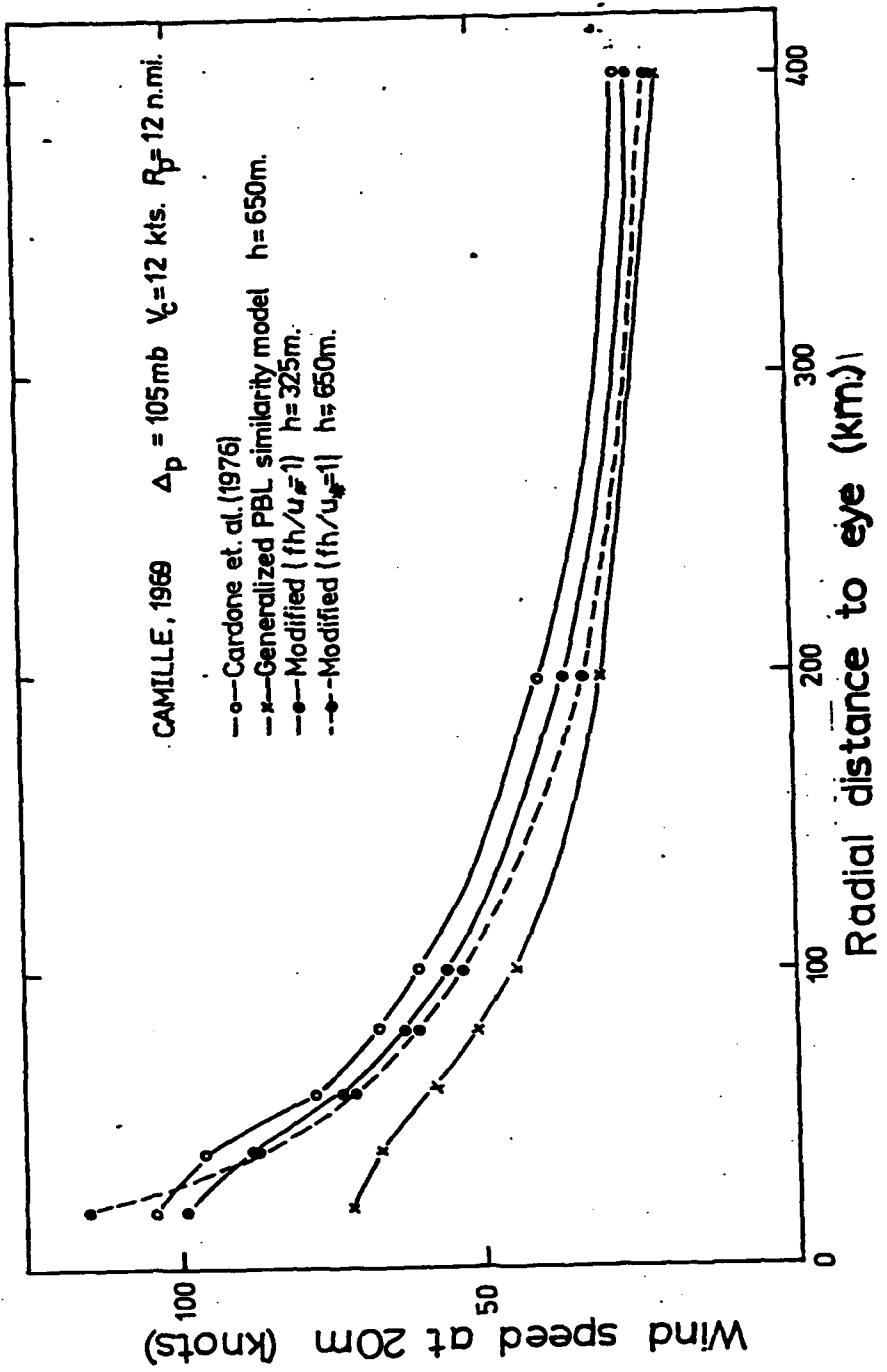


Figure 8. Predictions of 20-m wind speed versus radial distance to eye in Camille from the model of Cardone et al. (1976) and from the model with revised PBL stress law

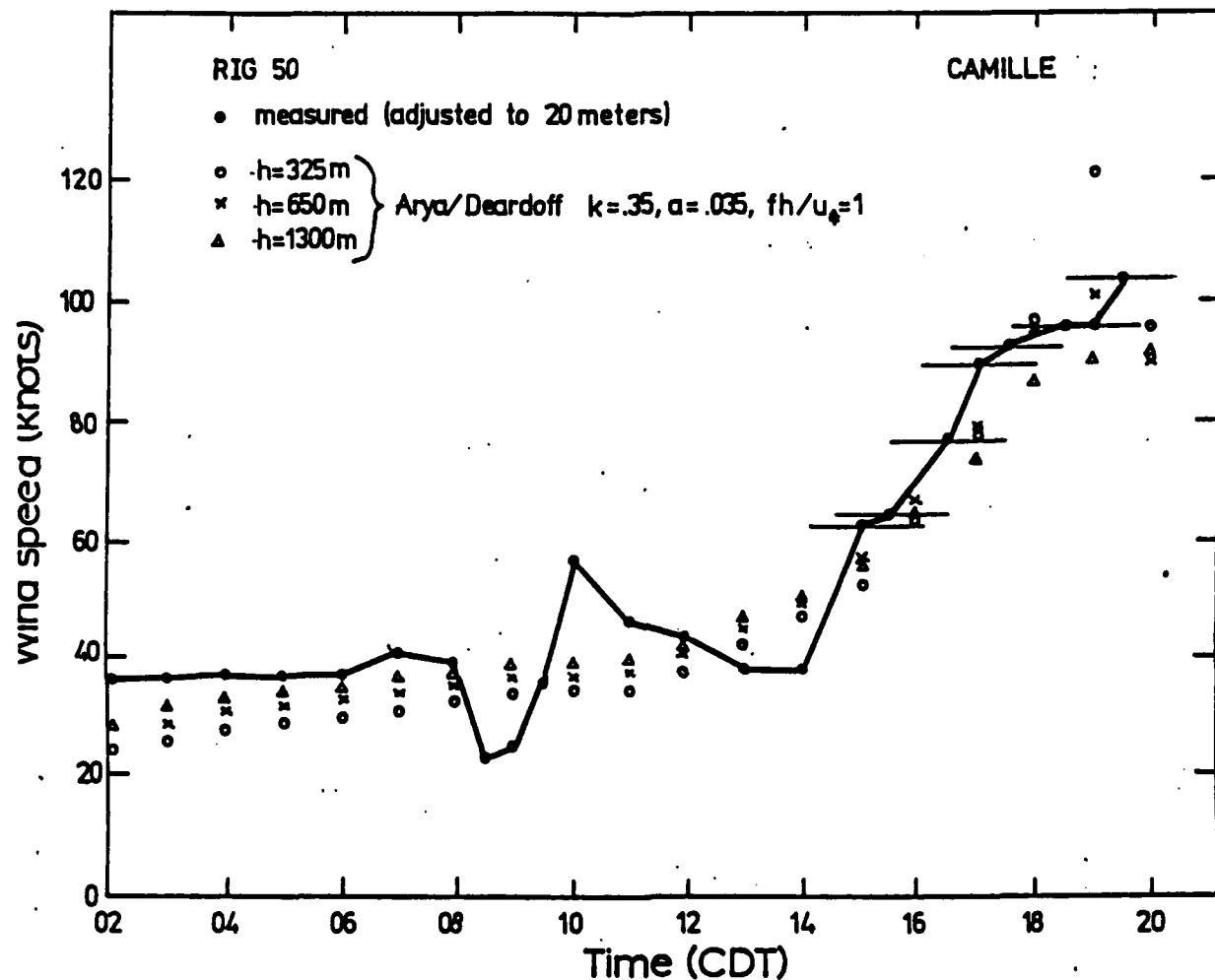


Figure 9. Comparison of modelled and measured surface wind speed at Rig 50 in Camille for vortex model with modified similarity model drag law and alternate boundary layer heights

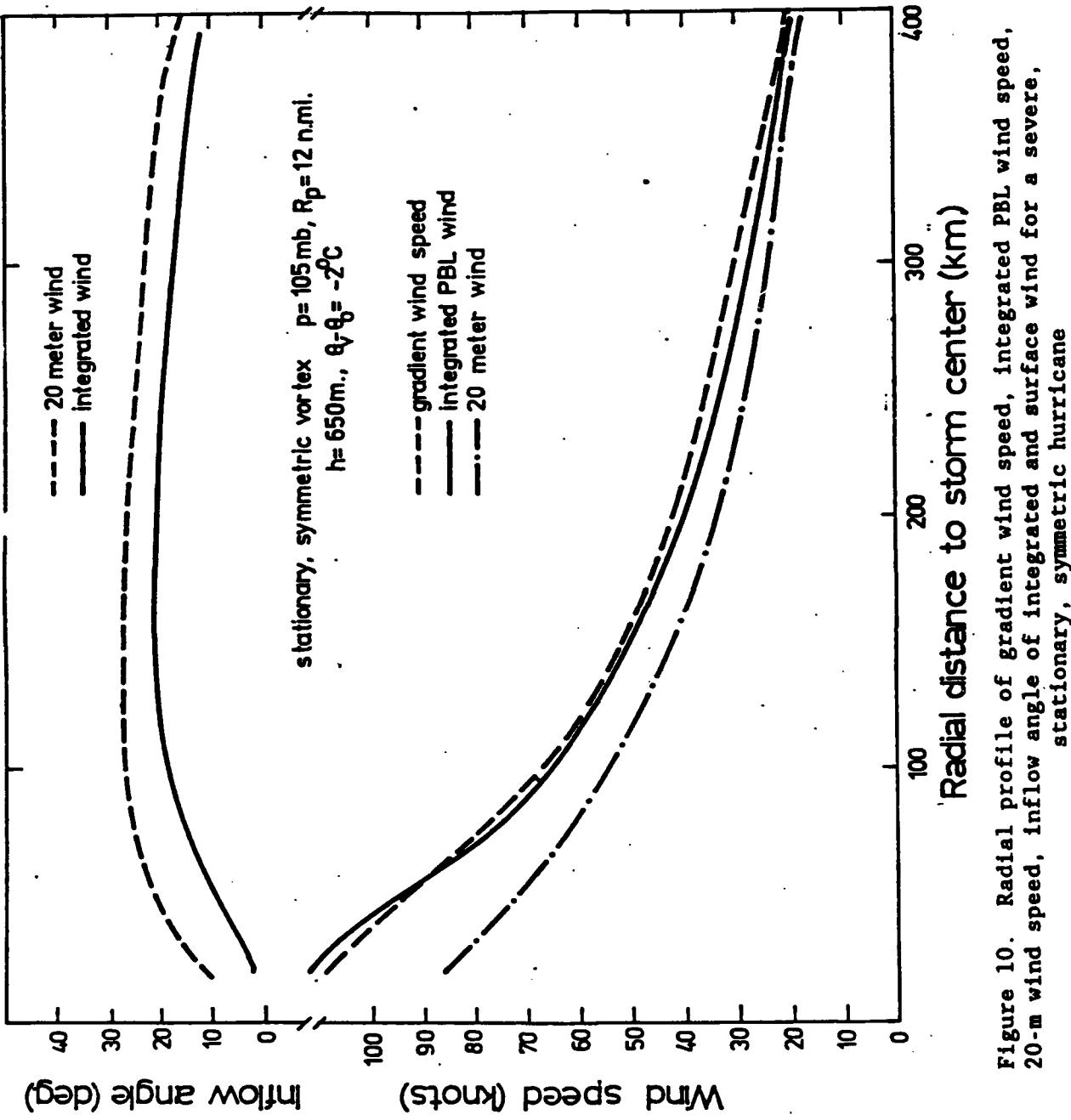


Figure 10. Radial profile of gradient wind speed, integrated PBL wind speed, 20-m wind speed, inflow angle of integrated and surface wind for a severe, stationary, symmetric hurricane

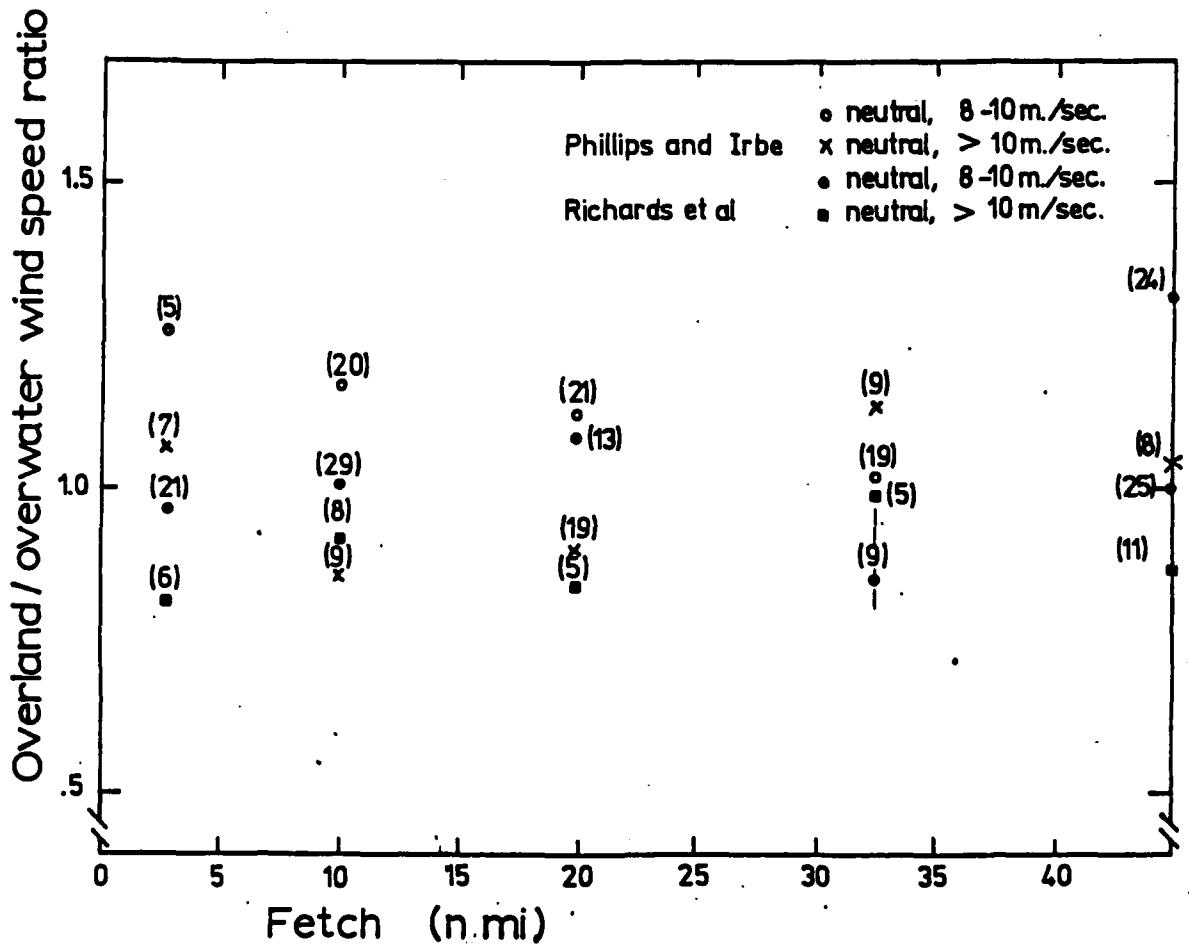
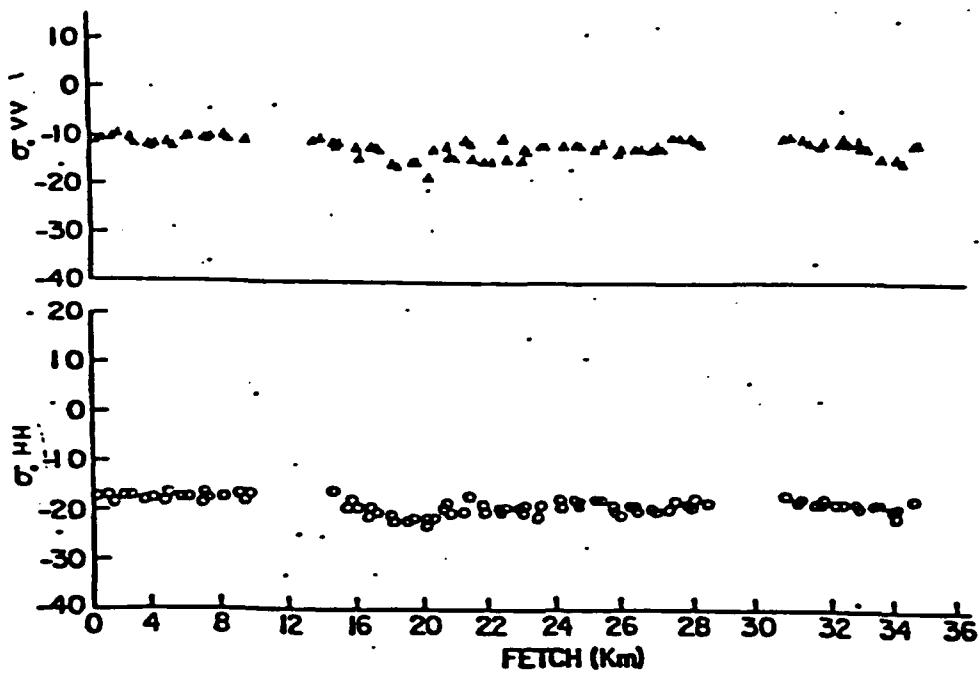
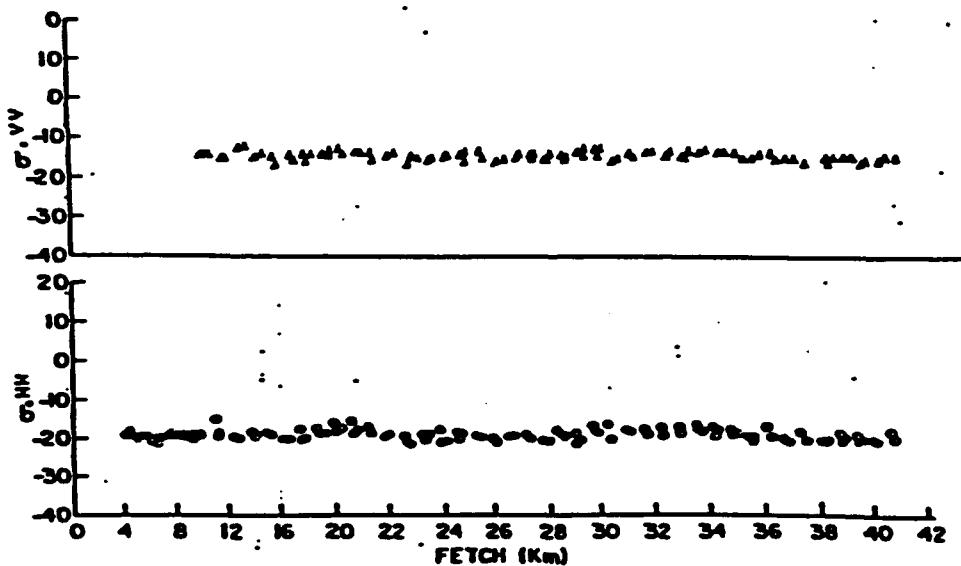


Figure 11. Ratio of overland/over-water wind speed versus fetch for near neutral moderate-high wind speed data of Richards et al. (1966) and Phillips and Irbe (1977)



Radar scattering cross section  $\sigma_0$  as a function of fetch at an incidence angle of  $40^\circ$ . Surface wind speeds were  $9.0 \text{ m s}^{-1}$ .



Radar scattering cross section  $\sigma_0$  as a function of fetch at an incidence angle of  $53^\circ$ . Surface wind speeds were  $13.0 \text{ m s}^{-1}$ .

Figure 12. Radar scattering cross section  $\sigma_0$  as a function of fetch at an incidence angle of  $40^\circ$  deg for wind speeds of  $9.0 \text{ m/sec}$  (above) and an incidence angle of  $53^\circ$  deg for wind speeds of  $13.0 \text{ m/sec}$  (below)

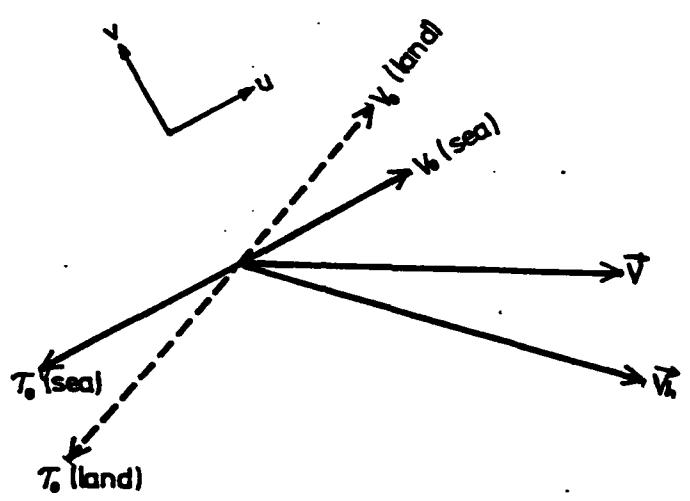


Figure 13. Coordinate system and relationship of wind stress components in the equilibrium model PBL over rough and smooth (sea) terrain

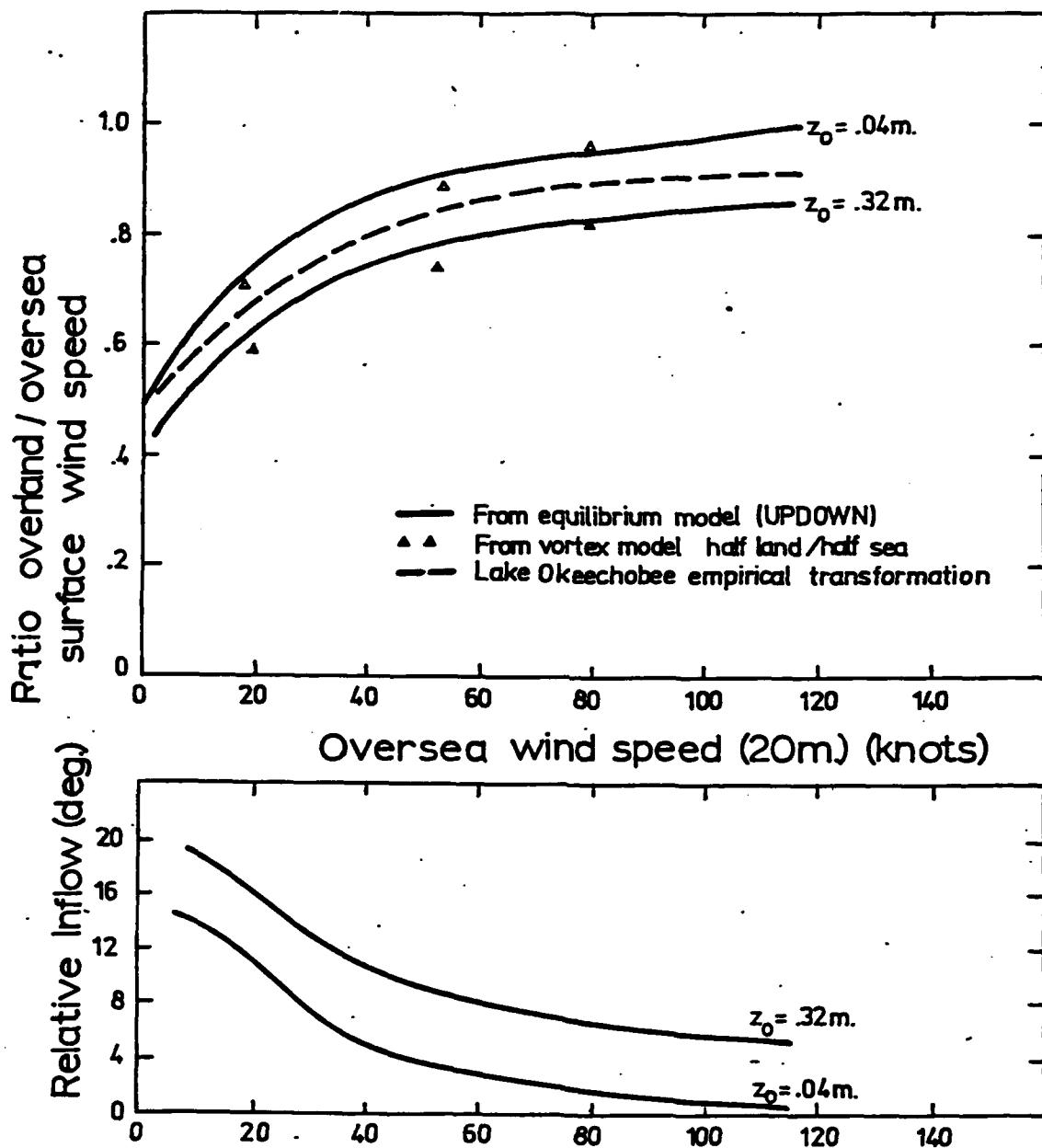


Figure 14. Ratio of the surface wind speed at the 20-m height overland to over sea (above) and the difference between the overland and over-sea inflow angle (below) from the equilibrium PBL model for two terrain roughnesses, from the numerical vortex model and the empirical wind speed ratio derived from measurements in hurricanes at Lake Okeechobee

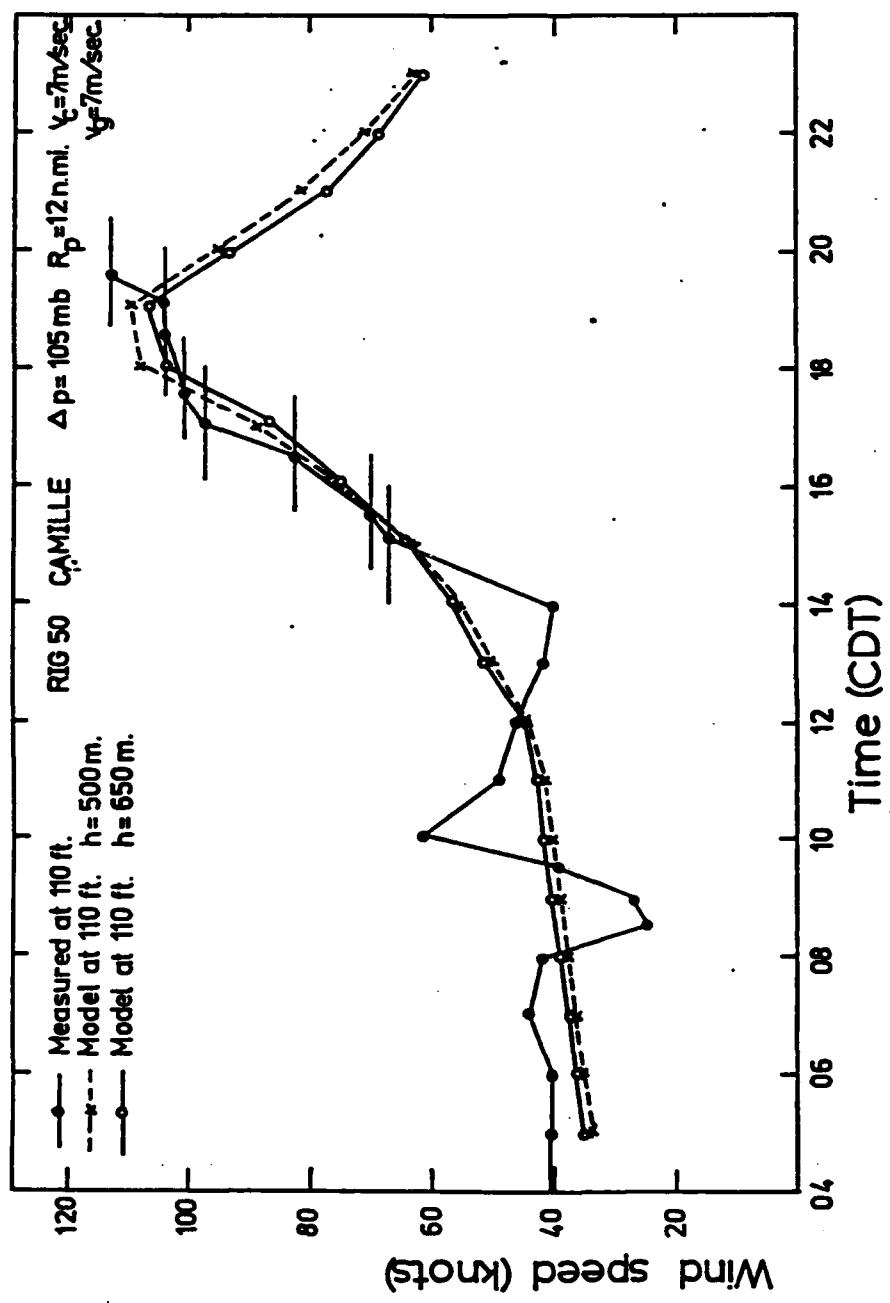


Figure 15. Comparison of measured and modelled wind speed at Rig 50 at measurement height in Camille

16 Z<sub>0</sub> Values for Typical Terrain Types (After ESDU 72026, 1972)

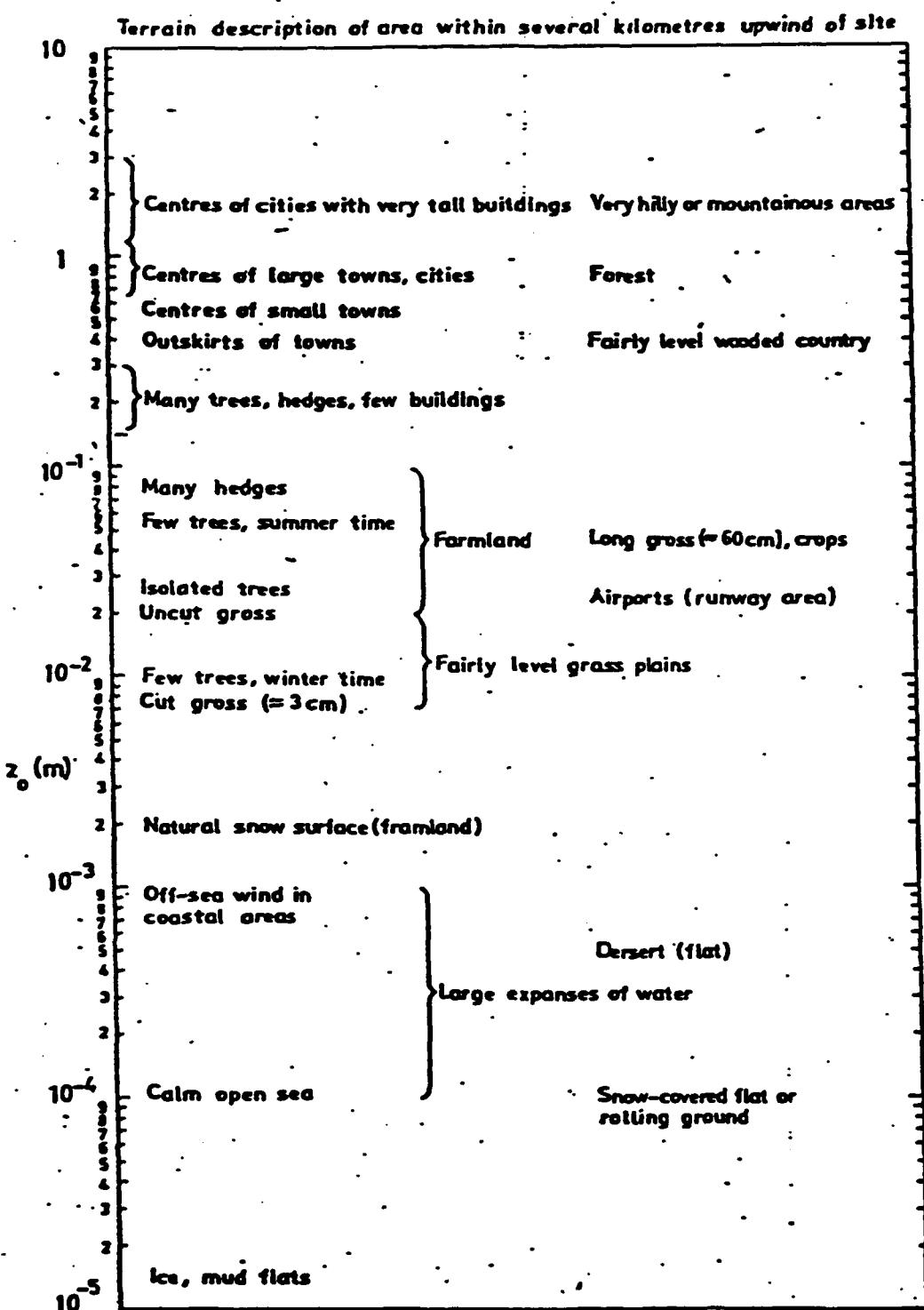


Figure 16. Roughness parameter values for typical terrain types  
(after ESDU 72026, 1972)

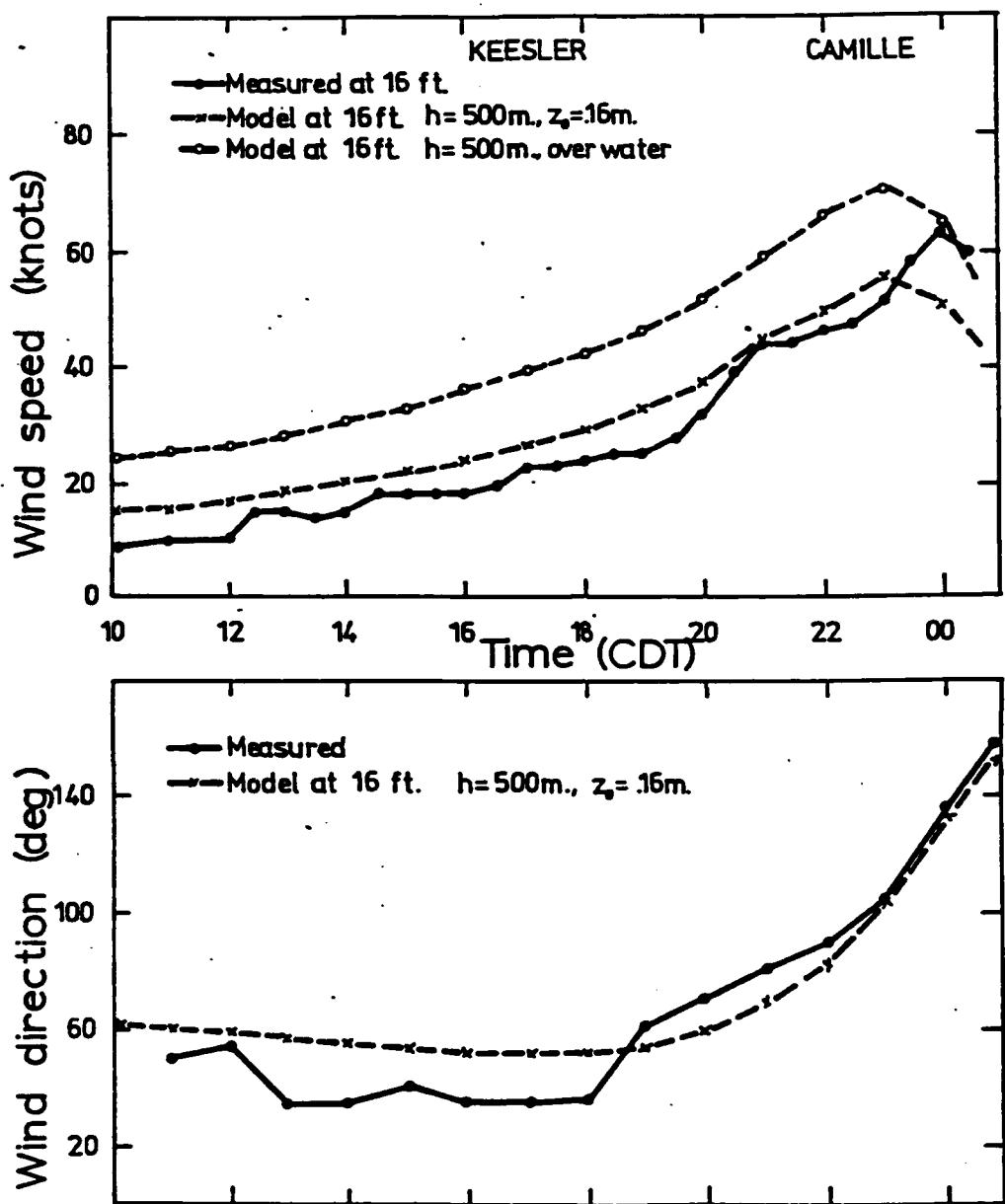


Figure 17. Comparison of measured and modelled winds at Keesler Air Force Base, Biloxi, MS, in Camille

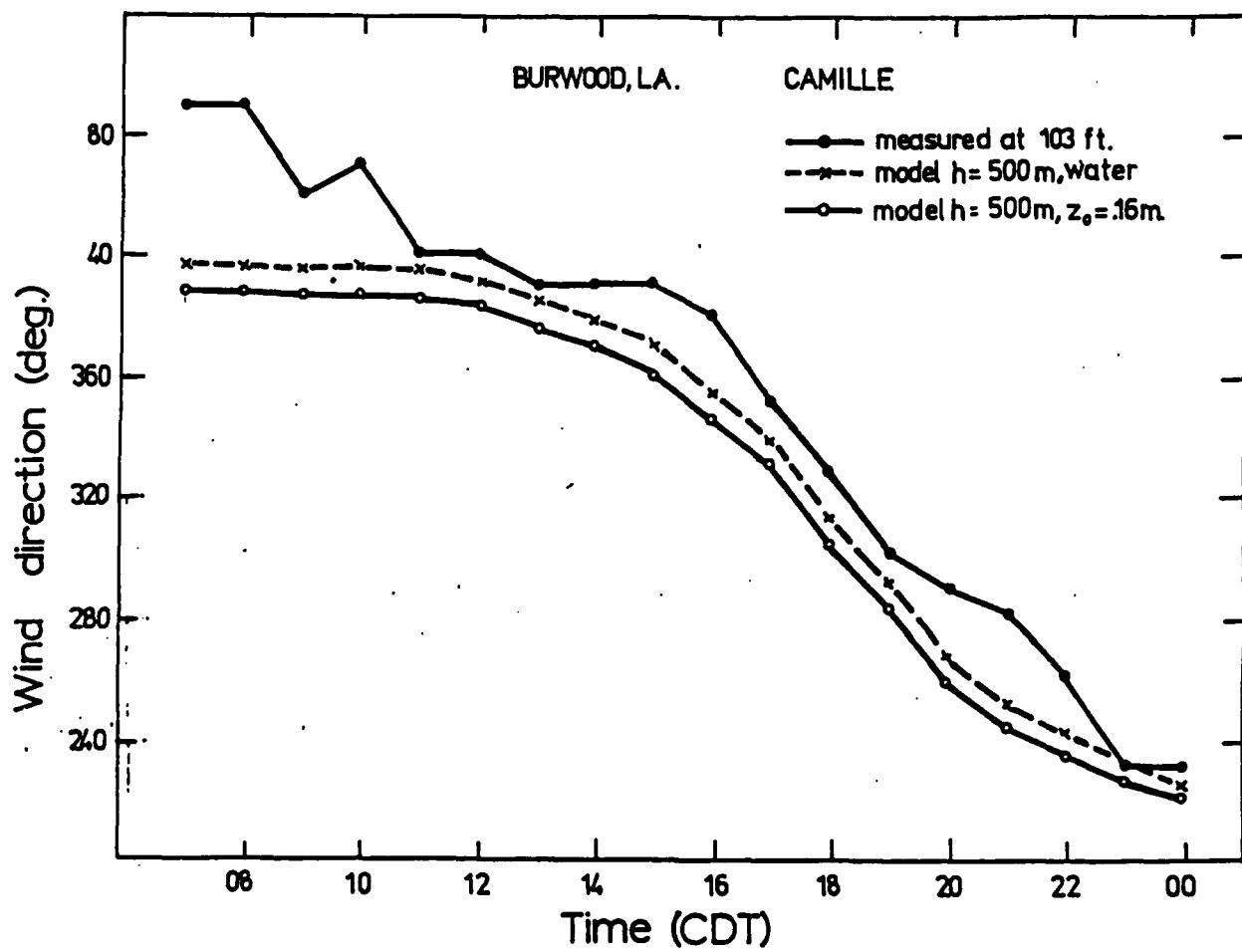


Figure 18. Comparison of measured and modelled wind direction at Burwood, LA, in Camille

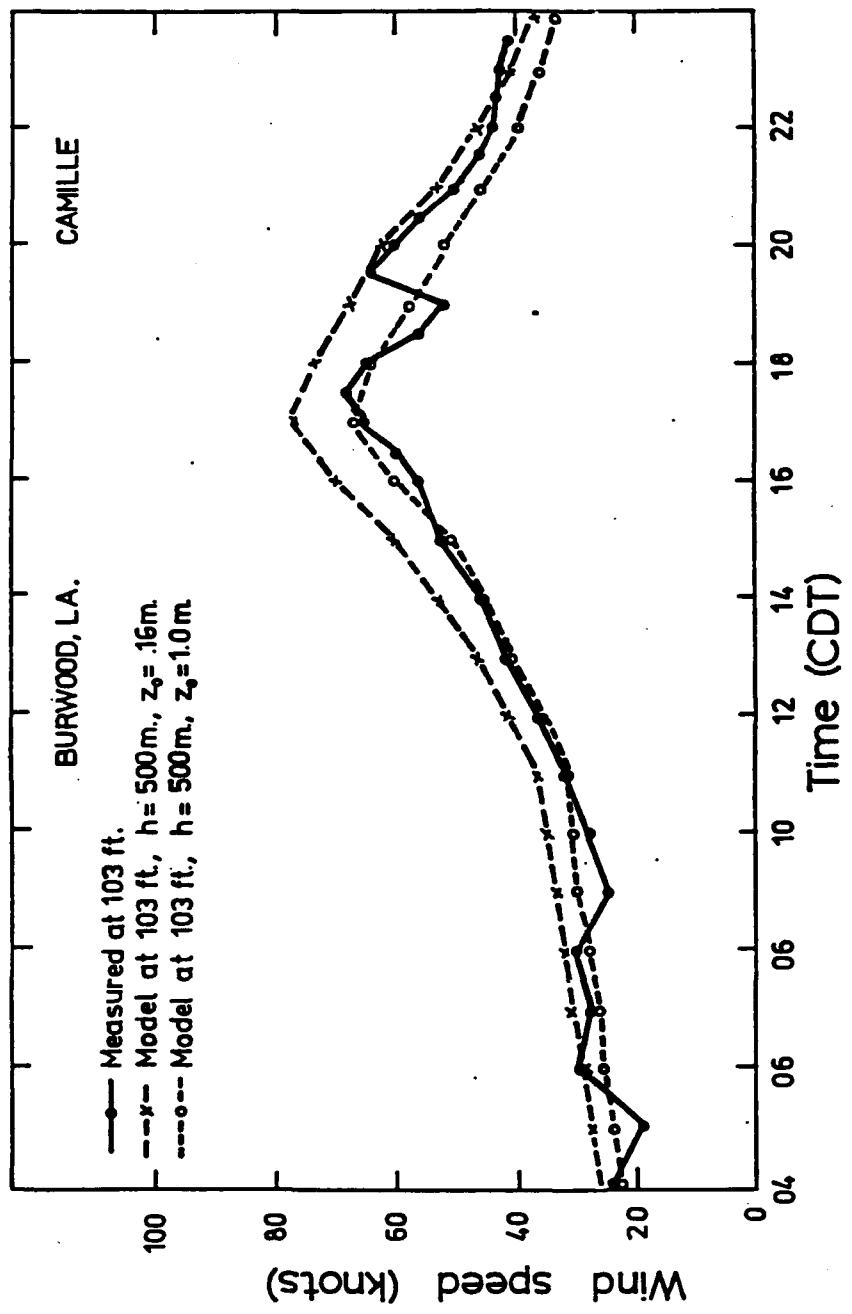


Figure 19. Comparison of measured and modelled wind speed at Burwood, LA, in Camille

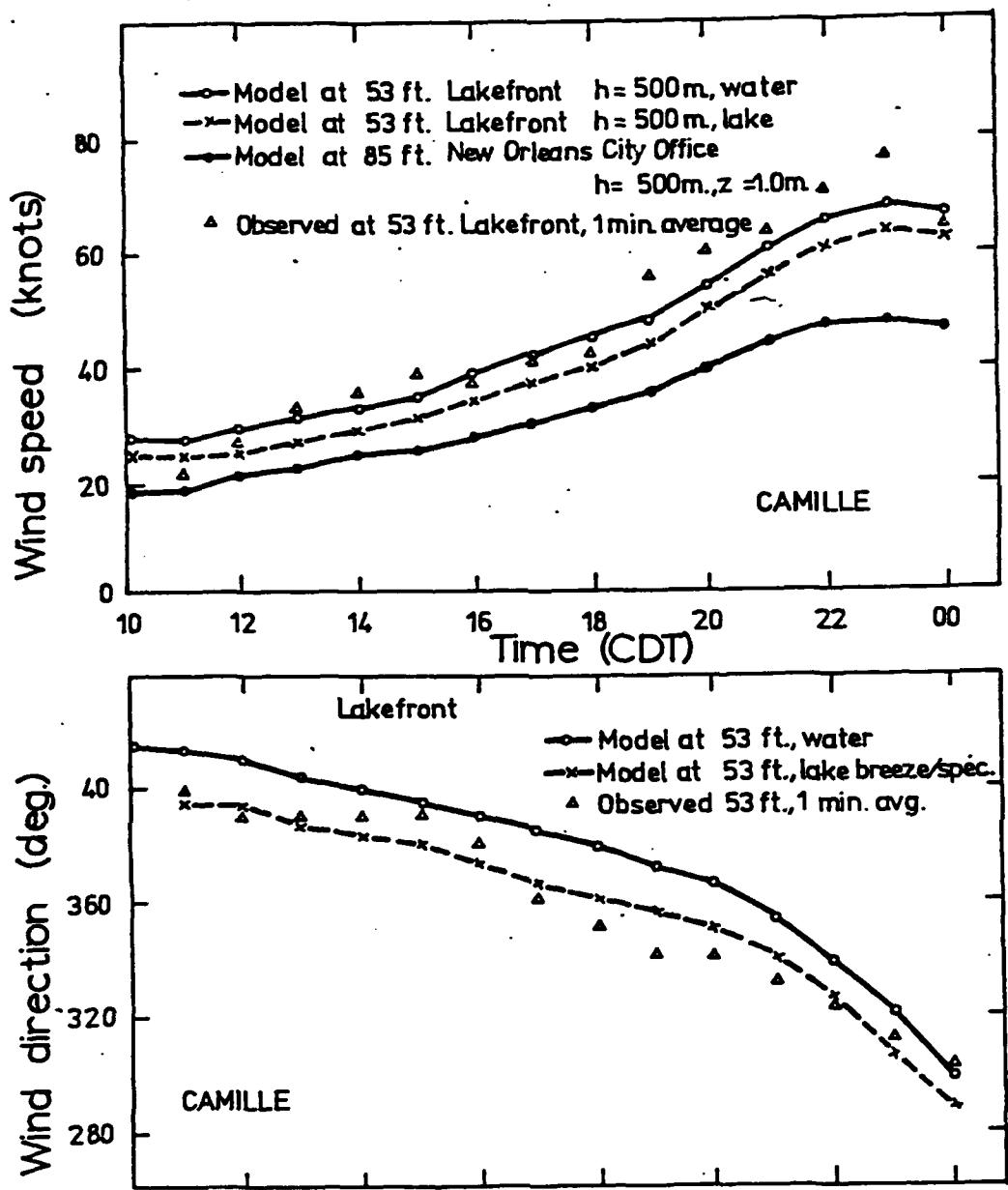


Figure 20. Comparison of measured and modelled wind speed (above) and wind direction (below) at New Orleans Lakefront Airport in Camille. Modelled wind at New Orleans for roughness parameter of 1 m shown

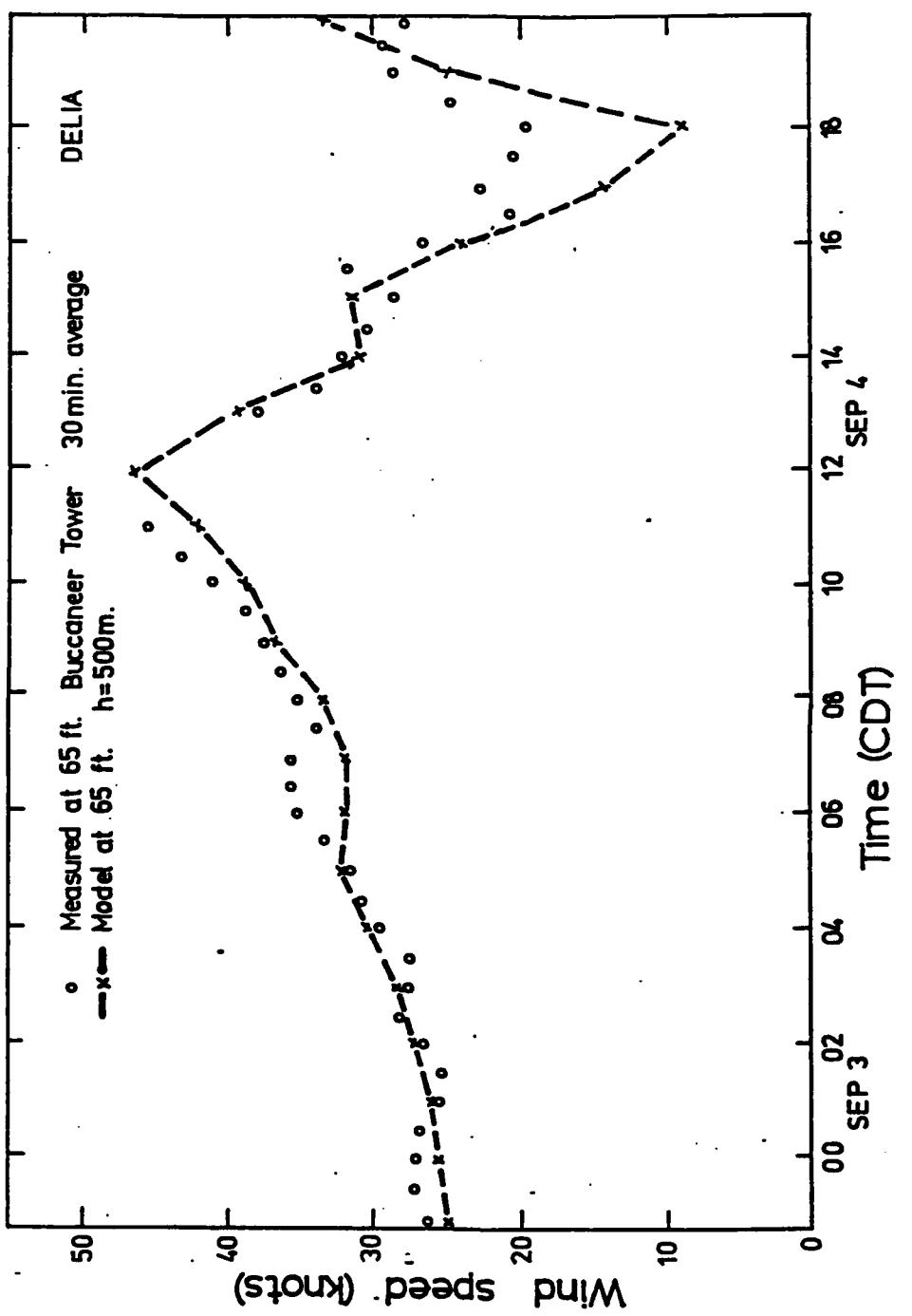


Figure 21. Comparison of measured and modelled wind speed at Buccanneer Tower in Delia

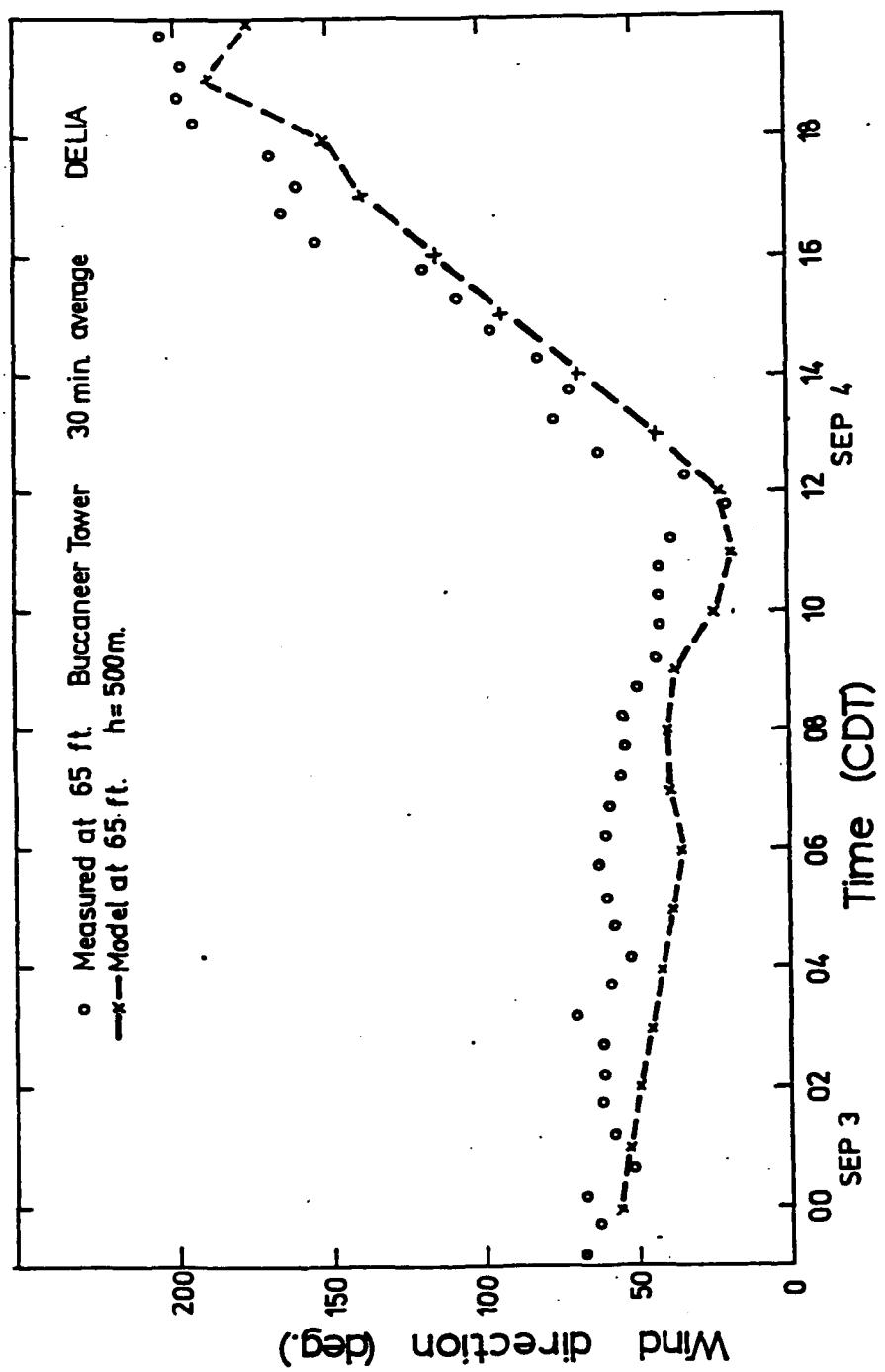
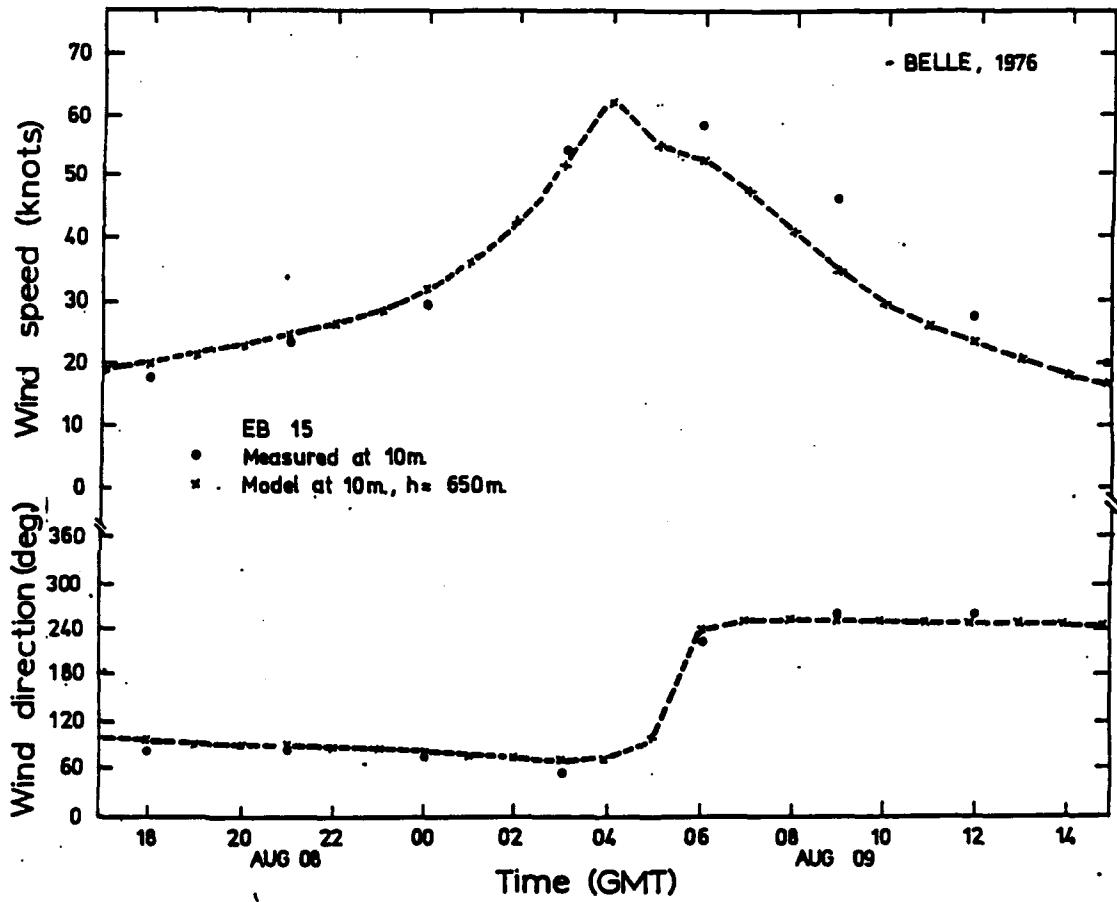


Figure 22. Comparison of measured and modelled wind direction at Buccanneer Tower in DELIA



**Figure 23.** Comparison of measured and modelled wind speed and direction at buoy EB15 in Belle

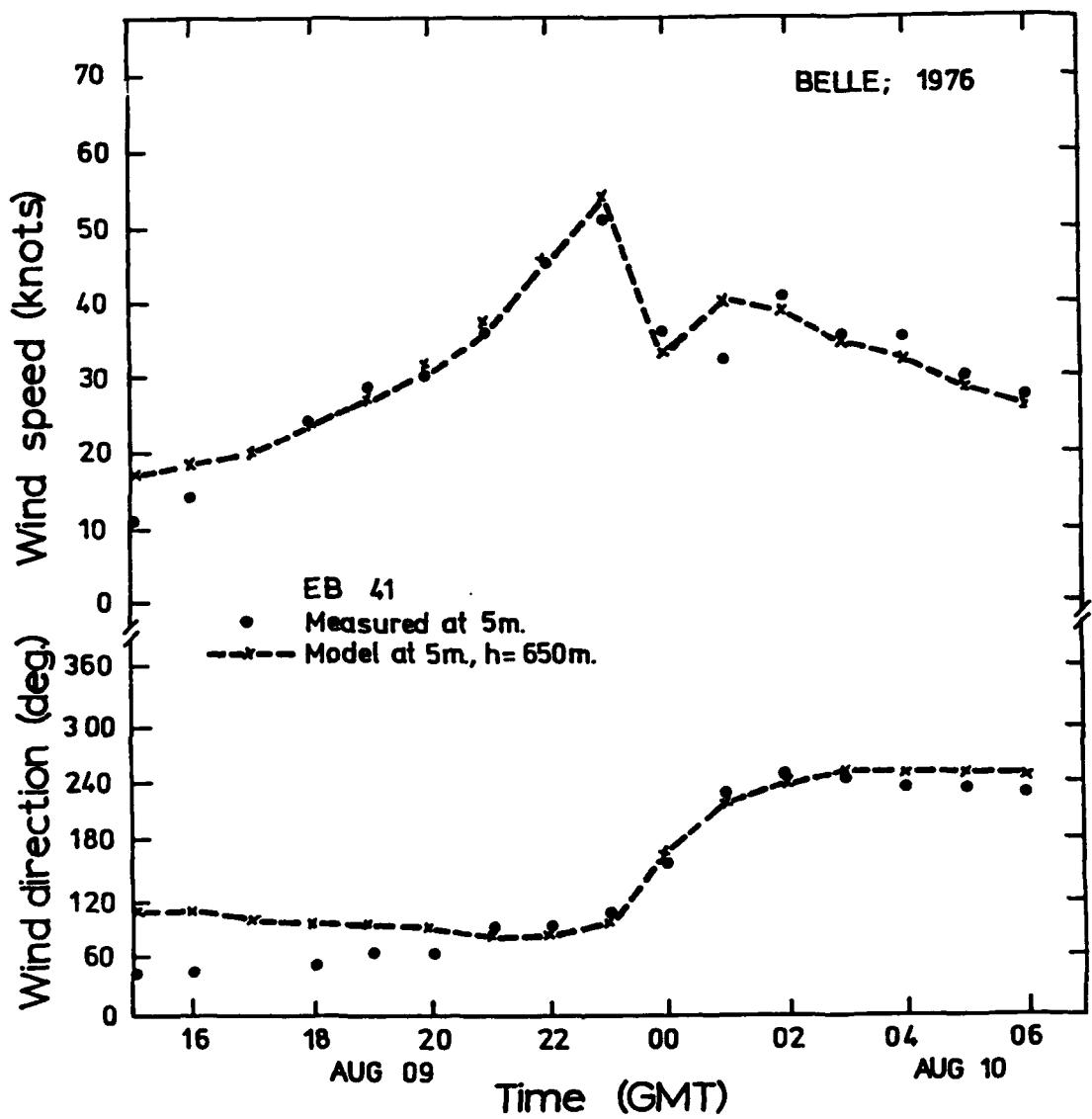


Figure 24. Comparison of measured and modelled wind direction at buoy EB41 in Belle

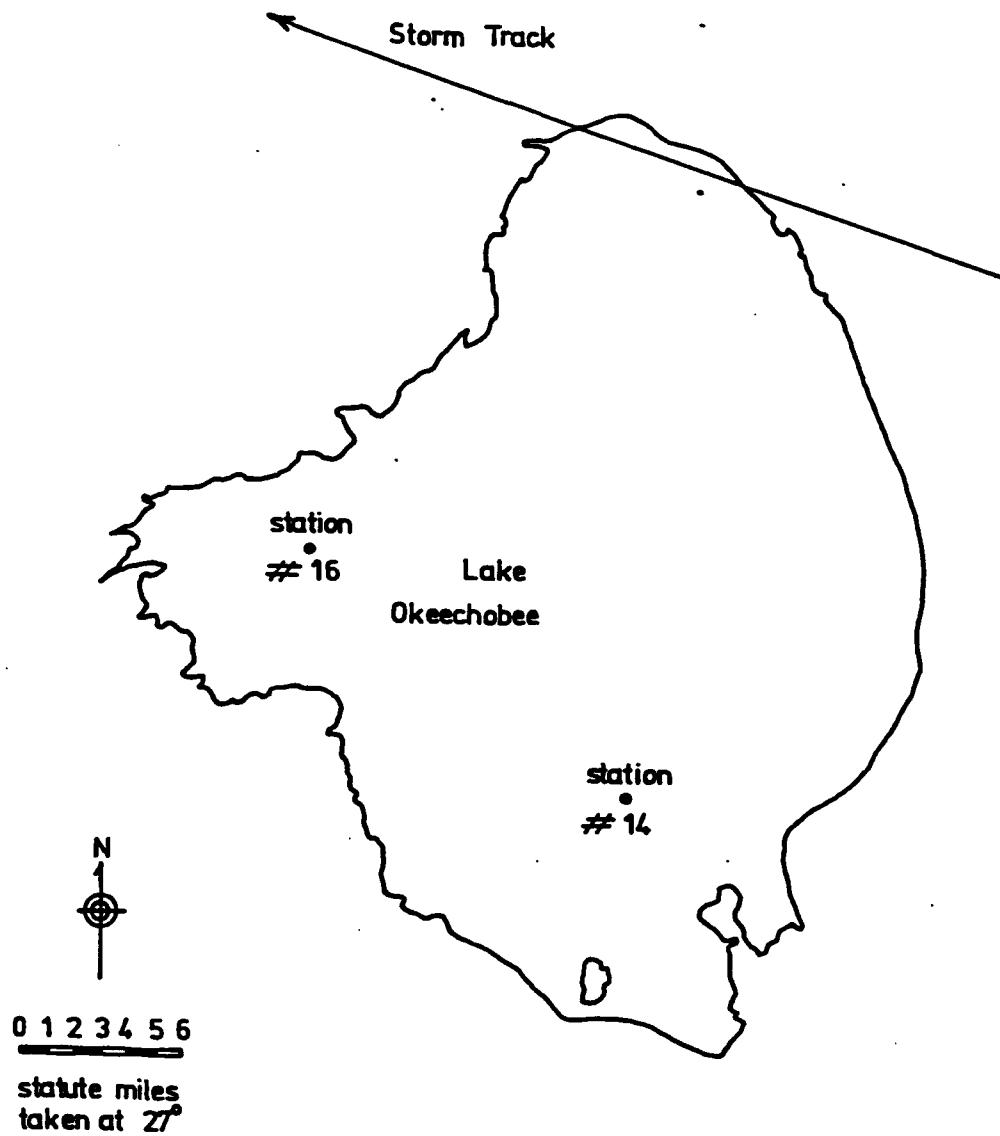


Figure 25. Location of measurement stations in Lake Okeechobee and path of the 1949 Lake Okeechobee hurricane

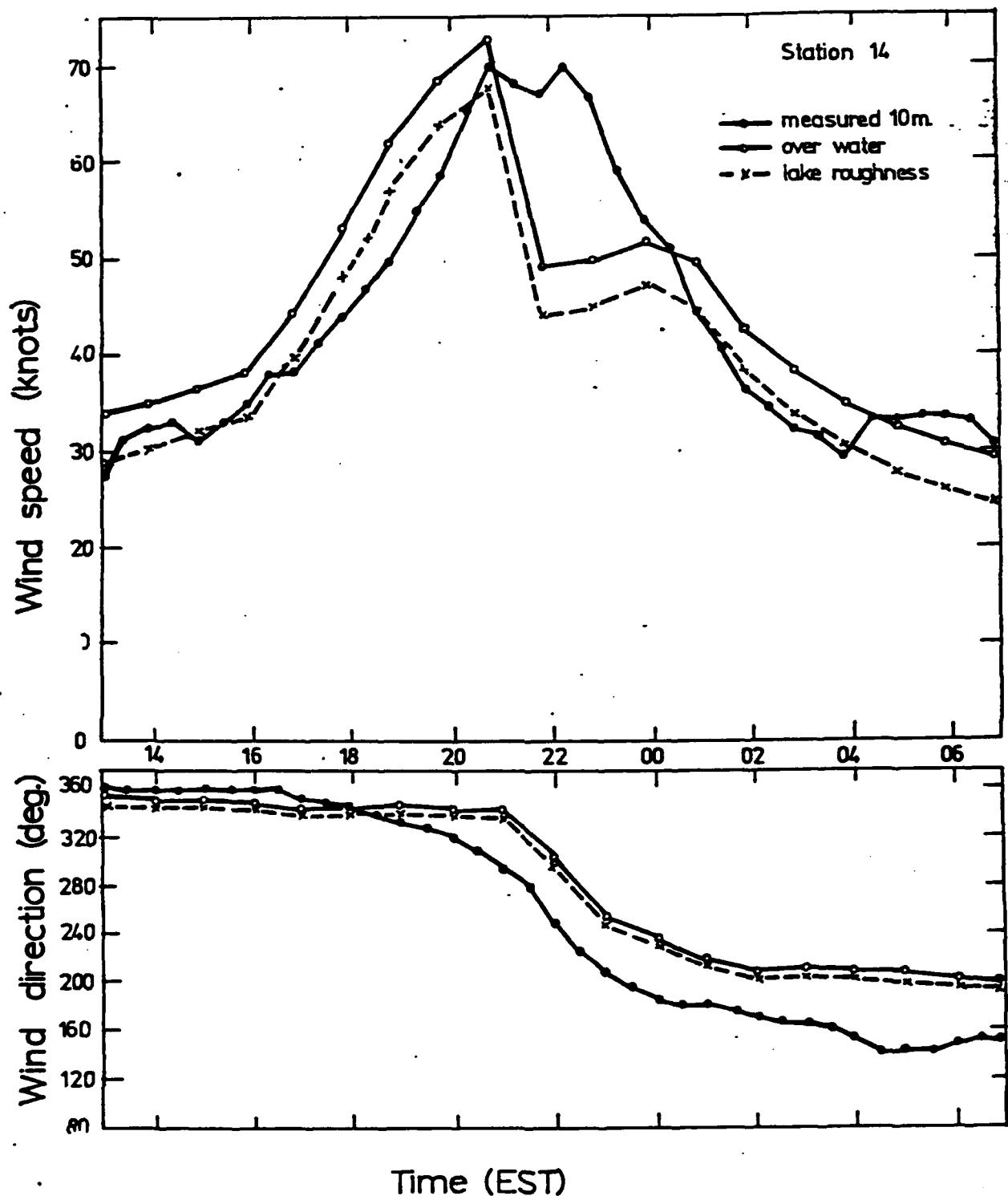


Figure 26. Comparison of measured and modelled wind speed and direction at station 14 in the 1949 Lake Okeechobee hurricane

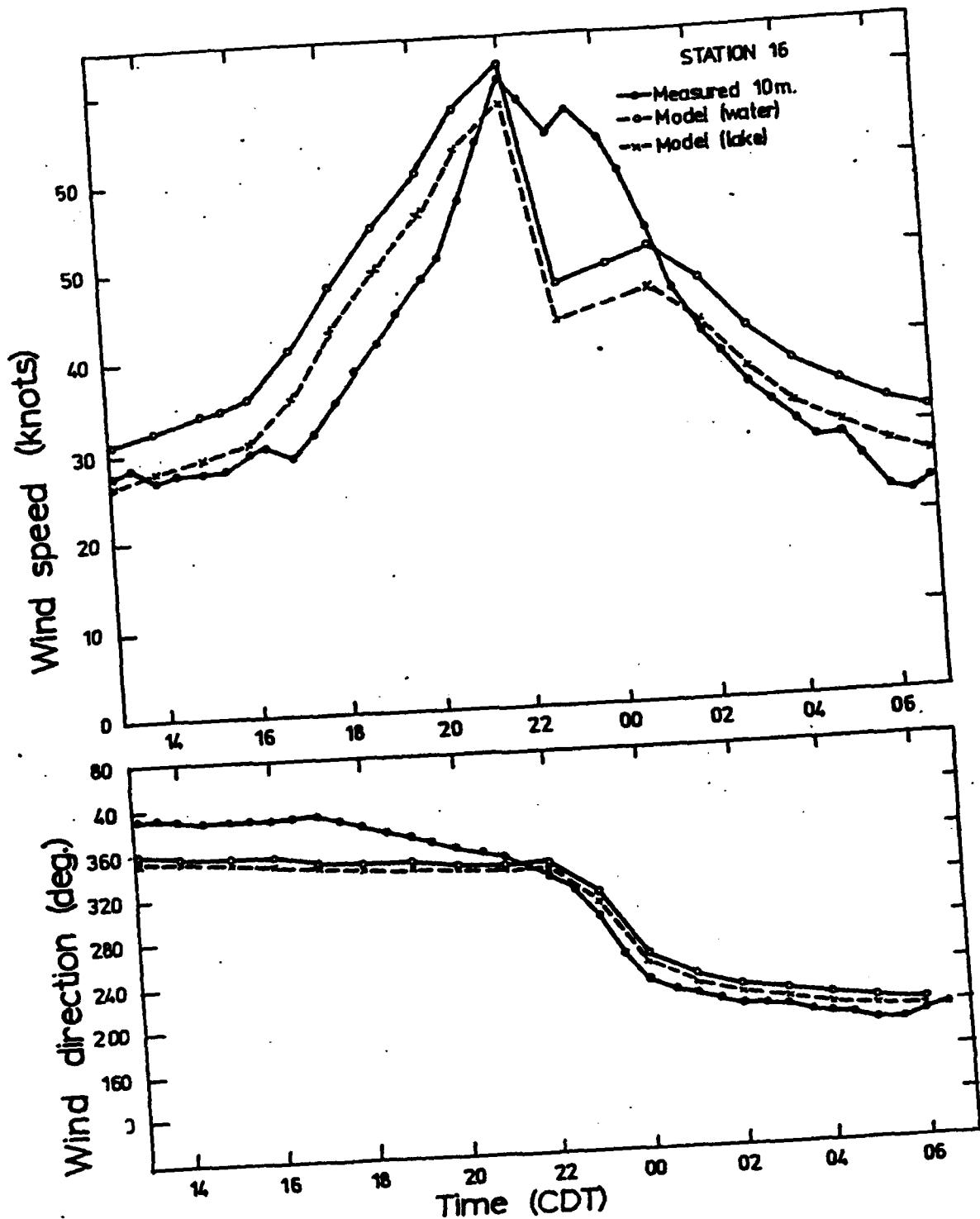


Figure 27. Comparison of measured and modelled wind speed and direction at station 15 in the 1949 Lake Okeechobee hurricane

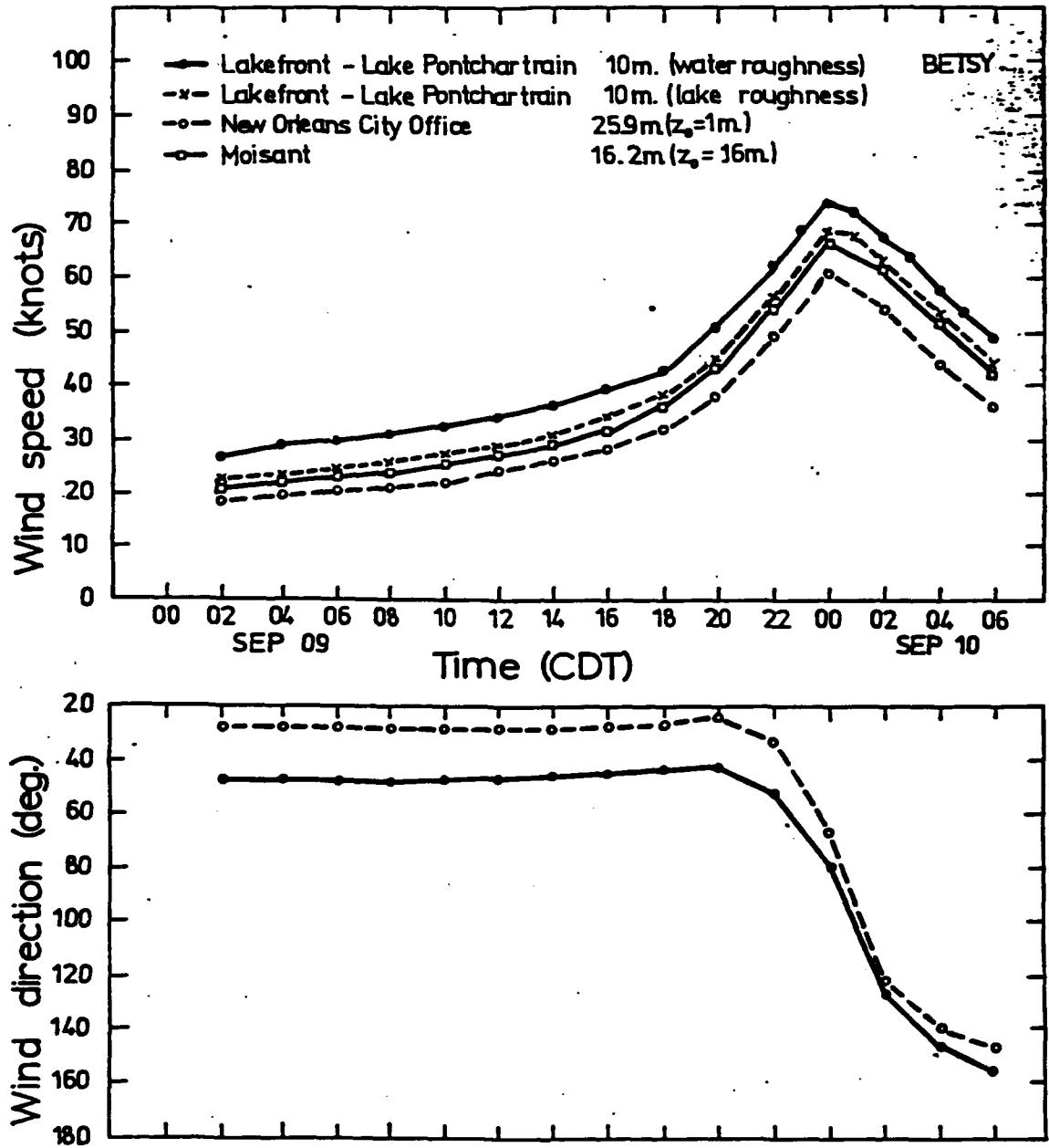


Figure 28. Modelled wind speed and direction for various terrain roughnesses in Hurricane Betsy

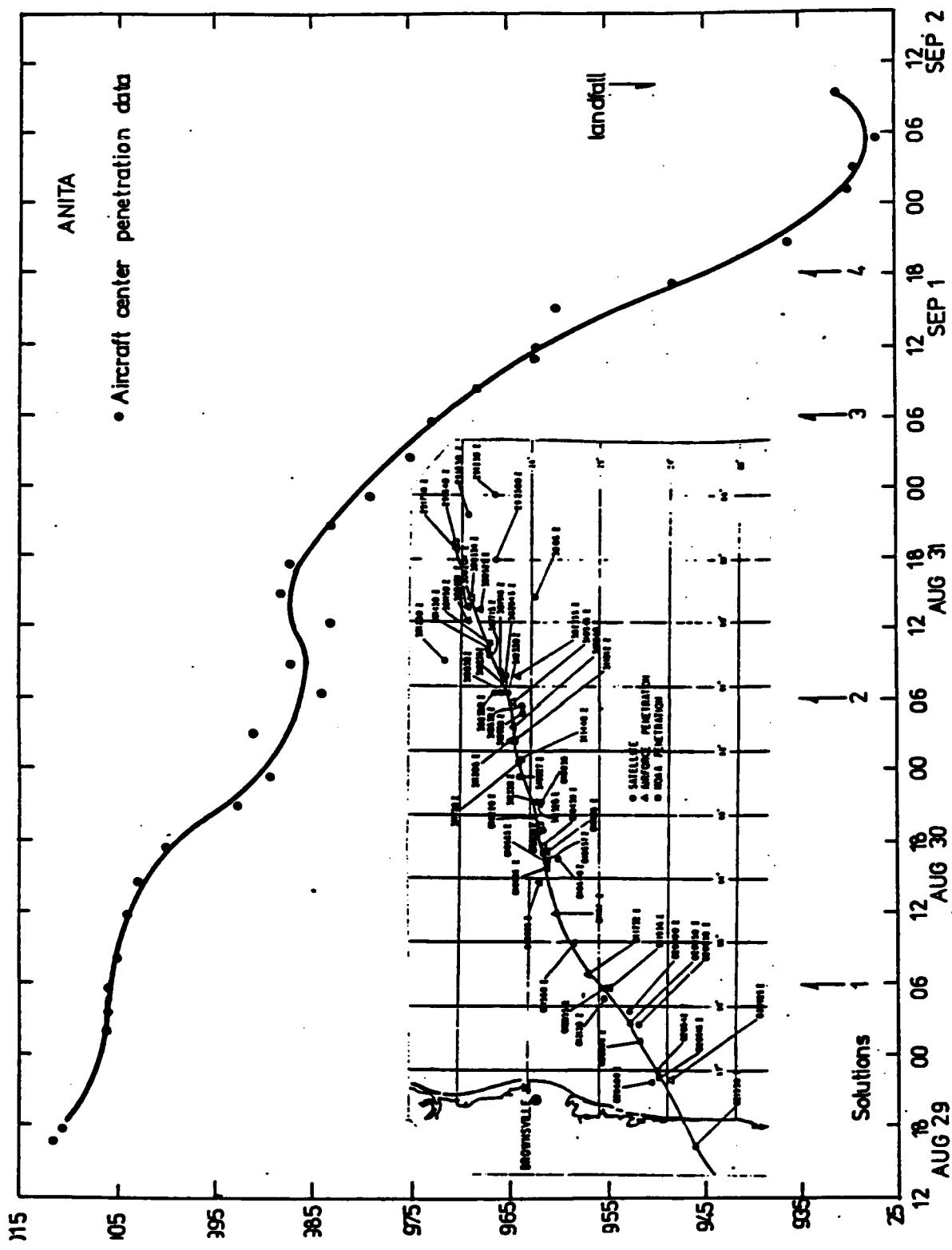


Figure 29. Path and variation of eye pressure with time in Hurricane Anita

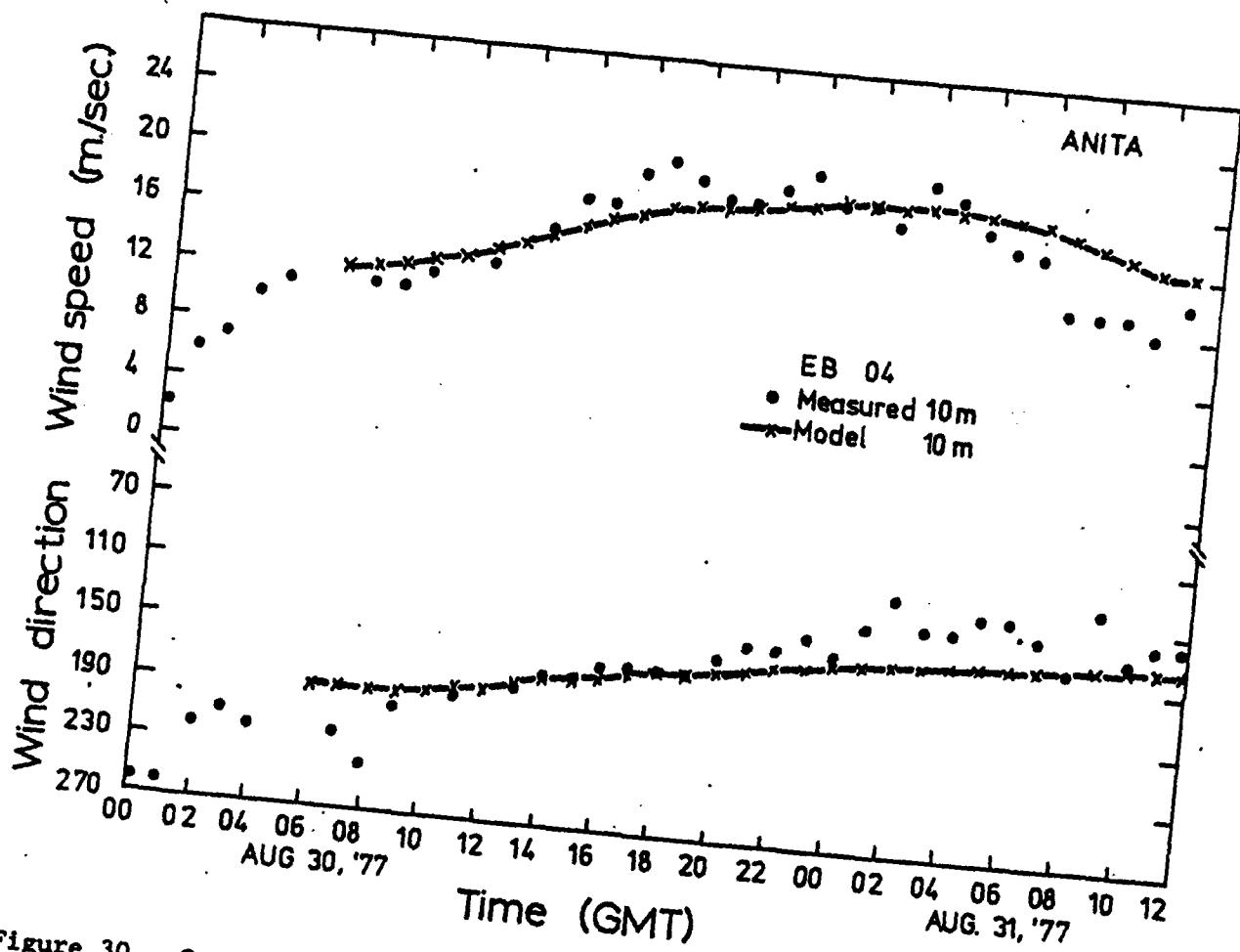


Figure 30. Comparison of measured and modelled wind speed and direction at buoy EB04 in Hurricane Anita

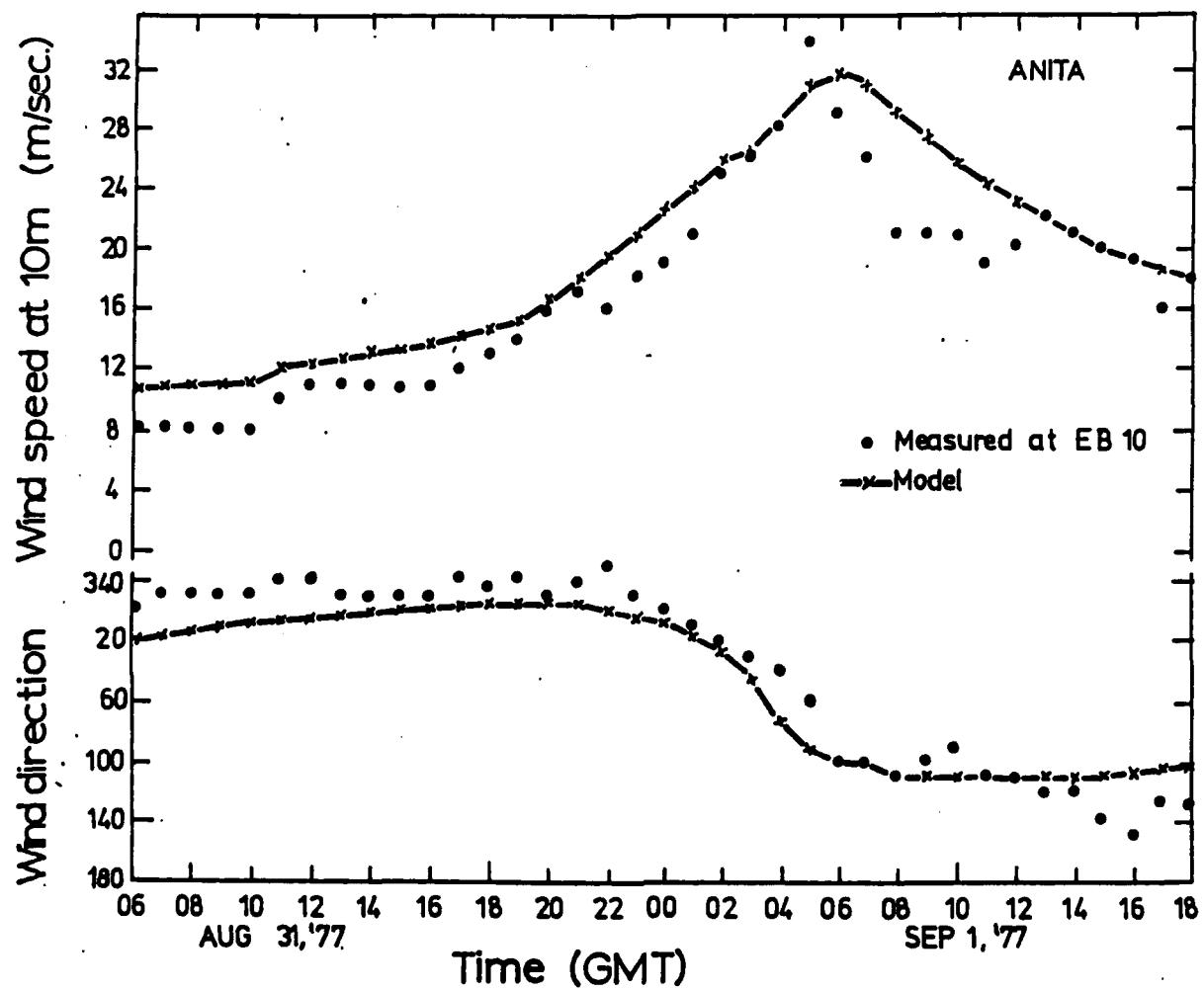


Figure 31. Comparison of measured and modelled wind speed and direction at buoy EB71 in Hurricane Anita

**APPENDIX A: NAMELIST**

1. Source. The material below is abridged from §6.4 of "Sperry Univac Series 1100, FORTRAN V Level 4 R1, programmer reference", edition of April 1979.

2. The nonexecutable NAMELIST statement and the associated forms of the formatted input/output statements provide a simplified means of transmitting an annotated list of data to and from peripheral units. The input/output statements take the form

READ (unit, x)  
and  
WRITE (unit, x)

where x is a namelist name. The list of items in the namelist is used on input to specify those items which may have their values defined in the records to be read. Not all items of the namelist list need be used in the input records nor must the input fields be in the same order as the list items. On output, each list item of the designated namelist is formatted in a standard fashion for output in the order specified by the list.

3. Namelist Statement. The general form of the statement is

NAMELIST /X/A,B,...,C/Y/D,E,...,F/Z/G,H...,I

where X, Y, Z,... are namelist names and A, B, C, D,... are simple variables, subscripted variables, or array names. An array must be dimensioned before appearing in a namelist. The following rules apply to defining and using a namelist:

- a. A namelist name consists of from one to six alphanumeric characters, the first of which must be alphabetic.
- b. Within a NAMELIST statement, a namelist name is enclosed in slashes. The list of variables associated with a namelist name ends when a new namelist name enclosed in slashes is encountered or with the end of the NAMELIST statement.
- c. A namelist name may be defined only once in a routine by its appearance in a NAMELIST statement. In the routine in which it is defined, a namelist name may appear only in

input/output statements and in the defining NAMELIST statement.

- d. A namelist name must not be the same as any other name in the routine in which it appears.
- e. A variable name, array element name, or an array name may be assigned to one or more namelist names. Array names must have been previously declared.
- f. The subscript(s) of an array element must consist of constants.

4. Namelist Input. The READ (i,X) statement causes the records that contain the input data for the variables and arrays that belong to the namelist name X to be read from input unit i. For READ (i,X) , the first character in each data record to be read is always ignored. The second character of the first record of a group of data records to be read must be a \$ immediately followed by the namelist name and a blank. The remainder of the first record and following records may contain any combination of the legal data items which are separated by commas (a comma after the last data group is ignored). The last input record is terminated by a blank followed by \$END.

5. The forms the data items may take are:

- a. Variable Name= Constant. Variable name is a simple variable name.
- b. Subscripted Variable=Constant. The array element appears in the NAMELIST statement. The subscripts in the input record must be constants.
- c. Array Name=Set of Constants. The set is represented by constants separated by commas, or k\*constant may represent k constants (k is an unsigned integer). The number of constants must be less than or equal to the number of elements in the array.

6. Constants used in the data items may take any of the following forms:

- a. Integer constants.
- b. Real constants. These are written with a decimal point, and, optionally, they are written with an exponent consisting of E , + , or - , or a combination of E and sign (for example, E+2, E2, +2).
- c. Hollerith constants. These are written as nHhhh...h where hh...h is a string of n alphanumeric characters, including blanks. Six characters can be stored in one

location. If less than six characters remain they are stored left-justified with the rest of the computer word filled out with blanks.

7. Hollerith constants may only be associated with integer or real variables. The data items that appear on the input records need not appear in the same order as the corresponding variable or array names in the namelist. All variable or array names of the namelist need not have a corresponding data item in the input records; if none appears, the contents of the variable or array are unchanged. Names that are equivalenced to these names may not be used in the input records unless they are part of the namelist. Blanks must not be embedded in a constant or a repeat constant field, but may be used freely elsewhere in a data record. The name of an array and the value of its first elements must appear on the same record. The last item on each record that contains data items must be a constant followed by a comma. The comma is optional in the record that contains or precedes the \$END sentinel.

8. Namelist Output. The WRITE (i,X) statement causes all names of variables and arrays (as well as their values) that belong to the namelist name X to be written on the output unit i. In the WRITE (i,X) statement, all variables and arrays, and their values belonging to the namelist name, are written out according to their types. The output data is written such that:

- a. The name of a variable and its value are written on one line.
- b. The name of an array is written, with the values of the elements of the array written in a convenient number of columns, in the order of the array in main storage, that is, with the left dimension [varying fastest].
- c. The data fields are large enough to contain all the significant digits.
- d. The output can be read by an input statement referencing the namelist name.

**APPENDIX B: PROGRAM LISTING**

```

SUBROUTINE AANGEL (I20)
C   CONVERTS U,V TO SPEED(CKN), DIRECTION.
C   COMPUTES SPEED AT 20 MTRS ALTITUDE IF I20 NE 0.
COMMON /C3/ U(21,21,5),V(21,21,5),UN(21,21,5),VN(21,21,5),
1  UX(21,21,5),UY(21,21,5),UW(21,21,5),ANG(21,21,5)
2  LW(21,21,5)

COMMON
/C4/ CUR(100,2),UX(100,2),TURN(100,2)
/C5/ FLAT,PTH,DTH(2),HH,20LAND,LS,VV(100),UX(3),UV(3),
S  DUV(3),K35,K2,G,6A,DE,N,VY2,HL,K123,20,ZLOG,AM,BM,CM,FF
REAL K35,K2
DO 1000 NEST=1,5
DO 1000 I=1,21
DO 1000 J=1,21
C COMPUTE VTN
VTN(I,J,NEST)=SQR(T(VN(I,J,NEST))**2+VN(I,J,NEST)**2)
IF(VTN(I,J,NEST) .GT. 860,840,860
840 ANG(I,J,NEST)=0.0
GO TO 1000
860 AN=57.29578*ATAN2(VN(I,J,NEST),UN(I,J,NEST))
ANG(I,J,NEST)=AH0D(270.-AN,360.)
C REDUCE SPEED TO 20 MTRS IF I20 NE 0
C REDUCE SPEED TO KNOTS IF I20 EQ 0
IF (I20 .NE. 0) GO TO 900
VTN(I,J,NEST) = VTN(I,J,NEST)*(3600./1852.)
GO TO 1000
900 SPP = 1.25*VTN(I,J,NEST)
KPP = SPP
SPP = SPP-KFP
LS = LW(I,J,NEST)
IF (KPP .NE. 0) GO TO 910
UXX = VTN(I,J,NEST)*UXV(I,LS)
TWIST = TURN(I,LS)
GO TO 930

```

```

910 IF (KPP .GE. 100) GO TO 920
      UXX = VTN(I,J,NEST)*((1.-SPP)*UXV(KPP,LS)+SPP*UXV(KPP+1,LS))
      TWIST = ((1.-SPP)*TURN(KPP,LS)+SPP*TURN(KPP+1,LS))
      GO TO 930
920 UXX = VTN(I,J,NEST)*UXV(100,LS)
      TWIST = TURN(100,LS)
930 Z20 = 19.5/20 AND
      IF (LS .EQ. 2) Z20 = 19.5/(6A*UXX*2)
      VTN(I,J,NEST) = 3600./1852.*AMIN1
      $ (UXX/K35*ALOG(Z20),VTN(I,J,NEST))
      ANG(I,J,NEST) = ANG(I,J,NEST)+7.29578*TWIST
      IF (ANG(I,J,NEST).LT.0.) ANG(I,J,NEST) = ANG(I,J,NEST)+360.
1000 CONTINUE
      RETURN
      END

```

## SUBROUTINE ABCC

COMMON

```

      S/C4 / CDR(100,2,2),UX,V(100,2),TURN(100,2),
      S/C5 / FLAT,PTH,D1H(2),HH,201AND,LS,VV(100),UX(3),UW(3),
      S/UW(3),K35,K2,G,GA,DEN,VV2,HL,K123,20,2LOG,A,M,BM,CH,FF
      REAL K35,K2
      IF (LS .EQ. 1) GO TO 53
      Z0 = GA+UX(K123)*2/HH
      2L06 = ALOG(20)
      53   HL = -HH*DEN/(UX(K123)*2*PTH*(ZL06+CH))
      IF (HL+2) 54,55,56
      HLOG = ALOG(-HL)
      AH = AHIN1(HLOG+1.5,-.875+ZL06)
      54   YH = 1.8*EXP(2*HL)
      CH = AHIN1(HLOG+3.7,-.875+ZL06)
      60 TO 60
      55   AH = 2*1931472
      BH = 1.2065761
      CK = 4.3931472
      GO TO 60

```

```

56 IF (HL-2.) 57,58,59
57 AM = 1.3865736-.4032868*HL
      CM = 2.5465736-HL*.923868
      BM = 1.953288+.37335598*HL
      CH = .7
      DM = 2.70
      GO TO 60
58 AM = .58
      GO TO 60
59 YLOG = 2L06+4.7
60 DUV(K123) = UX(K123)+2*X2*(2L06+AM)+2*BM+2)
61 UV(K123) = UV(K123)-V2
62 RETURN
63 END
64
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SUBROUTINE ABCCC
COMMON /C57/ PTH,DTH,HH,ZCOFF(3,5),LAKE,VV(100),UX(3),UV(3),
      DUV(3),K35,K2,6,6A,DEN,VV2,HL,K123,20,ZLOG,AH,BM,CH,UXX(100,6),
      TURN(100,6),COST(100),SINT(100)
      REAL K2,K35
      20 = 6A*UX(K123)*2/HH
      ZLOG = ALOG(ZD)
      HL = -HH*DEN/(UX(K123)*2*PTH*(ZLOG+CH))
      IF (HL+2.) 54,55,56
      53   HL0G = ALOG(-HL)
      AH = AMIN1(HL0G+1.5,-.875*ZLOG)
      BH = 1.*B*EXP(.2*HL)
      CH = AMIN1(HL0G+3.7,-.875*ZLOG)
      GO TO 60
      55   AH = 2.1931472
      BH = 1.2065761
      CH = 4.3931472
      GO TO 60
      56   IF (HL-2.) 57,58,59
      57   AH = 1.3665736-.4032868*HL
      BH = 1.953288+.37335598*HL
      CH = 2.5465736-HL+.9232868
      GO TO 60
      58   AH = .58
      BH = 2.70
      CH = .7
      GO TO 60
      59   YL05 = 2LCG+4.7
      HL = .25*(YL0G+SORT(YLOG**2+8.*HH*DEN/(UX(K123)**2*PTH)))
      AH = AMAX1(2.5-.96*HL,.99.)
      BH = AMIN1(1.1+.8*HL,.99.)
      CH = 4.7-2.*HL
      UV(K123) = UX(K123)**2/K2*((ZLOG+AH)**2+BH**2)
      UUV(K123) = UV(K123)-VV2
      RETURN
      END

```

```

C SUBROUTINE BLOWUP PRODUCES THE WIND FIELD FOR MOVING VORTEX IN THE PLANETARY
C BOUNDARY LAYER. ITS U(EASTWARD) AND V(NORTHWARD) COMPONENTS
C ARE IN ARRAYS UN, VN RESPECTIVELY UPON EXIT.
C SGW IS MAGNITUDE OF SURFACE GEOSTROPHIC WIND
C AN1 IS ANGLE BETWEEN SGW AND EAST
C ST12 IS DISTANCE IN KM FROM AXIS OF STORM TO HALF MAGNITUDE OF SGW
C (UG, VG) IS THE SFC GEOSTROPHIC WIND.
C CS IS SPEED OF STORM MOVEMENT.
C COMMON/C1/NAME,NSMAP,D1(2),F,SGW,AN1,UC,VC,UG,VG,CS,NM,IB
C ST12
COMMON /C3/ U(21,21,5),V(21,21,5),UN(21,21,5),VN(21,21,5),
1 ,PX(21,21,5),PY(21,21,5),VTN(21,21,5),ANG(21,21,5)
2 ,LY(21,21,5)

COMMON
$ /C4/ CDR(100,2,2),UXV(100,2),TURN(100,2)
$ /C5/ FLAT,PTH,DTI(2),MH,ZOLAND,LS,UV(100),UX(3),UV(3),
$ DUV(3),K35,K2,G,GA,DEN,VV2,HL,K123,Z0,ZLOG,AM,BM,CM,FF
REAL K35,K2
DIMENSION LEVEL(16)
REAL COH2(2)
DATA LEVEL/1.5,1.2,1.3,1.2,1,4,1.2,1.3,1.2/
DEGG = AN1/57.29578
UG=SGW*CUSIDE(GG)
VG=SGW*SIN(DEGG)
CALL SHURE
CALL PXYH
IF (IB,NC,0) CALL OUTOUT(0,NAME,4HINIT,NSMAP)
SPP = 1.25*SGW
KPP = SPP
SPP = SPP-KPP
DO 1020 LS = 1,2
  COH2(LS) = (((1.-SPP)*UXV(KPP,LS)+SPP+UXV(KPP+1,LS))*2/HH)*2
CONTINUE
CS=SQRT((UC**2+VC**2)
DO 1026 I = 1,21

```

```

DO 1026 J = 1,21
IF (I .EQ. 1) GO TO 1025
IF (I .EQ. 21) GO TO 1025
IF (J .EQ. 1) GO TO 1025
IF (J .EQ. 21) GO TO 1025
60 TO 1026

1025 CONTINUE
      LS = L(L,I,J,5)
      AA = CDH2(LS)*(PX(I,J,5)+2*PY(I,J,5)+2)
      F2=F+2/2.0
      SK=AA/(F2+SQRT(F2+AA))
      BK=SQRT(SK)
      V(I,J,5)=(F+PX(I,J,5)-BK*PY(I,J,5))/(F+2+SK)
      U(I,J,5)=-PY(I,J,5)/(F-V(I,J,5)+BK/F)
      V(I,J,5)=V(I,J,5)-VC
      U(I,J,5)=U(I,J,5)-UC
1026 CONTINUE
      DO 1026 I = 1,21
      DO 1028 J = 1,21
      UN(I,J,5)=U(I,J,5)
      VN(I,J,5)=V(I,J,5)
      C
1028 CONTINUE
      DO 1030 NEST = 1,5
      DO 1030 I = 1,21
      DO 1030 J = 1,21
      PX(I,J,NEST)=PX(I,J,NEST)-F+VC
      PY(I,J,NEST)=PY(I,J,NEST)+F+UC
      C
      DO 1430 K=1,NM
      KCK = MOD(K,16)
      CALL COMOUT (LEVEL(KCK+1))
      1430 CONTINUE
      CALL OUTFLO
      C
      SOLUTION OF WIND FIELD ON COORDINATE SYSTEM MOVING
      C
      WITH STORM IS NOW COMPLETE.

```

C COMPUTE SOLUTION WITH RESPECT TO SEA SURFACE.  
DO 1450 NEST=1,5  
DO 1450 I=1,21  
DO 1450 J=1,21  
UN(I,J,NEST)=UN(I,J,NEST)+UC  
VN(I,J,NEST)=VN(I,J,NEST)+VC  
1450 CONTINUE  
CALL OUTQUT(1,NAME,4HSNAP,NSMAP)  
CALL OUTGUT(0,NAME,4HSNAP,NSMAP)  
RETURN  
END

```

SUBROUTINE BREEZE
REAL LA0,LA1,L00,L01,LB,LL
COMMON /C57/ PTH,PTH,HH,ZCOEFF(3,5),LAKE,UV(100),UX(3),UV(3),
$ DUV(3),K35,K2,G,GA,DEN,VV2,HL,K123,Z0,ZL06,AM,BM,CM,UXV(100,6),
$ TURW(100,6),COST(100),SINT(100)

REAL K2,K35
COMMON /D1/LA0,L00,ROT,LA1,L01,DX,STHT,LANSEA,W1,TH1,D,AL,USt
COMMON /D2/ XX(21,21,10)
EQUIVALENCE (U1,X,Y),(V1,XY(2)),(UX,X,USt) INPUT (RADIANs)
C LA0,L00 ARE LAT, LON OF EYE (PT0) INPUT (RADIANs)
C LA1,L01 ARE LAT, LON OF POINT AT WHICH WIND IS WANTED (PT1)-INPUT(RAD)
C LONGITUDE IS POSITIVE WEST.
C DX IS GRID SPACING OF INNERMOST NESt - INPUT (KILOMETERS)
C W1 IS WIND SPEED AT PT1,20 METERS - OUTPUT (KNOTS)
C D IS DISTANCE BETWEEN PT0 AND PT1 - OUTPUT (KM)
C ROT IS ANGLE FROM TRUE NORTH TO Y-AXIS OF NESTED RECTANGULAR
C GRID WIND FIELD (CLOCKWISE,DEGREES).
C TH1 IS DIRECTION TO WHICH WIND BLOWS, CLOCKWISE FROM SOUTH(DEG)
C AL IS BEARING OF POINT 1 FROM POINT 0, CLOCKWISE FROM NORTH(DEG)
C LANSEA IS TERRAIN CODE AT POINT 1
C 1 IS OCEAN, 2 TO 6 ARE VARIOUS GROUNDS AND LAKES
C DIMENSION XY(2)
NAMELIST/DBG2/1,J,F1,F2,XY,SPP,LANSEA,UX,X,TWIST,STHT,GA,K35
HAV(X) = (SIN(.5*X))**2
AHAV(X) = 2.*ASIN(SQRT(X))
SARU(X) = SQRT(ABS(X))
LL = LA0+LA1
LB = LA0-LA1
R = AHAV(HAV(LL)+COS(LA0)*COS(LA1))*HAV(LL-LO1)
IF (R.EQ.0.) GO TO 16
IF (R.LT.-1E-4) AB = ATAN2((LO0-LO1)*COS(.5+LL),-LB)
AB = AB+.5
IF (R.GE.+1E-4) AB = ATAN2(SQRU((COS(.5*(LL+R))+SIN(.5*(R+LB))),1
$ SQRU((COS(.5*(LL-R))+SIN(.5*(R-LB))))),
IF ((LO1 .GT. LO0) .AND. AB = -AB
AL = 2.*57.29578*AB

```

```

1 IF (AL .LT. 0.) AL = AL + 360.
2 DE = (AL-ROT)/57.29578
3 D = R*1.052*3437.7468
4 AA = 10.*D*Dx
5 IZ = 1
6 XY(1) = D*SIN(DE)
7 XY(2) = D*COS(DE)
8 AB = AMAX1(ABS(XY(1)),ABS(XY(2)))
9 IF (AB .LT. AA) GO TO 11
10 AA = AA+AA
11 IZ = IZ+1
12 GO TO 10
13 IF (IZ .LE. 5) GO TO 12
14 U1 = 0.
15 V1 = 0.
16 W1 = 0.
17 TH1 = 0.
18 RETURN
19 DO 13 IA = 1,2
20 KY(IA) = (AA*XY(IA))*10./AA
21 I = XY(1)
22 J = XY(2)
23 F1 = XY(1)-I
24 F2 = XY(2)-J
25 G11 = (1.-F1)*(1.-F2)
26 G12 = (1.-F1)*F2
27 G21 = F1*(1.-F2)
28 G22 = F1*F2
29 DO 14 IA = 1,2
30 IB = 5*IA+IZ-5
31 XY(IA) = G11*XX(1+IA,J+IA)+G12*XX(1+IA,J+2,IA)+G21*
32 *XX(1+2,JA+IA)+G22*XX(1+2,JA+2,IA)
33 W1 = U1*U1+V1*V1
34 IF (W1*EC .GT. 0.) GO TO 110
35 W1 = SQRT(W1)

```

```

C REDUCE W1 TO STATION HEIGHT
900   SPP = 1.25*W1
      KPP = SPP
      SPP = SPP-KPP
      IF (KPP .NE. 0) GO TO 910
      UXX = W1*UXV(1,LANSEA)
      TWIST = TURN(1,LANSEA)
      GO TO 930
910   IF (KPP .GE. 100) GO TO 920
      UXX = W1*((1.-SPP)*UXV(KPP,LANSEA)+SPP*UXV(KPP+1,LANSEA))
      TWIST = ((1.-SPP)*TURN(KPP,LANSEA)+SPP*TURN(KPP+1,LANSEA))
      GO TO 930
920   UXX = W1*UXV(100,LANSEA)
      TWIST = TURN(100,LANSEA)
930   IF (LANSEA .EQ. 1) Z0 = CA*UST**2
      IF (LANSEA .NE. 1) Z0 = 2COEFF(1,LANSEA-1)/UST+
      + 2COEFF(2,LANSEA-1)*UST**2+2*COEFF(3,LANSEA-1)
      Z20 = STHT/Z0
      C WRITE(6,DBG2)
      S W1 = (3600./1852.)*AMIN1
      S UXX/K35*ALOG(220),W1)
      TH1 = (ATAN2(-U1,-V1)+TWIST)*57.29578+ROT
      IF (TH1 .LT. 0.) TH1 = TH1+360.
      RETURN
16    U1=XX(11,11,1)
      V1=XX(11,11,6)
      D=0.
      AL=0.
      GO TO 15
      END

```

SUBROUTINE BREEZE

```

C     *** THIS VERSION OF BREEZE USES TURNING FROM TERRAIN TYPE 6 ***
C     *** FOR TERRAIN TYPE 2
C
C     REAL LA0,LA1,L00,L01,LB,LL
C     COMMON /C57/ PTH,DTH,HH,ZCOEFF(3,5),LAKE,VV(100),UX(3),UV(3),
C     $ DUV(3),K35,K2,G,GA,DEN,VV2,HL,K123,20,2LOG,AM,BM,CM,UXX(100,6),
C     $ TURN(100,6),COST(100),SINT(100)
C
C     REAL K2,K35
C     COMMON /D1/LA0,LUU,ROT,LAI,L01,DX,STHT,LANSEA,W1,TH1,D,AL,UST
C     COMMON /D2/ X(21,21,10)
C     EQUIVALENCE (U1,XY), (V1,XY(2)), (UX,X,UST)      INPUT (RADIAN)
C     LA0,L00 ARE LAT,LON OF EYE(PT0)      INPUT (RADIAN)
C     LA1,L01 ARE LAT,LON OF POINT AT WHICH WIND IS WANTED(PT1)-INPUT(RAD)
C     LONGITUDE IS POSITIVE WEST.
C
C     DX IS GRID SPACING OF INNERMOST NEST - INPUT(KILOMETERS)
C     W1 IS WIND SPEED AT PT1,20 METERS - OUTPUT(KNDS)
C     D IS DISTANCE BETWEEN PT0 AND PT1 - OUTPUT(KM)
C     ROT IS ANGLE FROM TRUE NORTH TO Y-AXIS OF NESTED RECTANGULAR
C     GRID WIND FIELD (CLOCKWISE,DEGREES).
C
C     TH1 IS DIRECTION TO WHICH WIND BLOWS, CLOCKWISE FROM SOUTH(DEG)
C     AL IS BEARING OF POINT 1 FROM POINT C, CLOCKWISE FROM NORTH(DEG)
C     LANSEA IS TERRAIN CODE AT POINT 1
C
C     1 IS OCEAN, 2 TO 6 ARE VARIOUS GROUNDS AND LAKES
C
C     DIMENSION XY(2)
C     NAMELIST/DBG2/I,J,F1,F2,XY,SPP,LANSEA,UXX,TWIST,STHT,GA,K35
C
C     HAV(X) = (SIN((5*X))**2
C     AHAV(X) = 2.*ASIN(SQRT(X))
C     SURU(X) = SQRT(ABS(X))
C
C     LL = LA0+LAI
C     LB = LA0-LAI
C     R = AHAV(HAV(LB))*COS(LA0)*COS(LL0-L01))
C     IF (R.EQ.0.) GO TO 16
C     IF (R.LT.-1E-4) AB = ATAN2((LL0-L01)*COS(.5+LL),-LB)
C     AB = AB+.5

```

```

IF (R . GE . 1E-4) AB = ATAN2 (SQRT ((COS (.5*(LL+R)) * SIN (.5*(R+LB))),  

1 SUR (COS (.5*(LL-R)) * SIN (.5*(R-LB))))  

1 IF (LUI . GT . L00) AB = -AB  

AL = 2 . 57 . 29578 * AB  

IF (AL . LT . 0.) AL = AL + 360.  

DE = (AL-ROT)/57.29578  

D = R+1 . 852 . 3437 . 7468  

AA = 10.0 * DX  

IZ = 1  

XY(1) = D * SIN(DE)  

XY(2) = D * COS(DE)  

AB = ABS (XY(1)) * ABS (XY(2))  

10 IF (AB . LT. AA) GO TO 11  

AA = AA+AA  

IZ = IZ+1  

GO TO 10  

11 IF (IZ . LE . 5) GO TO 12  

U1 = 0.  

V1 = 0.  

W1 = 0.  

TH1 = 0.  

110 RETURN  

12 DO 13 IA = 1,2  

13 XY(IA) = (AA+XY(IA))+10./AA  

I = XY(1)  

J = XY(2)  

F1 = XY(1)-I  

F2 = XY(2)-J

```

```

- G11 = (1.-F1)*(1.-F2)
G12 = (1.-F1)*F2
G21 = F1*(1.-F2)
G22 = F1*F2
DO 14 IA = 1,2
IB = 5*IA+IZ-5
XY(IA) = G11*XX(I+1,IB)+G12*XX(I+1,J+2,IB)+G21*
*XX(I+2,J+1,IB)+G22*XX(I+2,J+2,IB)
14 W1 = U1+V1+V1+V1
IF (W1.EQ.0.) GO TO 110
15 W1 = SORT(W1)
REDUCE W1 TO STATION HEIGHT
C SPP = 1.25*W1
KPP = SPP
SPP = SPP-KPP
LSSUB=LANSEA
1 IF(LSSUB.EQ.2) LSSUB=6
IF (KPP .NE. 0) GO TO 910
UXV = W1*UXV(1,LANSEA)
TWIST = TURN(1,LSSUB)
TWIST = TURN(1,LSSUB)
GO TO 930
1 IF (KPP .GE. 100) GO TO 920
UXV = W1+((1.-SPP)*UXV(KPP,LANSEA)+SPP*UXV(KPP+1,LANSEA))
TWIST = ((1.-SPP)*TURN(KPP,LSSUB)+SPP*TURN(KPP+1,LSSUB))
900
910
920
930
B16

```

C WRITE(6,0B62)  
W1 = (3600./1852.)\*AMIN1  
\$ (UX/X/K35\*AL06(220)\*W1)  
TH1 = (ATAN2(-U1,-V1)\*TWIST)\*57.\*2957R\*ROT  
IF (TH1 .LT. 0.) TH1 = TH1+360.  
RETURN  
U1=XX(11,11,1)  
V1=XX(11,11,6)  
D=0.  
AL=0.  
GO TO 15  
END

16

SUBROUTINE CCROSS

COMMON

```

$ /C4/ CDR(100,2,2),UX V(100,2),TUR:(10H,2)
$ /C5/ FLAT,PTH,DTH(2),HH,ZOLAND,L5,VV(100),UX(3),UV(3),
$ DUV(3),K35,K2,G,GA,DEN,VV2,HIL,K123,Z0,7LOG,AM,BM,CM,FF
KEAL K35,K2

```

```
FF = AMAX1(ABS(FLAT),1.832E-6)
```

```
DO 10 IA = 1,100
```

```
VV(IA) = .8*FLOAT(IA)
```

CONTINUE

```
DO 41 LS = 1,2
```

```
DEN = G*K2*DTH(LS)
```

```
DO 40 IA = 1,100
```

```
IB = 101-IA
```

```
VV2 = VV(IB)*2
```

```
IF (IA .NE. 1) GO TO 20
```

```
CN = 2.55
```

```
K123 = 1
```

```
UX(1) = 2.74
```

```
IF (LS .EQ. 1) UX(1) = 3.38
```

```
Z0 = ZOLAND/HH
```

```
ZLOG = ALOG(Z0)
```

```
CALL ABCC
```

```
K123 = 2
```

```
UX(2) = UX(1)*80./SORT(UV(1))
```

```
GO TO 30
```

```
UX(1) = UX(3)
```

```
UV(1) = UV(3)
```

```
DUV(1) = UV(1)-VV2
```

```
K123 = 2
```

```
UX(2) = UX(1)*VV(IB)/VV(1B+1)
```

```
CALL ABCC
```

```
KSTOP = 0
```

```
K123 = 3
```

30

```

31      IF (DUV(1) .EQ. DUV(2)) GO TO 32
      UX(3) = AMAX1(.5 * AMIN1(UX(1),UX(2)),
      $      AMIN1(2 * AMAX1(UX(1),UX(2)),
      $      (UX(1)*DUV(2)-UX(2)*DUV(1))/IDUV(2)-IDUV(1)))
      CALL ABCC
      IF (KSTOP .NE. 0) GO TO 32
      IF (ABS(DUV(3)).LT. .1E-4 * VV2) KSTOP = 1
      UX(1) = UX(2)
      UX(2) = UX(3)
      DUV(1) = DUV(2)
      DUV(2) = DUV(3)
      GO TU 31
      AU = SQRT((2LOG+AH)**2+BH**2)
      UXV(IB,LS) = K35/AB
      AB = UXV(IB,LS)**2/AB
      CDR(IB,LS,1) = (ZLOG+AH)*AB
      CDR(IB,LS,2) = BH*AB
      TURN(IB,LS) = ATAN(BH/(ZLOG+AH))
      40  CONTINUE
      41  CONTINUE
      RETURN
      END

```

```

SUBROUTINE COMOUT (LEVEL)
COMMON/C1/DM1(2),DX,DT,F,DM2(2),UC,VC,DM3(5)
, ST12
COMMON /C3/ U(21,21,5),V(21,21,5),UN(21,21,5),VN(21,21,5)
1 ,PX(21,21,5),PY(21,21,5),VT(21,21,5),ANG(21,21,5),
2 ,LV(21,21,5)

COMMON
$ /C4/ CDR(100,2,2),UXV(100,2),TURN(100,2)
$ /C5/ FLAT,PTH,DTL(12),HH,ZOLAND,LS,VV(100),UX(3),UV(3),
$ DUV(3),K35,K2,6,GA,DE,HL,K123,20,ZLOG,AM,BM,CH,FF
DIMENSION HKL(21,21),DRAG(2)
DATA E1,E2/0.5,0.5/
DATA VONK/0.4/
C
SQQ IS APPROXIMATION TO SORT OF SUM OF SQUARES
SQQ(A,B) = (ABS(A)+ABS(B))/2.51 + (ABS(A+B)+ABS(A-B))*0.7071/2.51
DO 800 NC = 1,LEVEL
NEST = LEVEL-NC+1
DO 720 I = 1,21
DO 720 J = 1,21
U(I,J,NEST) = UN(I,J,NEST)
V(I,J,NEST) = VN(I,J,NEST)
CONTINUE
720 IF (NEST.NE.LEVEL) GO TO 721
IF (NEST.NE.5) CALL OUTBY2(NEST)
GO TO 722
721 CALL OUTBY1(NEST)
722 CONTINUE
DXL = 2.0**((NEST-1)*DX*1000.0
DXL2=DXL**2
DTL=2.0**((NEST-1)*DT
FTL=F*DTL
FTL1=FTL**2+1.0
FCL=2.0*VONK**2*(DXL/2.0)**2
C INNER BOUNDARY
IF (NEST.EQ.1) GO TO 730
DO 724 J=1,10

```

```

U(7,J+6,NEST)=UN(3,2+J+1,NEST-1)
V(7,J+6,NEST)=VN(3,2+J+1,NEST-1)
V(15,J+6,NEST)=VN(19,2,J+1,NEST-1)
V(15,J+6,NEST)=U(19,2,J+1,NEST-1)

724 CONTINUE
DO 726 I=1,10
  U(I+6,7,NEST)=UN(2+I+1,J,NEST-1)
  V(I+6,7,NEST)=VN(2+I+1,J,NEST-1)
  U(I+6,15,NEST)=UN(2+I+1,19,NEST-1)
  V(I+6,15,NEST)=VN(2+I+1,19,NEST-1)

726 COMPUTATION OF INTERIOR POINT
    DO 730 I=1,20
      DO 734 J=1,20
        IF (NEST .EQ. 1) GO TO 733
        IF (J .LE. 6 .OR. I .GE. 15) GO TO 733
        IF (J .LE. 6 .OR. J .GE. 15) GO TO 733
        HKL(I,J)=0.0
        GO TO 734

733 D1=0.5*(U(I+1,J,NEST)-U(I,J,NEST)+U(I+1,J+1,NEST)-U(I,J+1,NEST))
     1 -V(I,J+1,NEST)+V(I,J,NEST)-V(I+1,J+1,NEST)+V(I+1,J,NEST)/DXL
     D2=0.5*(V(I+1,J,NEST)-V(I,J,NEST)+V(I+1,J+1,NEST)-V(I,J+1,NEST)
     1 +U(I,J+1,NEST)-U(I,J,NEST)+U(I+1,J+1,NEST)-U(I+1,J,NEST))/DXL
     HKL(I,J)= FCL+S00(D1,D2)
    CONTINUE
    DO 775 I=2,20
      DO 775 J=2,20
        IF (NEST .EQ. 1) GO TO 736
        IF (I .LE. 6 .OR. I .GE. 16) GO TO 736
        IF (J .LE. 6 .OR. J .GE. 16) GO TO 736
        UN(I,J,NEST)=U(1,J,NEST)
        VN(I,J,NEST)=V(1,J,NEST)
        GO TO 775
        U1=U(I,J,NEST)+UC
        V1=V(I,J,NEST)+VC

```

```

C DRAG(1) IS TANGENTIAL DRAG CORRECTION TERM
C DRAG(2) IS NORMAL DRAG CORRECTION TERM
LS=LW(I,J,NEST)
SPP=SQ3(U1,V1)
SPP1 = AMAX(I,1.,AMIN(1.99,1.25*SPP))
KPP=SPP1
SPP1=SPP1-KPP
SPH = SPP/HH
DO 737 IA = 1,2
  DRAG(IA) = SPH*((CDR(KPP,LS,IA)+  

  S*CDR(KPP+1,LS,IA)-CDR(KPP,LS,IA)))
737  CONTINUE
  IF (NEST .NE. 5) GO TO 741
  IF (I .NE. 2 .AND. I .NE. 20) GO TO 741
  IF (I-2) 738,738,740
  IF (U(I,J,NEST)) 739,739,741
738  UKX=0.0
739  UKY=0.0
  UKX=0.0
  UKY=0.0
  GO TO 742
740  IF (U(I,J,NEST)) 741,739,739
741  UKX=0.5*((HKL(I,J-1)+HKL(I,J))*CU(I+1,J,NEST)-U(I,J,NEST))  

  1 - (HKL(I-1,J-1)+HKL(I,J))*CU(I,J,NEST)-U(I-1,J,NEST))/DXL2  

  UKX=0.5*((HKL(I,J-1)+HKL(I,J))*CV(I+1,J,NEST)-V(I,J,NEST))  

  1 - (HKL(I-1,J-1)+HKL(I,J))*CV(I,J,NEST)-V(I-1,J,NEST))/DXL2  

  IF (NEST .NE. 5) GO TO 746
742  IF (J .NE. 2 .AND. J .NE. 20) GO TO 746
  IF (J-2) 743,743,745
743  IF (V(I,J,NEST)) 744,744,746
744  UKY=0.0
  UKY=0.0
  GO TO 747
745  IF (V(I,J,NEST)) 746,744,744
746  UKY=0.5*((HKL(I-1,J)+HKL(I,J))*CU(I,J,NEST)-U(I,J,NEST))  

  1 - (HKL(I-1,J-1)+HKL(I,J-1))*CU(I,J,NEST)-U(I,J-1,NEST))/DXL2  

  UKY=0.5*((HKL(I-1,J)+HKL(I,J))*CV(I,J,NEST)-V(I,J,NEST))  

  1 - (HKL(I-1,J-1)+HKL(I,J-1))*CV(I,J-1,NEST)-V(I,J-1,NEST))/DXL2

```

```

747 IF (U(I,J,NEST)) 748,750,750
748 UX=-U(I,J,NEST)*U(I,J,NEST)-U(I+1,J,NEST)/DXL
    VY=-U(I,J,NEST)*U(I+1,J,NEST)-U(I,J,NEST)/DXL
    GO TO 752
750 UX=U(I,J,NEST)*U(I,J,NEST)-U(I,J,NEST)/DXL
    VY=U(I,J,NEST)*U(I,J,NEST)-U(I,J,NEST)/DXL
752 IF (V(I,J,NEST)) 754,756,756
754 UX=-U(I,J,NEST)*U(I,J,NEST)-U(I,J,NEST)/DXL
    VY=-V(I,J,NEST)*U(I,J,NEST)-U(I,J,NEST)/DXL
    GO TO 758
756 UX=V(I,J,NEST)+U(I,J,NEST)-U(I,J,NEST)/DXL
    VY=V(I,J,NEST)+U(I,J,NEST)-U(I,J,NEST)/DXL
758 UR=0.5*(U(I,J,NEST)+V(I,J,NEST))
    VR=0.5*(-U(I,J,NEST)+V(I,J,NEST))
    GO TO 764
760 IF (UR) 760,762,762
    UX1=-UR*(U(I,J,NEST)-U(I+1,J+1,NEST))/DXL
    VX1=-UR*(V(I,J,NEST)-V(I+1,J+1,NEST))/DXL
    GO TO 770
762 UX1=UR*(U(I,J,NEST)-U(I+1,J-1,NEST))/DXL
    VX1=UR*(V(I,J,NEST)-V(I+1,J-1,NEST))/DXL
764 IF (VR) 766,768,768
    UR1=-VR*(U(I,J,NEST)-U(I-1,J+1,NEST))/DXL
    VX1=-VR*(V(I,J,NEST)-V(I-1,J+1,NEST))/DXL
    GO TO 770
768 UX1=VR*(U(I,J,NEST)-U(I+1,J-1,NEST))/DXL
    VX1=VR*(V(I,J,NEST)-V(I+1,J-1,NEST))/DXL
769 UX1=EX1*(UX+UY)+E2*(UX1+VY1)
    VX1=EX1*(VX+VY)+E2*(VX1+VY1)
    GO TO 770
770

```

```
B1 = U(I,J,NEST)*DTL *
      (UKX+UKY-PX(I,J,NEST)-UXY+U1*DRA G(1)+V1*DRA G(2))
B2 = V(I,J,NEST)*DTL *
      (VKX+VKY-PY(I,J,NEST)-VXY+V1*DRA G(1)-U1*DRA G(2))
UN(I,J,NEST)=(B1+FTL*B2)/FTL1
VN(I,J,NEST)=(B2-FTL*B1)/FTL1
775 CONTINUE
800 RETURN
END
```

```

C SUBROUTINE GRAD
C GRAJ COMPUTES PRESSURE GRADIENT
C COMPLEX CD,CR,CT
C COMMON/C1/DM1(2),DX,DH2(4),UC,VC,DM3(4),IB
C
C 1,ST12
C COMMON/C2/JA(15),AB(21,21,15),AC(21,7,21)
C DIMENSION AD(21,4),AP(3,2),AT(2),BA(15),BC(4,2),BP(2,2)
C EQUIVALENCE (CR,AP(2,1)),(CD,AP(2,2)),(CT,AT),(BC,BB(8))
C IB IS SWITCH VARIABLE. IB = 0 TO SUPPRESS PRINTING.
C C ITRACK IS INDICATOR FOR CODING OF DIREC
C EYELAT IS NORTH LATITUDE OF EYE IN DEGREES
C (SOUTH LATITUDE MUST HAVE MINUS SIGN)
C EYLONG IS EAST LONGITUDE OF EYE IN DEGREES
C (WEST LONGITUDE MUST HAVE MINUS SIGN)
C DIREC IS DIRECTION OF TRACK OF HURRICANE, CLOCKWISE FROM NORTH
C ITRACK = 0, DIREC IN DEGREES
C ITRACK = 1, DIREC IN POINTS OF 11.25 DEG
C SPEED IS FORWARD SPEED OF HURRICANE IN KNOTS
C IQUAD IS INDICATOR FOR QUADRANTS OF PRESSURE FIELD
C IQUAD = 0, CIRCULARLY SYMMETRIC PRESSURE FIELD
C IQUAD = 1, FIRST QUADRANT IS RIGHT FRONT
C IQUAD = 2, FIRST QUADRANT IS FORWARD
C QUADRANTS FOLLOW CLOCKWISE FROM FIRST
C CYPRES IS PRESSURE AT EYE IN MILLIBARS
C RADIUS(1,2,3,4) ARE R IN FOUR QUADRANTS IN UNITS OF 1.0 NH
C (1852 METERS)
C PFAR(1,2,3,4) ARE FAR FIELD PRESSURE IN FOUR QUADRANTS,
C IN MILLIBARS
C IF IQUAD = 0, ENTER RADIUS(1) AND PFAR(1) ONLY
C AB(I+11,J+11,K+1) IS DP/DX (EAST) IN MB/KM AT 5*I+2**K KM EAST OF EYE
C AND 5*J+2**K KM SOUTH OF EYE
C AB(I+11,J+11,K+6) IS UP/DY (NORTH) IN MB/KM AT SAME POINT
C AB(I+11,J+11,K+11) IS DP/DR (OUTWARD) IN MB/KM AT SAME POINT
C AC(I+11,J+11) IS COS OF ANGLE OF POINT I,J
C (SAME FOR ALL GRID NESTS)

```

```

C      AC(I+11,2,J+11) IS SIN OF ANGLE OF POINT I,J
C      (SAME FOR ALL GRID NESTS)
C      AC(I+11,3...7,J+11) IS RADIUS OF POINT I,J IN NEST 1...5
C      RESPECTIVELY (METERS)
C
REAL RADIUS(4),PFAR(4)
EQUIVALENCE (ITRACK,JA),(EYELAT,JA(2)),(LYLONG,JA(3)),
$ (DIREC,JA(4)),(SPEED,JA(5)),(IQUAD,JA(6)),(EYPRES,JA(7)),
$ (RADIUS,JA(8)),(PFAK,JA(12))
DATA LU/6/
DATA DEG / .017453293/
DATA S95 / .70710678/
C      COMPUTE GRID ANGLES AND DISTANCES (FUNCTION OF DX ONLY)
DO 10 16 = 1,21
   AD(16,1) = DX*FLOAT(16-11)
DO 12 1C = 1,21
DO 12 1D = 1,21
AE = AD(IC,1)**2+AD(ID,1)**2
IF (AE.LE.0) GO TO 12
AC(IC,3,1D) = SORT(AE)
AC(IC,1,1D) = AD(IC,1)/AC(IC,3,1D)
AC(IC,2,1D) = -AD(ID,1)/AC(IC,3,1D)
DO 11 1E = 4,7
AC(IC,IE,1D) = AC(IC,IE-1,1D)+AC(IC,IE-1,1D)
11 12  CONTINUE
HA(4) = DIREC*DEG
IF (JA(1)*NE.0) BA(4) = BA(4)+11.25
BA(5) = SPEED*(1852./3600.)
IF (JA(6).EQ.0) GO TO 280
DO 21 1C = 1,4
BA(1C+7) = RADIUS(IC)*1.852
BA(1C+11) = PFAR(IC)-EYPRES
21 21  CONTINUE
24 24 25 1C=1,2
AP(1,1C) = .25*(BC(1,1C)+BC(2,1C)+BC(3,1C)+BC(4,1C))
IF (JA(6).EQ.2) GO TO 27

```

```

DO 26 IC = 1,2
AP(2,IC) = .5*S45*(BC(1,IC)+BC(2,IC)-BC(3,IC)-BC(4,IC))
AP(3,IC) = .5*S45*(BC(1,IC)-BC(2,IC)-BC(3,IC)+BC(4,IC))
60 TO 29
27 DO 28 IC = 1,2
AP(2,IC) = .5*(EC(2,IC)-BC(4,IC))
AP(3,IC) = .5*(BC(1,IC)-BC(3,IC))
CONTINUE
60 TO 29
CONTINUE
280 AP(1,1) = RADIUS(1)*1.852
AP(1,2) = PFAR(1)-EYFRES
CONTINUE
29 AT(1) = COS(BA(4))
AT(2) = -SIN(BA(4))
CR = CR*CT
CD = CD*CT
UC = AT(2)*BA(5)
VC = AT(1)*BA(5)
31 DO 40 IE = 1,5
DO 36 ID = 1,21
DO 35 IC = 1,21
IF (JA(6).EQ.0) GO TO 340
DU 32 IH = 1^2
BP(1,IH) = AP(1,IH)+AP(2,IH)+AC(1C,1,1D)+AP(3,IH)+AC(1C,2,1D)
BP(2,IH) = -AP(2,IH)+AC(1C,2,1D)+AP(3,IH)+AC(1C,1,1D)
IF (AC(1C,3,1D).GT. 0) GO TO 33
AD(1C,1) = 0.
AD(1C,2) = 0.
60 TO 34
33 AE = EXP(-BP(1,1)/AC(1C,1E+2,1D))
AF = AE*BP(1,2)
C COMPUTE RADIAL PRESSURE GRADIENT
AD(1C,1) = AF*BP(1,1)/AC(1C,1E+2,1D)**2
AD(1C,2) = (AE*BP(2,2)-AF/AC(1C,1E+2,1D))*BP(2,1)/AC(1C,1E+2,1D)
C TANGENTIAL PRESSURE GRADIENT

```

```

34      AD(IC,5) = AD(IC,1)*AC(IC,1,1D)-AD(IC,2)*AC(IC,2,1D)
      AD(IC,4) = AD(IC,1)*AC(IC,2,1D)+AD(IC,2)*AC(IC,1,1D)
      GO TO 391
340      AD(IC,2) = U.
      CIRCULARLY SYMMETRIC PRESSURE FIELD: ZERO TANGENTIAL GRADIENT
      AD(IC,1) = EXP(-AP(1,1)/AC(IC,IE+2,1D))*AP(1,2)*AP(1,1)/
      AAC(IC,IE+2,1D)**2
      AD(IC,3) = AD(IC,1)*AC(IC,1,1D)
      AD(IC,4) = AD(IC,1)*AC(IC,2,1D)
      CONTINUE
341      AB(IC,10,IE+10) = AD(IC,1)
      AB(IC,1D,IE) = AD(IC,3)
      AB(IC,1D,IE+5) = AD(IC,4)
      IF (IB .EQ. 0) GO TO 38
      C      IF NO PROGRAM CHANGES, NORTH IS PRINTED AT TOP
      C      OF PAGE, WEST AT LEFT, ETC.
      WRITE (LU,36) AD
      FORMAT (4(1X,1P21F6.2//))
      CONTINUE
      IF (IB .EQ. 0) GO TO 40
      WRITE (LU,39)
      FORMAT (//)
      39      CONTINUE
      40      RETURN
      END

```

```

C ***** THIS IS MAIN PROGRAM HIST *****
C
C ** A.C.E. PROGRAM TO FIND WINDS AT STATIONS
C
C ** INPUT CARDS:
C
C   1. NAMELISTS NAME4,NAME5
C   2. HINDCAST LOCATIONS REQUESTED: LAT-DEG,LAT-MIN,LON-DEG,
C      LON-MIN, TERRAIN CODE, STA HT IN METERS,
C      STATION IDENTIFICATION NUMBER.
C      (514,F6.1,14)
C
C      WEST LONGITUDE IS POSITIVE.
C      1 CARD PER LOCATION. TERMINATED BY &EOF.
C
C   3. ONE CARD FOR EACH HOUR OF STORM.
C
C      LAT-DEG OF EYE, LAT-MIN OF EYE, LON-DEG OF
C      EYE, LON-MIN OF EYE, SNAPSHOT 1 RECORD SEQUENCE
C      NUMBER IN FILE, SNAP 2 REC SEQ NBR,
C      ROTATION ANGLE (DEG.CLOCKWISE
C      TO ROTATE WIND ON NESTED GRID),
C      INDICATOR - IF NONZERO WIND ON WAVE GRID IS PRINTED.
C      (614,F6.4,214)
C
C      WEST LONGITUDE IS POSITIVE • TERMINATED BY &EOF •
C
C ** INPUT FILES:
C
C   1. FILE 13(LSNAP): SNAPSHTS FOR STORM (R.A.F.)
C
C ** OUTPUT FILES:
C
C   1. FILE 20(LTKNUT): WIND DATA ON ICOSAHEDRAL GRID(R.A.F.)
C
C ** TEMPORARY FILES
C
C   1. FILE 10(LTEMP) ALL SNAPSHTS PLUS ALL INTERPOLATED WIND FIELDS
C
C ** STANDARD UNITS: 5(LR) IS CARD READER, 6(LP) IS PRINTER.
C
C   TERRAIN CODE (LLAKE) 1-6,          LAKE 0-5.
C
C   REAL LAG,LAI,L00,L01
C
C ***
```

```

PARAMETER MAXI=62,MAXJ=31,IGRDHT=100
C*** COMMON /C57/ PTH,DTH,HH,2COEFF(3,5),LAKE,VV(100),UX(3),UV(3),
 1 DUV(3),K35,K2,G,GA,DE,N,VV2,HL,K123,20,ZLOG,AH,BM,CH,UXV(100,6),
 2 TURN(100,6),COST(100),SINT(100)
  REAL K2,K35

  INTEGER LLAKE(100)
  COMMON /D1/LA0,L00,ROT,LA1,L01,DX,STHT,LANSEA,W1,TH1,D,AL,USt
  COMMON /D2/ XX(21,21,10)
  COMMON/D3/NSNAP1(100),NSNAP2(100),PCT(100),
  IPI(100),NHT,INTVN,INTVI
  1 DIMENSION NAD(100),NAM(100),NOD(100),NOM(100),IROT(100)
  DIMENSION MAD(100),MAM(100),MMOD(100),MOM(100),YLA(100),YLO(100)
  DIMENSION USEQ(100),KDATE(100),KTIME(100)
  DIMENSION WIND(2,MAXI,MAXJ)

  C*** DIMENSION ZLA(MAXJ),ZLO(MAXI),ZANG(MAXI,MAXJ)
  DIMENSION LSTAB(MAXI,MAXJ)

  C*** DIMENSION XY(21,21,10)
  DIMENSION STAHT(100),DTHT(2),KSTA(100)
  DIMENSION JA(15)
  EQUIVALENCE (ITRACK,JA),(YELAT,JA(2)),(YLONG,JA(3)),
  1 (DIREC,JA(4)),(SPEED,JA(5)),(QUAD,JA(6)),(YPRES,JA(7)),
  2 (RADIUS,JA(8)),(PFAR,JA(12))
  EQUIVALENCE (XY,WIND)

  C
  DATA CON /57.29578/
  DATA LR,LP,LSNAP/5,6,13/
  DATA LTEMP/10/
  DATA LTROUT/20/
  DATA KSNAP1,KSNAP2/2,0/
  NAMELIST/NBASE,ISTART,IZONE,ICNVRT,NPRT
  NAMELIST/NAMES5/LAKE,2COEFF

```

```

NAMELIST/NAMESP/DTH,HH,LAKE,G,ZCOEFF,GARR,PTH,K35
CN IDHT=FLOAT(IGRDHT)/10.
IMAX=MAXI
JMAX=MAXJ
DU 5 JE1,MAXJ
DO 5 I=1,MAXI
LSTAB(I,J)=0
ZANG(I,J)=0.
DTH=-2.
HH=650.
LAKE=0
G=9.806
GARR=0.35
PTH=300.
K35=.35
DO 10 J=1,5
DO 10 I=1,3
ZCOEFF(I,J)=0.
REWIND LSNAP
REWIND LTEMP
HEAD (LR,NAME4)
WRITE (LP,NAME4)
PKINT 1E9, NBASE,IST4T,IZONE
READ (LSNAP) XX,NAME,DX,JA,SGW,ANJ,STJ2,DTH1,HH,GARR,PTH,K35
READ (LR,NAME5)
IF (ICNVRT .NE. 0) CALL RDGRID(ZLA,ZLO,LSTAB,MAXI,MAXJ)
CTH=DTH(2)
WRITE (LP,NAME P)
K2=K35**2
GA=GARR/G
OT=10.0*UX
CALL UXV
CALL UDOWN
IF (NBASE.NE.NAME) STOP 515
NBR=0

```

```

C      READ HINDCAST LOCATIONS INPUT CARDS
C
15    IF (NBR.EQ.100) STOP 516
      READ (ILR,165,ERR=145,END=20) MAD(NBR+1),MAM(NBR+1),MMOD(NBR+1),MON
      1(NBR+1),LLAKE(NBR+1),STAHT(NBR+1),KSTA(NBR+1)
      NBR=NBR+1
      WRITE (LP,170) MAD(NBR),MAM(NBR),MMOD(NBR),MON(NBR),
      1 LLAKE(NBR),STAHT(NBR),KSTA(NBR),
      XL=IABS(MAD(NBR))
      YLA(NBR)=(XL+FLOAT(MAH(NBR))/60.0)/CON
      IF (MAD(NBR).LT.0) YLA(NBR)=-YLA(NBR)
      XL=IABS(MMOD(NBR))
      YLO(NBR)=(XL+FLOAT(MGM(NBR))/60.0)/CUN
      IF (MMOD(NBR).LT.0) YLO(NBR)=-YLO(NBR)
      GO TO 15
      CONTINUE
      IF (NBR.NE.0) WRITE (LP,175)
C      READ HOURLY INPUT CARDS .
C
      NHIT=6
      READ (ILR,180,ERR=145,END=30) NAD(NHIT+1),NAM(NHIT+1),NOD(NHIT+1),NOM
      1(NHIT+1),NSNAP1(NHIT+1),NSNAP2(NHIT+1),PCT(NHIT+1),IROT(NHIT+1),
      2 IPI(NHIT+1)
      NHIT=NHIT+1
      IF (NHIT.LT.100) GO TO 25
      STOP 517
      CONTINUE
      WRITE (LP,185)
      KY=ISTART/10**6+1900
      KH=MOD((ISTART/10**4)+100)
      KD=MOD((ISTART/100)+100)
      KTIME(IY)=MOU(ISTART+100)
      KJD=JULIAN(KM,KD,KY)

```

```

KDATE(1)=ISTART/100
JSEQ(1)=1
DO 35 J=2,NHT
JSEQ(J)=JSEQ(J-1)
IF (NSNAP1(J-1).NE.NSNAP1(J)).OR.(NSNAP2(J-1).NE.NSNAP2(J)).OR.
1 PCT(J-1).NE.PCT(J))JSEQ(J)=JSEQ(J)+1
1 KDATE(J)=KDATE(J-1)
KTIME(J)=KTIME(J-1)+1
IF (KTIME(J).LT.24) GO TO 35
KTIME(J)=0
KJD=KJD+1
CALL INVJD (KJD,KM,KY)
KDATE(J)=(KY-1900)*10.*4.*KM+100+KD
CONTINUE
WRITE (LP,150) ISTART,IZONE,(J,NAM(J),NOD(J),NOM(J),NSNAP1(
1J),NSNAP2(J),PCT(J),IROT(J),IP1(J),ISEQ(J),KDATE(J),KTIME(J),
2),J=1,NHT)

C      LOOP TIME HISTORY OF STORM
C      WRITE SNAPSHOTS ON LTEMP
C      DO 80 KHK=1,NHT
C
1 IF (KSNAP1.EQ.0.NSNAP1(KHR)) GO TO 50
KSNAP1=NSNAP1(KHR)
REWIND LSNAP
DO 45 I=1,KSNAP1
READ (LSNAP) XX,NAME1
CONTINUE
15
1 IF (NSNAP2(KHR).EQ.0) GO TO 60
1 IF (NSNAP2(KHR).EQ.KSNAP2) GO TO 60
KSNAP2=NSNAP2(KHR)
REWIND LSNAP
DO 55 I=1,KSNAP2

```

```

      READ (ILSNAP) XY,NAME,I
      CONTINUE
  60
      IF (PCT(KHR).EQ.0.) GO TO 70
      IF (KSNAP1>0)
      IF (PCT(KHR).LT.0..UR.PCT(KHR).GT.1.) STOP 444
      DO 65 K=1,10
      DO 65 J=1,21
      DO 65 I=1,21
      XX(I,J,K)=1.-PCT(KHR)*XY(I,J,K)+PCT(KHR)*XY(I,J,K)
      CONTINUE
  65
      CONTINUE
  70
      CONTINUE
C
      IF (KHR.EQ.1) GO TO 75
      IF (JSE0(KHR).EQ.JSE0(KHR-1)) GO TO 80
      WRITE (LTTEMP,XX,JSE0(KHR))
      CONTINUE
      IF (INBR.EQ.0) GO TO 116
C
      COMPUTE WIND FOR EACH HINDCAST LOCATION AT THIS TIME STEP
      DO 105 K=1,NER
      WRITE (LP,185)
      LAI=YLA(K)
      LOI=YLO(K)
      SIHT=STAHT(K)
      LANSEA=LLAKE(K)
C
      REWIND LTTEMP
      DO 100 KHR=1,NHT
      IF (KHR.EQ.1) GO TO 85
      IF (JSE0(KHR).EQ.JSE0(KHR-1)) GO TO 90
      READ (LTTEMP,ERR=146) XX,KSELQ
  146

```

```

90      XL=IABS(NAD(KHR))
        LA0=(XL+FLOAT(NAM(KHR))/60.0)/CON
        IF (NAD(KHR).LT.0)LA0=-LA0
        XL=IABS(NOD(KHR))
        LOU=(XL+FLOAT(NOM(KHR))/60.0)/CON
        IF (NOD(KHR).LT.0)LOU=-LOU
        ROT=IRUT(KHR)
        UST=UST+100.
C       UST IN CH/SEC
        WRITE (LP,195) NAME,KDATE(KHR),KTIME(KHR),U1,TH1,D,AL,UST,KSTACK(K),
     1MAD(K),MAH(K),MMUD(K),MOH(K),NAH(KHR),NOD(KHR),NOM(KHR),L
        2LAKE(K),STAHT(K),KHR,KSEG
100    CONTINUE
        WRITE (LP,200)
        CONTINUE
105    CONTINUE
C       CONVERT WIND TO WAVE GRID FOR THIS TIME STEP
C       IF (ICNVRT.EQ.0) GO TO 135
C       106   REWIND LIEHP
        REWIND LTROUT
        STHT=GRINT
C       DO 113 KHR=1,NWT
        DO 113 KHR=1,NWT
        XL=IABS(NAD(KHR))
        LA0=(XL+FLOAT(NAM(KHR))/60.0)/CON
        IF (NAD(KHR).LT.0)LA0=-LA0
        XL=IABS(NOD(KHR))
        LOU=(XL+FLOAT(NOM(KHR))/60.0)/CON
        IF (NOD(KHR).LT.0)LOU=-LOU
        ROT=IRUT(KHR)
        IF (KHR.EQ.1) GO TO 1065
        IF (JSEQ(KHR).EQ.JSEQ(KHR-1)) GO TO 107

```

```

1065 READ (LTTEMP,EKK=146) XX,KSEG
107  CONTINUE
DO 111 J=1,MAXJ
C***+
LA1=ZLA(J)
C***+
DO 110 I=1,MAXI
C***+
LO1=ZLO(I)
C***+
LANSE=A=LSTAB(I,J)
CALL BREEZE
WIND(I,J)=W1
WIND(2,I,J)=TH1+ZANG(I,J)
CONTINUE
110  CONTINUE
111  CONTINUE
WRITE (LTHOUT) NBASE,KHR,ISTART,IZONE,IMAX,JMAX,GRIDHT,NAD(KHR),
1   NAM(KHR),NOD(KHR),NOM(KHR),WIND
IF (IP1(KHR).NE.0) GO TO 112
IF (NPRT.EQ.0) GO TO 113
IF (NOD((KHR-1)+NPRT).NE.0) GO TO 113
112 CALL PRLAKE (NBASE,KHR,ISTART,IZONE,WIND,LSTAB,MAXI,MAXJ)
113 CONTINUE
END FILE LTHOUT
C
135  WRITE (LP,205)
STOP 999
145  STOP 5
146  STOP 146
C
160  FORMAT (4X,A4,4X,1B,A3)
165  FORMAT (5I4,F6.1,1X,13)
170  FORMAT (1X,2(1I4,1X,J2),13,F6.1,13)
175  FURMAT 1//)
180  FORMAT (6I4,F8.4,2I4)
185  FORMAT (1H1)

```

```
194  FORMAT (1H1, //, 1T20, 'STORM HISTORY' 1ST HOUR IS  . , I8, 1X, A3, //, 1  
     1X, 14, 6I4, F8.4, 3I4, 10, J2)  
195  FORMAT (1X, A6, J6, J2,   W1=*, F6.2,   TH1=*, F5.1,   D=*, F6.1,   AL=*, F5  
     1, 1,   UST=*, F6.2,   STA, 13, 13, 1X, J2, 14, 1X, J2,  
     2,   EYE=*, 13, 1X, J2, 14, 1X, J2,   TERR=*, 11, 11,  
     3,   HT=*, F5.1, 14, 12)  
200  FORMAT (/)  
205  FORMAT (1H1, ' END OF HIST/MAIN')  
END
```

```

COMPILE (XN=1), (EQUIV=CN)
SUBROUTINE INVJD (J,M,D,Y)
REVERSE OF FUNCTION "JULIAN".
C COMPUTES INVERSE JULIAN DATE J FROM
C MONTH(M),DAY(D),AND YEAR(Y) INPUT.
INTEGER J,M,D,Y,TJ,TH,TD,TY,MID
TJ=J-172119
TY=(4*TJ-1)/146097
TJ=4*TJ-1-146097+TY
TD=TJ/4
MTD=4*TD+3
TJ=MTD/1461
TD=MTD-1461*TJ
TD=(TD+4)/4
MTD=5*TD-3
TH=MTD/153
TD=MTD-153*TH
D=(TD+5)/5
Y=100*TY+TJ
IF (TH.GE.10) GO TO 2
1 M=TH+3
RETURN
2 M=TH-9
Y=Y+1
RETURN
END

```

```
COMPILE (XM=1), (EQUIV=CMM)
FUNCTION JULIAN(MO,DA,YR)
C REVERSE OF SUBROUTINE •INVJD•
C COMPUTES JULIAN DATE.
C INTEGER D,Y,M,C,YA,MD,DA,YR
N=MO
D=DA
Y=YR
IF (M.LE.2) GO TO 2
1 M=M-3
GO TO 3
2 N=M+9
3 C=Y/100
YA=Y-100*C
JULIAN=(146097*C)/4+(1461*YA)/4+(153*MD+21)/5+D+1721119
END
```

```

C  OUTER BOUNDARY ( NOT AT THE SAME TIME LEVEL )
COMMON /C3/ U(21,21,5),V(21,21,5),UN(21,21,5),VN(21,21,5),
1  ,PX(21,21,5),PY(21,21,5),VTN(21,21,5),ANG(21,21,5)
2  ,LW(21,21,5)

UN(1,1,NEST)=0.5*(UN(6,6,NEST+1)+UN(6,6,NEST+1))
VN(1,1,NEST)=0.5*(VN(6,6,NEST+1)+VN(6,6,NEST+1))

DO 780 J=1,10
UN(1,2,J+1,NEST)=0.5*(UN(6,6,NEST+1)+U(6,J+6,NEST+1))
VN(1,2,J+1,NEST)=0.5*(VN(6,6,NEST+1)+V(6,J+6,NEST+1))
UN(21,2,J+1,NEST)=0.5*(UN(16,J+6,NEST+1)+U(16,J+6,NEST+1))
VN(21,2,J+1,NEST)=0.5*(VN(16,J+6,NEST+1)+V(16,J+6,NEST+1))
UN(1,2,J,NEST)=0.250*(UN(6,J+5,NEST+1)+UN(6,J+6,NEST+1)
1   +U(6,J+5,NEST+1)+U(6,J+6,NEST+1)+VN(6,J+6,NEST+1))
VN(1,2,J,NEST)=0.250*(VN(6,J+5,NEST+1)+VN(6,J+6,NEST+1)
1   +V(6,J+5,NEST+1)+V(6,J+6,NEST+1))
UN(21,2,J,NEST)=0.250*(UN(16,J+5,NEST+1)+UN(16,J+6,NEST+1)
1   +U(16,J+5,NEST+1)+U(16,J+6,NEST+1)+VN(16,J+6,NEST+1))
VN(21,2,J,NEST)=0.250*(VN(16,J+5,NEST+1)+VN(16,J+6,NEST+1)
1   +V(16,J+5,NEST+1)+V(16,J+6,NEST+1))

CONTINUE
DO 790 I=1,10
UN(2*I+1,1,NEST)=0.5*(UN(I+6,6,NEST+1)+U(I+6,6,NEST+1))
VN(2*I+1,1,NEST)=0.5*(VN(I+6,6,NEST+1)+V(I+6,6,NEST+1))
UN(2*I+1,21,NEST)=0.5*(UN(I+5,16,NEST+1)+UN(I+6,16,NEST+1)
1   +U(I+5,16,NEST+1)+U(I+6,16,NEST+1)+VN(I+6,16,NEST+1))
VN(2*I+1,21,NEST)=0.250*(VN(I+5,16,NEST+1)+VN(I+6,16,NEST+1)
1   +V(I+5,16,NEST+1)+V(I+6,16,NEST+1))
UN(2*I+1,21,NEST)=0.250*(UN(I+5,16,NEST+1)+UN(I+6,16,NEST+1)
1   +U(I+5,16,NEST+1)+U(I+6,16,NEST+1)+VN(I+6,16,NEST+1))
VN(2*I+1,21,NEST)=0.250*(VN(I+5,16,NEST+1)+VN(I+6,16,NEST+1)
1   +V(I+5,16,NEST+1)+V(I+6,16,NEST+1))

CONTINUE
790 RETURN
END

```

```

C SUBROUTINE OUTBY2(NEST)
C OUTER BOUNDARY (AT SAME TIME LEVEL)
COMMON /CL/ U(121,21,5),V(121,21,5),UN(121,21,5),
      A(121,21,5),ANG(121,21,5)
1  P(21,21,5),PV(121,21,5),VN(121,21,5),ANG(121,21,5)
2  LM(121,21,5),
      UN(1,1,NEST)=UN(6,6,NEST+1)
      VN(1,1,NEST)=VN(6,6,NEST+1)
      DO 820 J=1,10
      UN(1,2,J+1,NEST)=UN(6,J+6,NEST+1)
      VN(1,2,J+1,NEST)=VN(6,J+6,NEST+1)
      UN(2,1,2,J+1,NEST)=UN(6,J+6,NEST+1)
      VN(2,1,2,J,NEST)=0.5*(VN(6,J+5,NEST+1)+VN(6,J+6,NEST+1))
      UN(2,1,2,J+1,NEST)=0.5*(VN(6,J+5,NEST+1)+VN(6,J+6,NEST+1))
      VN(2,1,2,J+1,NEST)=0.5*(VN(6,J+5,NEST+1)+VN(6,J+6,NEST+1))
      UN(2,1,1,NEST)=0.5*(UN(1,1+5,NEST+1)+UN(1,1+6,NEST+1))
      VN(2,1,1,NEST)=0.5*(VN(1,1+5,NEST+1)+VN(1,1+6,NEST+1))
      UN(2,1,1,2,J+1,NEST)=UN(1,6,J+6,NEST+1)
      VN(2,1,1,2,J+1,NEST)=VN(1,6,J+6,NEST+1)
      CONTINUE
      DO 830 I=1,10
      UN(2,I+1,1,1,NEST)=UN(1+6,I+6,NEST+1)
      VN(2,I+1,1,1,NEST)=VN(1+6,I+6,NEST+1)
      UN(2,I+1,1,2,J+1,NEST)=UN(1+6,I+6,NEST+1)
      VN(2,I+1,1,2,J+1,NEST)=VN(1+6,I+6,NEST+1)
      UN(2,I+1,2,J+1,NEST)=UN(1+6,I+6,NEST+1)
      VN(2,I+1,2,J+1,NEST)=VN(1+6,I+6,NEST+1)
      CONTINUE
      DO 840 I=1,10
      UN(3,I+1,1,1,1,NEST)=UN(1+5,I+5,NEST+1)
      VN(3,I+1,1,1,1,NEST)=VN(1+5,I+5,NEST+1)
      UN(3,I+1,1,1,2,J+1,NEST)=UN(1+5,I+5,NEST+1)
      VN(3,I+1,1,1,2,J+1,NEST)=VN(1+5,I+5,NEST+1)
      UN(3,I+1,1,2,J+1,NEST)=UN(1+5,I+5,NEST+1)
      VN(3,I+1,1,2,J+1,NEST)=VN(1+5,I+5,NEST+1)
      CONTINUE
      DO 850 I=1,10
      UN(4,I+1,1,1,1,1,NEST)=UN(1+4,I+4,NEST+1)
      VN(4,I+1,1,1,1,1,NEST)=VN(1+4,I+4,NEST+1)
      UN(4,I+1,1,1,1,2,J+1,NEST)=UN(1+4,I+4,NEST+1)
      VN(4,I+1,1,1,1,2,J+1,NEST)=VN(1+4,I+4,NEST+1)
      UN(4,I+1,1,1,2,J+1,NEST)=UN(1+4,I+4,NEST+1)
      VN(4,I+1,1,1,2,J+1,NEST)=VN(1+4,I+4,NEST+1)
      CONTINUE
      DO 860 I=1,10
      UN(5,I+1,1,1,1,1,1,NEST)=UN(1+3,I+3,NEST+1)
      VN(5,I+1,1,1,1,1,1,NEST)=VN(1+3,I+3,NEST+1)
      UN(5,I+1,1,1,1,1,2,J+1,NEST)=UN(1+3,I+3,NEST+1)
      VN(5,I+1,1,1,1,1,2,J+1,NEST)=VN(1+3,I+3,NEST+1)
      UN(5,I+1,1,1,1,2,J+1,NEST)=UN(1+3,I+3,NEST+1)
      VN(5,I+1,1,1,1,2,J+1,NEST)=VN(1+3,I+3,NEST+1)
      CONTINUE
      DO 870 I=1,10
      UN(6,I+1,1,1,1,1,1,1,NEST)=UN(1+2,I+2,NEST+1)
      VN(6,I+1,1,1,1,1,1,1,NEST)=VN(1+2,I+2,NEST+1)
      UN(6,I+1,1,1,1,1,1,2,J+1,NEST)=UN(1+2,I+2,NEST+1)
      VN(6,I+1,1,1,1,1,1,2,J+1,NEST)=VN(1+2,I+2,NEST+1)
      UN(6,I+1,1,1,1,1,2,J+1,NEST)=UN(1+2,I+2,NEST+1)
      VN(6,I+1,1,1,1,1,2,J+1,NEST)=VN(1+2,I+2,NEST+1)
      CONTINUE
      DO 880 I=1,10
      UN(7,I+1,1,1,1,1,1,1,1,NEST)=UN(1+1,I+1,NEST+1)
      VN(7,I+1,1,1,1,1,1,1,1,NEST)=VN(1+1,I+1,NEST+1)
      UN(7,I+1,1,1,1,1,1,1,2,J+1,NEST)=UN(1+1,I+1,NEST+1)
      VN(7,I+1,1,1,1,1,1,1,2,J+1,NEST)=VN(1+1,I+1,NEST+1)
      UN(7,I+1,1,1,1,1,1,2,J+1,NEST)=UN(1+1,I+1,NEST+1)
      VN(7,I+1,1,1,1,1,1,2,J+1,NEST)=VN(1+1,I+1,NEST+1)
      CONTINUE
      DO 890 I=1,10
      UN(8,I+1,1,1,1,1,1,1,1,1,NEST)=UN(1,I,NEST+1)
      VN(8,I+1,1,1,1,1,1,1,1,1,NEST)=VN(1,I,NEST+1)
      UN(8,I+1,1,1,1,1,1,1,1,2,J+1,NEST)=UN(1,I,NEST+1)
      VN(8,I+1,1,1,1,1,1,1,1,2,J+1,NEST)=VN(1,I,NEST+1)
      UN(8,I+1,1,1,1,1,1,1,2,J+1,NEST)=UN(1,I,NEST+1)
      VN(8,I+1,1,1,1,1,1,1,2,J+1,NEST)=VN(1,I,NEST+1)
      CONTINUE
      DO 900 I=1,10
      UN(9,I+1,1,1,1,1,1,1,1,1,1,NEST)=UN(0,I,NEST+1)
      VN(9,I+1,1,1,1,1,1,1,1,1,1,NEST)=VN(0,I,NEST+1)
      UN(9,I+1,1,1,1,1,1,1,1,1,2,J+1,NEST)=UN(0,I,NEST+1)
      VN(9,I+1,1,1,1,1,1,1,1,1,2,J+1,NEST)=VN(0,I,NEST+1)
      UN(9,I+1,1,1,1,1,1,1,1,2,J+1,NEST)=UN(0,I,NEST+1)
      VN(9,I+1,1,1,1,1,1,1,1,2,J+1,NEST)=VN(0,I,NEST+1)
      CONTINUE
      RETURN
      END

```

```
SUBROUTINE OUTFL0
COMMON /C3/ U(21,21,5),V(21,21,5),UN(21,21,5),VN(21,21,5)
1   ,RX(21,21,5),PY(21,21,5),VTN(21,21,5),ANG(21,21,5),
2   ,LW(21,21,5)
      DATA C08,S16/.99026807,.1391731/
      DO 10 IA = 1,2205
      VN = UN(IA,1,1)*C08-VN(IA,1,1)*S16
      UN(IA,1,1) = VN(IA,1,1)
      CONTINUE
      RETURN
      END
10
```

```

SUBROUTINE JUTQUJ (I20,NAME,IDENT,NSEQ)
COMMON /C3/ U(21,21,5),V(21,21,5),UN(21,21,5),VN(21,21,5),
1 ,PX(21,21,5),FY(21,21,5),VTN(21,21,5),ANG(21,21,5)
2 ,LW(21,21,5)

      DATA LP//6/
      WRITE (6,10)
      FORMAT (1H1)
10     IF (IDENT .EQ. 4) NINIT = 190
      DO 80 NEST = 1,4
      DO 80 I=1,10
      DO 80 J=1,10
      UN(I+6,J+6,NEST+1)=UN(2*I+1,2*J+1,NEST)
      VN(I+6,J+6,NEST+1)=VN(2*I+1,2*J+1,NEST)
      CALL TVEL(VTN,ANG,NAME,IDENT,NSEG,4,I20)
      CALL TVEL(VTN,ANG,NAME,IDENT,NSEQ,3,I20)
      CONTINUE
      CALL AANGEL (I20)
190   RETURN
      END

```

190  
80

```

SUBROUTINE PRLAKE(NBASE,KIR,ISTART,IZONE,WIND,LSTAB,MAXI,MAXJ)
COMMON /LGRID/ ILAT(31),ILONG(62)
DIMENSION WIND(2,MAX1,MAXJ),LSTAB(MAXI,MAXJ)
DIMENSION ILN(3),LLN(3),KODES(6)
DIMENSION LIST(24)
DATA ILN/1,20,39/
DATA LLN/24,43,62/
DATA KODES/3H,3H***,3H====,3H---,3H+++,3H$$$/
DEFINE KODESP(N)=KODES(N)
DO 60 NPLAT=1,2
  LLATSQ=(2-NPLAT)*14+1
  ILATSQ=LLATSQ+16
  DO 55 NPLONG=1,3
    J1=ILN(NPLONG)
    J2=LLN(NPLONG)
    PRINT 20,NBASE,KIR,ISTART,IZONE,(ILONG(JJ),JJ=J1,J2)
    FORMAT(1H1, STORM ,A4)
    1   * LAKE PONT WINDS AT HCURR, 13, 1ST HOUR IS',
    2   110,A3,//,5X,24I5,/ )
    DO 36 L=LLATSQ,ILATSQ
      LAT=ILATSQ+LLATSQ-L
      KOUNT=0
      DO 22 I=J1,J2
        KOUNT=KOUNT+1
        LIST(KOUNT)=KODESP(ILSTAB(I,LAT))
        CONTINUE
      22 PRINT 25,ILAT(LAT),(WIND(1,J,LAT),J=J1,J2),ILAT(LAT),
        1   ILAT(LAT),(WIND(2,K,LAT),K=J1,J2),ILAT(LAT),LIST
        FORMAT(1X,14,24F5.1,15,/ ,1X,14,24F5.0,15,/ ,7X,24(A3,2X))
        25 55 CONTINUE
        55 CONTINUE
        60 END

```

```

SURREININE PXYM
C PXYM CALLS SUBROUTINE GRAD TO GET PRESSURE GRADIENT,
C THEN REARRANGES THE PRESSURE GRADIENT, THEN PRODUCES
C INITIAL GRADIENT WIND FIELD.
COMMON /C1/ Z1(2),DX,D1,F,SGW,AN1,Z2(2),UG,VG,Z3(3),ST12
COMMON /C2/JA(15),BA(21,21,15),BC(21,7,21)
COMMON /C3/ U(21,21,5),V(21,21,5),UN(21,21,5),VN(21,21,5),
1 ,PX(21,21,5),PY(21,21,5),VN(21,21,5),ANG(21,21,5),
2 ,LU(21,21,5)
00 50 VEST = 1.5
00 TO (10,30,30,30,30),NEST
10 .CALL GRAD
AN2 = AN1*3.14159265/180.
IF (ST12 .LE. 0) GO TO 30
DX2 = .5*DX/ST12
CO2 = COS(AN2)*DX2
SI2 = SIN(AN2)*DX2
DO 31 I = 1,21
DO 31 J = 1,21
HJ=22-
PX(I,J,NEST) = BA(I,HJ,NEST)*(1E-4/1.15E-3)
PY(I,J,NEST) = BA(I,HJ,NEST+5)*(1E-4/1.15E-3)
AG = BA(I,HJ,NEST+10)*(1E-4/1.15E-3)
AH=BC(I,NEST+2,HJ)*F*500.
AI=AG+BC(I,NEST+2,HJ)*1000.
IF (AI .NE. 0.) AI = AI/(AI+SORT(AH+AI))
UI(I,J,NEST) = AI*BC(I,2,HJ)
V(I,J,NEST) = AI*BC(I,1,HJ)
UN(I,J,NEST) = U(I,J,NEST)
VN(I,J,NEST) = V(I,J,NEST)
31 SGW NON-ZERO FOR STEERING FLOW.
32 IF (SGW.EQ.0.) GO TO 50
AG = F*UG
AH = F*VG

```

```

IF (ST12 .GT. 0) GO TO 34
DO 33 I = 1,21
DO 33 J = 1,21
  PX(I,J,NEST) = PY(I,J,NEST)+AH
  PY(I,J,NEST) = PY(I,J,NEST)-AG
  GO TO 50
33
  C02 = C02+C02
  SI2 = SI2+SI2
DO 36 I = 1,21
  AI = (I-11)*SI2
DO 35 J = 1,21
  AJ = (J-11)*C02
  FADE = -.69314718*(AJ-AI)**2
  FADE = EXP(FADE)
IF (FADE .LE. 0.) GO TO 35
  PX(I,J,NEST) = PY(I,J,NEST)+AH*FADE
  PY(I,J,NEST) = PY(I,J,NEST)-AG*FADE
CONTINUE
35
CONTINUE
50
CONTINUE
99
RETURN
END

```

```

SUBROUTINE RUGRID (ZLA, ZLO, LSTAB, MAXI, MAXJ)
COMMON /LGRID/ ILAT(31), ILUNG(62)
DIMENSION LSTAB(MAXI), MAXJ
DIMENSION ZLA(MAXJ), ZLO(MAXI)
DO 25 J=1,4
   I1=(J-1)*15+1
   I2=I1+14
   READ 15, (ILONG(I1), I=I1, I2), KSEQ
   FURMAT(1615)
   IF (KSEQ.EQ.0) GO TO 25
   PRINT 20, KSEQ, J, (ILONG(I1), I=I1, I2)
   FORMAT(/, GRID INPUT ERROR*, 1615)
   STOP 21

25 CONTINUE
   READ 30, ILONG(61), ILONG(62), KSEQ
   FCXHAT(215, 65), I5)
   IF (KSEQ.EQ.5) GO TO 35
   PRINT 20, KSEQ, ILONG(61), ILONG(62)
   STOP 21

30   CONTINUE
   READ 15, (ILAT(I), I=1, 15), KSEQ
   IF (KSEQ.NE.1) GO TO 40
   READ 15, (ILAT(I), I=16, 30), KSEQ
   IF (KSEQ.NE.2) GO TO 40
   READ 37, ILAT(31), KSEQ
   FORMAT(15, 70X, I5)
   IF (KSEQ.EQ.3) GO TO 45
   PRINT 20, ILAT, KSEQ
   STOP 23

45 CONTINUE
   DO 55 J=1, MAXJ
      L=1ABS(ILAT(J))/100
      M=1ABS(MOD(ILAT(J), 100))
      ZLA(J)=(FLDAT(L)+FLDAT(M)/60.)*.0174532
      IF (ILAT(J).LT.0) ZLA(J)=-ZLA(J)
      READ 50, (LSTAB(I,J), I=1, 62), KSEQ
      FORMAT(10X, 62I1, 6X, 12)

```

```
IF (KSEQ.EQ.J) GO TO 55
PRINT 51,J,KSEQ,(LSTAB(I,J),I=1,62)
FORMAT(, LS INPUT ERROR, 214,2X,6211)
51 STOP 51
      IC
CONTINUE
DO 60 I=1,MAXI
L=IABS(ILONG(I))/100
M=IABS(MOD(ILONG(I),100))
ZLO(I)=FLOAT(L)*FLOAT(M/60.)*.0174532
IF(ILONG(I).LT.0) ZLO(I)=-ZLO(I)
CONTINUE
END
60
```

```
SUBROUTINE SHORE
COMMON /C3/ U(21,21,5),V(21,21,5),UN(21,21,5),VN(21,21,5)
1  PX(21,21,5),PY(21,21,5),VTN(21,21,5),ANG(21,21,5)
2  LU(21,21,5)
DO 20 I=1,2205
LU(I)=2
RETURN
END
```

20

\*\*\*\* THIS IS MAIN PROGRAM SNAP \*\*\*\*

C C MAIN PROGRAM FOR SNAPSHOT WINDS ON NESTED GRID.  
C INPUT IB IS SWITCH VARIABLE IB = 0 TO SUPPRESS PRINTING  
C OF PRESSURE FIELD  
C AND INITIAL WIND FIELD  
C  
C NZ IS NUMBER OF WIND SNAPSHOTS TO PRODUCE.  
C FOR THE QUANTITIES EQUIVALENTED TO JA, SEE COMMENTS TO  
C SUBROUTINE GRAD  
C NM IS NUMBER OF TIMES TO CYCLE WIND COMPUTATION IN  
C INNERMOST GRID NEST.  
C  
C SGW IS MAGNITUDE OF SURFACE GEOSTROPHIC WIND (MTRS/SEC)  
C AN1 IS DIRECTION OF SGW, COUNTERCLOCKWISE FROM EAST (DEG)  
C ST12 IS DIST (KM) FROM AXIS TO 1/2 MAGNITUDE OF SGW) FROM AXIS TO 1  
C DX IS GRID DISTANCE OF INNERMOST NEST (KILOMETERS).  
C NAME IS SNAPSHOT NAME FORMAT YY=YEAR, L=1ST LETTER  
C OF HURRICANE NAME).

C FOR ALL ARRAYS IN COMMON C3 :  
C 1ST DIMENSION INCREASES FROM WEST TO EAST  
C 2ND DIMENSION INCREASES FROM SOUTH TO NORTH  
C 3RD DIMENSION (NEST NUMBER) INCREASES FROM  
C INNERMOST TO OUTERMOST  
C  
COMMON/C1/NAME,NSMAP,DX,DT,F,SGW,AN1,UC,UG,VG,CS,NM,IB,ST12  
COMMON/C2/JA(15),AB(21,21,15),AC(21,7,21)  
COMMON /C3/ U(21,21,5),V(21,21,5),UN(21,21,5),VN(21,21,5)  
1 \*PX(21,21,5),PY(21,21,5),VTN(21,21,5),ANG(21,21,5)  
2 \*LU(21,21,5)  
COMMON  
\$/C4/ CDR(100,2,2),UV(100,2),TURN(100,2)  
\$/C5/ FLAT,PTH,DT,H(H,20LAND,LS,VV(100),UX(3),UV(3)),  
\$ DUV(3),K35,K2,G,GA,DEN,VV2,HL,K123,Z0,2LOG,AM,BH,CH,FF  
REAL RADIUS(4),PFAR(4)  
EQUIVALENCE (ITRACK,JA),(EYELAT,JA(2)),(EYLONG,JA(3)),  
S (IREC,JA(4)),(SPEED,JA(5)),(IGUARD,JA(G)),(CYPRESS,JA(7)),  
S (RADIUS,JA(8)),(PPFAR,JA(12))

```

C      REAL K35,K2
C      DATA LR,LP,LSNAP,LSHORE/5,6,13,14/
C      DATA PTH/300./
C      DATA HH/650./
C      DATA ZOLAND/.08/
C      DATA K35/.35/
C      DATA G/9.805/
C      DATA GARR/.0144/
C      DATA DTH/0.,-2./

C      NAMELIST/NAME1/1B,NZ
C      NAMELIST/NAME2/DTH,HH,ZOLAND,GARR,PTH,K35
C      NAMELIST/NAME3/SGW,AN1,NAME,
C      S EYELAT,EYLONG,DIREC,SPEED,EYPRES,RADIUS,PFAR,
C      S NM,DX,ST12,ITRACK,IQUAD

C      REWIND LSNAP
C      READ (LR,NAME1)
C      WRITE (LP,NAME1)
C      FORMAT (2I5)
C      READ (LR,NAME2)
C      WRITE (LP,NAME2)
C      K2=K35*2
C      GA = GARR/G

C      DO 20 NSHAP=1,NZ
C      WRITE (LP,15),
C      1      NM = 800
C      DX = 5.
C      ST12 = 0.
C      ITRACK = 0
C      IQUAD = 0
C      READ (LR,NAME3)
C      WRITE (LP,NAME3)
C      DT = 10.0*DX

```

```
PHI = EYELAT/57.29578
C   PHI IS LATITUDE IN RADIANS
      F=2.*7.29 E-5*SIN(PHI)
C   F IS CORIOLIS FORCE
      FLAT=F
      CALL CCRUSS
      WRITE (LIP,15)
15    FORMAT(1H1)
      CALL BLOWUP
      WRITE (LSNAP) UN,VN,NAME,DX,JA,SGW,AN1,ST12
      1,0TH,HH,GAKR,PTH,K35
      1,CONTINUE
C
      WRITE (LIP,25)
25    FORMAT('1 END OF SNAP/MAIN')
      END FILE LSNAP
      STOP 999
      END
```

```

C   SUBROUTINE TVEL (VTN, ANG, NBASE, IDENT, KSEQ0, LV, J20)
C   IF IN ARRAYS VTN, ANG 1ST DIMENSION INCREASES EASTWARD AND
C   2ND DIMENSION INCREASES NORTHWARD. SUBROUTINE TVEL PRINTS
C   WEST AT TOP OF PAGE, NORTH AT RIGHT OF PAGE, ETC.
C   NEST LV IS PRINTED AS INNER NEST, LV+1 AS OUTER NEST.
C   DIMENSION VTN(21,21,5),ANG(21,21,5)
C   DIMENSION KREDUC(2)/4HNOT,1H /
C   DATA LU/6/
      WRJTE (LU,100)
      100  FORMAT(1H11)
      LVP=LV+1
      120P=120+1
      WRITE (LU,200) NBASE,IDENT,KSEQ0,LVP,KREDUC(120P)
      200  FORMAT(1H0,20X,A4,5X,A4,1X,14,, LEVEL,13,5X,A4,,REDUCD)
      1190  WRITE (LU,120U) (J, J=1,12)
      1200  FORMAT(1H0,11X,12(12,UX))
      U0 1210 I=1,5
      1210  WRITE (LU,1220) ((VTN(I,J,LVP),J=1,12),(ANG(I,J,LVP),J=1,12))
      1220  FORMAT(1H0,6X,12,12(1PF5.0,5X)/9X,12(0PF5.0,5X),/1H0,/)
      1240  00 124B I=1,11
      IL=I+5
      10=2*1-1
      IE=2+1
      IF (I-11) 1242,1246,1248
      1242  WRITE (LU,1244) IL,(VTN(IL,J,LVP),J=1,5),(VTN(10,N,lv),N=1,13)
      1  *(ANG(IL,J,LVP),J=1,5),(ANG(10,N,lv),N=1,13)
      2  ,(VTN(IE,N,lv),N=1,13),(ANG(IE,N,lv),N=1,13)
      1244  FORMAT(1H0,6X,12,5(1PF5.0,5X),13(1PF5.0)/9X,5(0PF5.0,5X),
      1 13(0PF5.0)/1H0,58X,13(1PF5.0)/59X,13(0PF5.0)
      GO TO 1248
      1246  WRITE (LU,1247) IL,(VTN(IL,J,LVP),J=1,5),(VTN(10,N,lv),N=1,13)
      1  *(ANG(IL,J,LVP),J=1,5),(ANG(10,N,lv),N=1,13)
      1247  FORMAT(1H0,6X,12,5(1PF5.0,5X),13(1PF5.0)/9X,5(0PF5.0,5X),
      1 13(0PF5.0)/1H0,/)
      124A  CONTINUE

```

```

DU 1300 I=17,21
      WRITE (LU,1220) (I,(VTN(I,J,LVP),J=1,12),(ANG(I,J,LVP),J=1,12))
      WRITE (LU,1330)
      FORMAT (1H1,/1H0,15X)
      WRITE (LU,1340) (J,J=13,21)
      FORMAT(1H0,16X,9(12,8X))
      DO 1345 I=1,5
      1345 WRITE (LU,1350) ((VTN(I,J,LVP),J=13,21),I,(ANG(I,J,LVP),J=13,21))
      1350 FORMAT(1H0,13X,9(1PFS.0,5X),12/14X,9(0PF5.0,5X),/1H0,/)
      1370 DO 1378 I=1,11
      IL=1+5
      IO=2*I-1
      IE=2*I
      IF(I=11) 1372,1376,1376
      1372 WRITE (LU,1374) (VTN(10,N,LV),N=14,21),(VTN(IL,J,LVP),J=17,21),IL,
      1 (ANG(10,N,LV),N=14,21),(ANG(IL,J,LVP),J=17,21),
      2 (VTN(IE,N,LV),N=14,21),(ANG(IE,N,LV),N=14,21)
      1374 FORMAT(1H0,8X,8(1PFS.0),5X,5(1PFS.0),5X,12/9X,8(0PF5.0),5X,
      1 5(0PF5.0,5X)/1H0,8X,8(1PFS.0),9X,8(0PF5.0))
      GO TO 1378
      1376 WRITE (LU,1377) (VTN(10,N,LV),N=14,21),(VTN(IL,J,LVP),J=17,21),IL,
      1 (ANG(10,N,LV),N=14,21),(ANG(IL,J,LVP),J=17,21)
      1377 FORMAT(1H0,6X,6(1PFS.0),5X,5(1PFS.0,5X),12/9X,8(0PF5.0),5X,
      1 5(0PF5.0,5X)/1H0,/)
      1378 CONTINUE
      DO 1400 I=17,21
      1400 WRITE (LU,1350) ((VTN(I,J,LVP),J=13,21),I,(ANG(I,J,LVP),J=13,21))
      RETURN
      END

```

```

SUBROUTINE UPDOWN
COMMON /C57/ PTH,DTN,HH,ZCOEFF(3,5),LAKE,VV(100),UX(3),UV(3),
      $ DUV(3),K35,K2,G,GA,DEN,VV2,HL,K123,ZD,ZLOG,AM,BM,CH,UV(100,6),
      $ TURN(100,6),COST(100),SINT(100)
REAL K2,K35
REAL VV2(100),TOL(100),TARN(100)
IF (LAKE .EQ. 0) RETURN
BH = 1.95+K35
BH2 = BH**2
HLOG = 1.39-ALOG(HH)
DO 61 IA = 1,100
  UM = VV(IA)*COST(IA)
  VN = VV(IA)*SINT(IA)
  VPLUS = -VV(IA)*UX(VIA,1)
  VTOP = VH+VPLUS
  VV2(IA) = UM**2+VTOP**2
  TOL(IA) = 1E-4*VV2(IA)
  TARN(IA) = UM*VPLUS/(UM**2+VH*VTOP)
CONTINUE
61  DO 70 INCH = 1,LAKE
      A2 = ZCOEFF(1,INCH)
      B2 = ZCOEFF(2,INCH)
      C2 = ZCOEFF(3,INCH)
      UX(1) = 1.
      IF (AZ*BZ .NE. 0.) UX(1) = CBRT(.5*AZ/BZ)
      ZLOG = HLOG+ALOG(AZ/UX(1)+BZ+UX(1)**2+C2)
      DO 69 IA = 1,100
        UX(1) = K35*SORT(VV2(IA)/(ZLOG**2+BH2))
        ZLOG = HLOG+ALOG(AZ/UX(1)+BZ+UX(1)**2+C2)
        UV(1) = UX(1)**2/K2*(ZLOG**2+BH2)
        DUV(1) = UV(1)-VV2(IA)
        UX(2) = K35*SQRT(VV2(IA)/(ZLOG**2+BH2))
        UV(2) = UX(2)**2/K2*(ZLOG**2+BH2)
        DUV(2) = UV(2)-VV2(IA)
      KSTOP = 0
    
```

```

62      IF (DUV(1) .EQ. DUV(2)) GO TO 63
UX(3) = (UX(1)*DUV(1)+DUV(2)-UX(2)*DUV(1))/(DUV(2)-DUV(1))
ZLOG = HLOG+ALOG(AZ/UX(3)+BZ*UX(3)**2+CZ)
DUV(3) = UX(3)**2/K2*(ZLOG**2+BH2)
IF (KSTOP .NE. 0) GO TO 63
IF (ABS(DUV(3)).LT. TOL) KSTOP = 1
UX(1) = UX(2)
UX(2) = UX(3)
DUV(1) = DUV(2)
DUV(2) = DUV(3)
GO TO 62
63      UXV(IA,INCH+1) = K35*SQRT(VV2(IA)/(ZLOG**2+BH2))/VV(IA)
TURN(IA,INCH+1) = (BH-ZLOG*TARN(IA))/(ZLOG+BH*TARN(IA))
CONTINUE
69      CONTINUE
70      RETURN
END

```

```

SUBROUTINE UXV
COMMON /C57/ PTM,DTH,MM,ZCOEFF(3,5),LAKE,VV(1,100),UX(3),
$ DUV(3),K35,K2*6,GA,DEN,VY2,HL,K123,Z0,ZN06,AM,BM,CM,UXV(100,6),
$ TURN(100,6),COST(100),SINT(100)
REAL K2,K35
DATA A125/1.25/
DO 10 IA = 1,100
  VV(IA) = FLOAT(IA)/A125
10 CONTINUE
DEN = 6*K2*DTH
DO 40 IA = 1,100
  IB = 101-IA
  VV2 = VV(IB)**2
  TOL = 1E-4*VV2
  IF (IA .NE. 1) GO TO 20
  CH = 2.55
  K123 = 1
  UX(1) = 2.74
  CALL ABCCC
  K123 = 2
  UX(2) = UX(1)*80./SQRT(UV(1))
  GO TO 30
20  UX(1) = UX(3)
  UV(1) = UV(3)
  DUV(1) = UDV(1)-VV2
  K123 = 2
  UX(2) = UX(1)*VV(IB)/VV(IB+1)
  CALL ABCCC
  KSTOP = 0
  K123 = 3
  IF (DUV(1).EQ.0. DUV(2)) GO TO 32
  UX(3) = AMAX1(.5*AMIN1(UX(1),UX(2)),
$ AMIN1(2.*AMAX1(UX(1),UX(2)),
$ UX(1)*DUV(2)-UX(2)*DUV(1))/(DUV(2)-DUV(1)))

```

```

CALL AHCCC
IF (KSTOP .NE. 0) GO TO 32
IF (ABS(DUV(3)) .LT. TOL) KSTOP = 1
UX(1) = UX(2)
UX(2) = UX(3)
DUV(1) = DUV(2)
DUV(2) = DUV(3)
GO TO 31
AB = SQR((ZLOG+AM)**2+BH**2)
UXV(1B+1) = K35/AB
TURN(1B+1) = ATAN(BH/(ZLOG+AM))
COST(1B) = -IZLOG+AM)/AB
SINT(1B) = BH/AB
CONTINUE
RETURN
END

```

32

40

**APPENDIX C: PROGRAM HIST LISTING  
OF TEST STORM (BETSY) INPUT  
AND SAMPLE ANNOTATED OUTPUT**



REPACK 'CORPS' -  
FUFUPUR 27A3A 633 ST 781 10/22/79 13:29:49  
6000 PACK. 1C71512.1UC61515H26K1515

ISPAÑA PREP.

c4

1. *WAS* *STARTED* *IN* *1901*.  
2. *WAS* *STARTED* *IN* *1901*.  
3. *WAS* *STARTED* *IN* *1901*.  
4. *WAS* *STARTED* *IN* *1901*.  
5. *WAS* *STARTED* *IN* *1901*.  
6. *WAS* *STARTED* *IN* *1901*.

COLPUS, LANTERNISCH & STRAT



Panel 4 (USPCH, MAMC/H/MS/15) -  
- 001452546  
- 00090909  
- 00000000

卷之三

ADD. P CORPUS-LAURENTIANUS II.

charname	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100	101	102	103	104	105	106	107	108	109	110	111	112	113	114	115	116	117	118	119	120	121	122	123	124	125	126	127	128	129	130	131	132	133	134	135	136	137	138	139	140	141	142	143	144	145	146	147	148	149	150	151	152	153	154	155	156	157	158	159	160	161	162	163	164	165	166	167	168	169	170	171	172	173	174	175	176	177	178	179	180	181	182	183	184	185	186	187	188	189	190	191	192	193	194	195	196	197	198	199	200	201	202	203	204	205	206	207	208	209	210	211	212	213	214	215	216	217	218	219	220	221	222	223	224	225	226	227	228	229	230	231	232	233	234	235	236	237	238	239	240	241	242	243	244	245	246	247	248	249	250	251	252	253	254	255	256	257	258	259	260	261	262	263	264	265	266	267	268	269	270	271	272	273	274	275	276	277	278	279	280	281	282	283	284	285	286	287	288	289	290	291	292	293	294	295	296	297	298	299	300	301	302	303	304	305	306	307	308	309	310	311	312	313	314	315	316	317	318	319	320	321	322	323	324	325	326	327	328	329	330	331	332	333	334	335	336	337	338	339	340	341	342	343	344	345	346	347	348	349	350	351	352	353	354	355	356	357	358	359	360	361	362	363	364	365	366	367	368	369	370	371	372	373	374	375	376	377	378	379	380	381	382	383	384	385	386	387	388	389	390	391	392	393	394	395	396	397	398	399	400	401	402	403	404	405	406	407	408	409	410	411	412	413	414	415	416	417	418	419	420	421	422	423	424	425	426	427	428	429	430	431	432	433	434	435	436	437	438	439	440	441	442	443	444	445	446	447	448	449	450	451	452	453	454	455	456	457	458	459	460	461	462	463	464	465	466	467	468	469	470	471	472	473	474	475	476	477	478	479	480	481	482	483	484	485	486	487	488	489	490	491	492	493	494	495	496	497	498	499	500	501	502	503	504	505	506	507	508	509	510	511	512	513	514	515	516	517	518	519	520	521	522	523	524	525	526	527	528	529	530	531	532	533	534	535	536	537	538	539	540	541	542	543	544	545	546	547	548	549	550	551	552	553	554	555	556	557	558	559	560	561	562	563	564	565	566	567	568	569	570	571	572	573	574	575	576	577	578	579	580	581	582	583	584	585	586	587	588	589	590	591	592	593	594	595	596	597	598	599	600	601	602	603	604	605	606	607	608	609	610	611	612	613	614	615	616	617	618	619	620	621	622	623	624	625	626	627	628	629	630	631	632	633	634	635	636	637	638	639	640	641	642	643	644	645	646	647	648	649	650	651	652	653	654	655	656	657	658	659	660	661	662	663	664	665	666	667	668	669	670	671	672	673	674	675	676	677	678	679	680	681	682	683	684	685	686	687	688	689	690	691	692	693	694	695	696	697	698	699	700	701	702	703	704	705	706	707	708	709	710	711	712	713	714	715	716	717	718	719	720	721	722	723	724	725	726	727	728	729	730	731	732	733	734	735	736	737	738	739	740	741	742	743	744	745	746	747	748	749	750	751	752	753	754	755	756	757	758	759	760	761	762	763	764	765	766	767	768	769	770	771	772	773	774	775	776	777	778	779	780	781	782	783	784	785	786	787	788	789	790	791	792	793	794	795	796	797	798	799	800	801	802	803	804	805	806	807	808	809	810	811	812	813	814	815	816	817	818	819	820	821	822	823	824	825	826	827	828	829	830	831	832	833	834	835	836	837	838	839	840	841	842	843	844	845	846	847	848	849	850	851	852	853	854	855	856	857	858	859	860	861	862	863	864	865	866	867	868	869	870	871	872	873	874	875	876	877	878	879	880	881	882	883	884	885	886	887	888	889	890	891	892	893	894	895	896	897	898	899	900	901	902	903	904	905	906	907	908	909	910	911	912	913	914	915	916	917	918	919	920	921	922	923	924	925	926	927	928	929	930	931	932	933	934	935	936	937	938	939	940	941	942	943	944	945	946	947	948	949	950	951	952	953	954	955	956	957	958	959	960	961	962	963	964	965	966	967	968	969	970	971	972	973	974	975	976	977	978	979	980	981	982	983	984	985	986	987	988	989	990	991	992	993	994	995	996	997	998	999	1000
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Sample listings of time history of  
surface wind field at individual stations

65B1      Storm identification  
65091001    Year, month, day, hour  
W1          Wind speed in knots  
TH1         Wind direction (meteorological) in degrees  
D           Distance of station from eye of hurricane (km)  
AL          Bearing of station from eye (degrees)  
UST         Friction velocity (cm/sec)  
STA         Station number, latitude (deg, min) longitude (deg, min)  
EYE         Eye latitude and longitude (deg, min)  
TERR        Terrain roughness category  
H           Anemometer height at station











6561	65090901	612 21.55	TH12	02	662.1	AL=315.0	UST1	60.4	514	62 30	V1 = 25
6561	65090902	612 22.07	TH12	02	6130.1	AL=316.0	UST1	61.0	514	62 30	V1 = 26
6561	65090903	612 22.61	TH12	02	6130.2	AL=317.0	UST1	61.6	514	62 30	V1 = 26
6561	65090904	612 23.05	TH12	02	6130.4	AL=318.0	UST1	62.0	514	62 30	V1 = 26
6561	65090905	612 23.49	TH12	02	6130.7	AL=319.0	UST1	62.4	514	62 30	V1 = 26
6561	65090906	612 24.23	TH12	02	6131.1	AL=320.1	UST1	63.0	514	62 30	V1 = 26
6561	65090907	612 24.62	TH12	02	6131.3	AL=321.0	UST1	63.6	514	62 30	V1 = 26
6561	65090908	612 25.56	TH12	02	6131.6	AL=322.0	UST1	64.2	514	62 30	V1 = 26
6561	65090909	612 26.22	TH12	02	6131.9	AL=323.0	UST1	64.8	514	62 30	V1 = 26
6561	65090910	612 27.01	TH12	02	6132.2	AL=324.0	UST1	65.4	514	62 30	V1 = 26
6561	65090911	612 27.65	TH12	02	6132.5	AL=325.0	UST1	66.0	514	62 30	V1 = 26
6561	65090912	612 28.46	TH12	02	6132.8	AL=326.0	UST1	66.6	514	62 30	V1 = 26
6561	65090913	612 29.46	TH12	02	6133.1	AL=327.0	UST1	67.2	514	62 30	V1 = 26
6561	65090914	612 30.06	TH12	02	6133.4	AL=328.0	UST1	67.8	514	62 30	V1 = 26
6561	65090915	612 30.66	TH12	02	6133.7	AL=329.0	UST1	68.4	514	62 30	V1 = 26
6561	65090916	612 31.26	TH12	02	6134.0	AL=330.0	UST1	69.0	514	62 30	V1 = 26
6561	65090917	612 31.87	TH12	02	6134.3	AL=331.0	UST1	69.6	514	62 30	V1 = 26
6561	65090918	612 32.47	TH12	02	6134.6	AL=332.0	UST1	70.2	514	62 30	V1 = 26
6561	65090919	612 33.08	TH12	02	6134.9	AL=333.0	UST1	70.8	514	62 30	V1 = 26
6561	65090920	612 33.68	TH12	02	6135.2	AL=334.0	UST1	71.4	514	62 30	V1 = 26
6561	65090921	612 34.28	TH12	02	6135.5	AL=335.0	UST1	72.0	514	62 30	V1 = 26
6561	65090922	612 34.89	TH12	02	6135.8	AL=336.0	UST1	72.6	514	62 30	V1 = 26
6561	65090923	612 35.49	TH12	02	6136.1	AL=337.0	UST1	73.2	514	62 30	V1 = 26
6561	65090924	612 36.09	TH12	02	6136.4	AL=338.0	UST1	73.8	514	62 30	V1 = 26
6561	65090925	612 36.69	TH12	02	6136.7	AL=339.0	UST1	74.4	514	62 30	V1 = 26
6561	65090926	612 37.29	TH12	02	6137.0	AL=340.0	UST1	75.0	514	62 30	V1 = 26
6561	65090927	612 37.89	TH12	02	6137.3	AL=341.0	UST1	75.6	514	62 30	V1 = 26
6561	65090928	612 38.49	TH12	02	6137.6	AL=342.0	UST1	76.2	514	62 30	V1 = 26
6561	65090929	612 39.09	TH12	02	6137.9	AL=343.0	UST1	76.8	514	62 30	V1 = 26
6561	65090930	612 39.69	TH12	02	6138.2	AL=344.0	UST1	77.4	514	62 30	V1 = 26
6561	65090931	612 40.29	TH12	02	6138.5	AL=345.0	UST1	78.0	514	62 30	V1 = 26
6561	65090932	612 40.89	TH12	02	6138.8	AL=346.0	UST1	78.6	514	62 30	V1 = 26
6561	65090933	612 41.49	TH12	02	6139.1	AL=347.0	UST1	79.2	514	62 30	V1 = 26
6561	65090934	612 42.09	TH12	02	6139.4	AL=348.0	UST1	79.8	514	62 30	V1 = 26
6561	65090935	612 42.69	TH12	02	6139.7	AL=349.0	UST1	80.4	514	62 30	V1 = 26
6561	65090936	612 43.29	TH12	02	6140.0	AL=350.0	UST1	81.0	514	62 30	V1 = 26
6561	65090937	612 43.89	TH12	02	6140.3	AL=351.0	UST1	81.6	514	62 30	V1 = 26
6561	65090938	612 44.49	TH12	02	6140.6	AL=352.0	UST1	82.2	514	62 30	V1 = 26
6561	65090939	612 45.09	TH12	02	6140.9	AL=353.0	UST1	82.8	514	62 30	V1 = 26
6561	65090940	612 45.69	TH12	02	6141.2	AL=354.0	UST1	83.4	514	62 30	V1 = 26
6561	65090941	612 46.29	TH12	02	6141.5	AL=355.0	UST1	84.0	514	62 30	V1 = 26
6561	65090942	612 46.89	TH12	02	6141.8	AL=356.0	UST1	84.6	514	62 30	V1 = 26
6561	65090943	612 47.49	TH12	02	6142.1	AL=357.0	UST1	85.2	514	62 30	V1 = 26
6561	65090944	612 48.09	TH12	02	6142.4	AL=358.0	UST1	85.8	514	62 30	V1 = 26
6561	65090945	612 48.69	TH12	02	6142.7	AL=359.0	UST1	86.4	514	62 30	V1 = 26
6561	65090946	612 49.29	TH12	02	6143.0	AL=360.0	UST1	87.0	514	62 30	V1 = 26
6561	65090947	612 49.89	TH12	02	6143.3	AL=361.0	UST1	87.6	514	62 30	V1 = 26
6561	65090948	612 50.49	TH12	02	6143.6	AL=362.0	UST1	88.2	514	62 30	V1 = 26
6561	65090949	612 51.09	TH12	02	6143.9	AL=363.0	UST1	88.8	514	62 30	V1 = 26
6561	65090950	612 51.69	TH12	02	6144.2	AL=364.0	UST1	89.4	514	62 30	V1 = 26
6561	65090951	612 52.29	TH12	02	6144.5	AL=365.0	UST1	90.0	514	62 30	V1 = 26
6561	65090952	612 52.89	TH12	02	6144.8	AL=366.0	UST1	90.6	514	62 30	V1 = 26
6561	65090953	612 53.49	TH12	02	6145.1	AL=367.0	UST1	91.2	514	62 30	V1 = 26
6561	65090954	612 54.09	TH12	02	6145.4	AL=368.0	UST1	91.8	514	62 30	V1 = 26
6561	65090955	612 54.69	TH12	02	6145.7	AL=369.0	UST1	92.4	514	62 30	V1 = 26
6561	65090956	612 55.29	TH12	02	6146.0	AL=370.0	UST1	93.0	514	62 30	V1 = 26
6561	65090957	612 55.89	TH12	02	6146.3	AL=371.0	UST1	93.6	514	62 30	V1 = 26
6561	65090958	612 56.49	TH12	02	6146.6	AL=372.0	UST1	94.2	514	62 30	V1 = 26
6561	65090959	612 57.09	TH12	02	6146.9	AL=373.0	UST1	94.8	514	62 30	V1 = 26
6561	65090960	612 57.69	TH12	02	6147.2	AL=374.0	UST1	95.4	514	62 30	V1 = 26
6561	65090961	612 58.29	TH12	02	6147.5	AL=375.0	UST1	96.0	514	62 30	V1 = 26
6561	65090962	612 58.89	TH12	02	6147.8	AL=376.0	UST1	96.6	514	62 30	V1 = 26
6561	65090963	612 59.49	TH12	02	6148.1	AL=377.0	UST1	97.2	514	62 30	V1 = 26
6561	65090964	612 60.09	TH12	02	6148.4	AL=378.0	UST1	97.8	514	62 30	V1 = 26
6561	65090965	612 60.69	TH12	02	6148.7	AL=379.0	UST1	98.4	514	62 30	V1 = 26
6561	65090966	612 61.29	TH12	02	6149.0	AL=380.0	UST1	99.0	514	62 30	V1 = 26
6561	65090967	612 61.89	TH12	02	6149.3	AL=381.0	UST1	99.6	514	62 30	V1 = 26
6561	65090968	612 62.49	TH12	02	6149.6	AL=382.0	UST1	100.2	514	62 30	V1 = 26
6561	65090969	612 63.09	TH12	02	6149.9	AL=383.0	UST1	100.8	514	62 30	V1 = 26
6561	65090970	612 63.69	TH12	02	6150.2	AL=384.0	UST1	101.4	514	62 30	V1 = 26
6561	65090971	612 64.29	TH12	02	6150.5	AL=385.0	UST1	102.0	514	62 30	V1 = 26
6561	65090972	612 64.89	TH12	02	6150.8	AL=386.0	UST1	102.6	514	62 30	V1 = 26
6561	65090973	612 65.49	TH12	02	6151.1	AL=387.0	UST1	103.2	514	62 30	V1 = 26
6561	65090974	612 66.09	TH12	02	6151.4	AL=388.0	UST1	103.8	514	62 30	V1 = 26
6561	65090975	612 66.69	TH12	02	6151.7	AL=389.0	UST1	104.4	514	62 30	V1 = 26
6561	65090976	612 67.29	TH12	02	6152.0	AL=390.0	UST1	105.0	514	62 30	V1 = 26
6561	65090977	612 67.89	TH12	02	6152.3	AL=391.0	UST1	105.6	514	62 30	V1 = 26
6561	65090978	612 68.49	TH12	02	6152.6	AL=392.0	UST1	106.2	514	62 30	V1 = 26
6561	65090979	612 69.09	TH12	02	6152.9	AL=393.0	UST1	106.8	514	62 30	V1 = 26
6561	65090980	612 69.69	TH12	02	6153.2	AL=394.0	UST1	107.4	514	62 30	V1 = 26
6561	65090981	612 70.29	TH12	02	6153.5	AL=395.0	UST1	108.0	514	62 30	V1 = 26
6561	65090982	612 70.89	TH12	02	6153.8	AL=396.0	UST1	108.6	514	62 30	V1 = 26
6561	65090983	612 71.49	TH12	02	6154.1	AL=397.0	UST1	109.2	514	62 30	V1 = 26
6561	65090984	612 72.09	TH12	02	6154.4	AL=398.0	UST1	109.8	514	62 30	V1 = 26
6561	65090985	612 72.69	TH12	02	6154.7	AL=399.0	UST1	110.4	514	62 30	V1 = 26
6561	65090986	612 73.29	TH12	02	6155.0	AL=400.0	UST1	111.0	514	62 30	V1 = 26
6561	65090987	612 73.89	TH12	02	6155.3	AL=401.0	UST1	111.6	514	62 30	V1 = 26
6561	65090988	612 74.49	TH12	02	6155.6	AL=402.0	UST1	112.2	514	62 30	V1 = 26
6561	65090989	612 75.09	TH12	02	6155.9	AL=403.0	UST1	112.8	514	62 30	V1 = 26
6561	65090990	612 75.									

Sample listing of surface wind speed direction on Wes grid:

. Six pages of output are required to list the wind field at one time level. Grid points are identified on each page by latitude (deg, min), left side margin, and longitude (deg, min), top margin. At each grid point location are printed the wind speed at 10 meter height in knots (top), the wind direction in degrees (middle) and the terrain classification code of the grid points (bottom). The sample listing shown is for the 24th hour of the test Betsy simulation, for the rectangular grid system provided by WES.

ՀԱՅՈՒԹՅԱՆ ՎՐԱ ԵՎ ՎՐԱ ԵՎ ՀԱՅՈՒԹՅԱՆ

S. M. HUSSAIN AND J. S. LEE

STORY CSII LARI P.O.I. WINES AT HOUK 24 1ST HOUR 15 65798761CB1

8927	8929	8928	8916	8912	8909	8905	8901	8816	8802	8807	8802	8802	8826	8821	8815	8810	8810	8805	8801	8757	8752	8746	8739
3865	39.6	39.3	39.1	38.9	38.6	38.4	38.1	37.8	37.6	37.3	37.0	36.7	36.3	35.7	35.3	34.9	33.6	32.2	31.4	30.9	30.4	29.6	30.5
3865	73.6	73.3	73.1	72.8	72.6	72.3	72.0	71.7	71.4	71.1	70.8	70.5	70.2	69.9	69.6	69.3	69.0	68.7	68.4	68.1	67.8	67.5	67.2
3865	66.9	66.6	66.4	66.2	66.0	65.8	65.6	65.4	65.2	65.0	64.8	64.6	64.4	64.2	64.0	63.8	63.6	63.4	63.2	63.0	62.8	62.6	62.4
3869	19.9	19.6	19.3	19.0	18.7	18.4	18.1	17.8	17.5	17.2	16.9	16.6	16.3	16.0	15.7	15.4	15.1	14.8	14.5	14.2	13.9	13.6	13.3
3879	75.2	74.9	74.6	74.3	74.0	73.7	73.4	73.1	72.8	72.5	72.2	71.9	71.6	71.3	71.0	70.7	70.4	70.1	69.8	69.5	69.2	68.9	68.6
3883	92.2	91.9	91.6	91.3	91.0	90.7	90.4	90.1	89.8	89.5	89.2	88.9	88.6	88.3	88.0	87.7	87.4	87.1	86.8	86.5	86.2	85.9	85.6
3883	70.0	69.7	69.4	69.1	68.8	68.5	68.2	67.9	67.6	67.3	67.0	66.7	66.4	66.1	65.8	65.5	65.2	64.9	64.6	64.3	64.0	63.7	63.4
3826	54.6	54.3	54.0	53.7	53.4	53.1	52.8	52.5	52.2	51.9	51.6	51.3	51.0	50.7	50.4	50.1	49.8	49.5	49.2	48.9	48.6	48.3	48.0
3826	86.6	86.3	86.0	85.7	85.4	85.1	84.8	84.5	84.2	83.9	83.6	83.3	83.0	82.7	82.4	82.1	81.8	81.5	81.2	80.9	80.6	80.3	80.0
3821	55.5	55.2	54.9	54.6	54.3	54.0	53.7	53.4	53.1	52.8	52.5	52.2	51.9	51.6	51.3	51.0	50.7	50.4	50.1	49.8	49.5	49.2	48.9
3821	66.3	66.0	65.7	65.4	65.1	64.8	64.5	64.2	63.9	63.6	63.3	63.0	62.7	62.4	62.1	61.8	61.5	61.2	60.9	60.6	60.3	60.0	59.7
3814	36.6	36.3	36.0	35.7	35.4	35.1	34.8	34.5	34.2	33.9	33.6	33.3	33.0	32.7	32.4	32.1	31.8	31.5	31.2	30.9	30.6	30.3	30.0
3816	86.6	86.3	86.0	85.7	85.4	85.1	84.8	84.5	84.2	83.9	83.6	83.3	83.0	82.7	82.4	82.1	81.8	81.5	81.2	80.9	80.6	80.3	80.0
3813	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0	53.7	53.4	53.1	52.8	52.5	52.2	51.9	51.6	51.3	51.0
3813	92.5	92.2	91.9	91.6	91.3	91.0	90.7	90.4	90.1	89.8	89.5	89.2	88.9	88.6	88.3	88.0	87.7	87.4	87.1	86.8	86.5	86.2	85.9
3819	60.8	60.5	60.2	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3
3819	92.6	92.3	92.0	91.7	91.4	91.1	90.8	90.5	90.2	89.9	89.6	89.3	89.0	88.7	88.4	88.1	87.8	87.5	87.2	86.9	86.6	86.3	86.0
3808	60.2	59.9	59.6	59.3	59.0	58.7	58.4	58.1	57.8	57.5	57.2	56.9	56.6	56.3	56.0	55.7	55.4	55.1	54.8	54.5	54.2	53.9	53.6
3808	92.6	92.3	92.0	91.7	91.4	91.1	90.8	90.5	90.2	89.9	89.6	89.3	89.0	88.7	88.4	88.1	87.8	87.5	87.2	86.9	86.6	86.3	86.0
3806	60.6	59.9	59.2	58.5	57.8	57.1	56.4	55.7	55.0	54.3	53.6	52.9	52.2	51.5	50.8	50.1	49.4	48.7	48.0	47.3	46.6	45.9	45.2
3806	92.5	92.2	91.9	91.6	91.3	91.0	90.7	90.4	90.1	89.8	89.5	89.2	88.9	88.6	88.3	88.0	87.7	87.4	87.1	86.8	86.5	86.2	85.9
3803	55.6	55.3	55.0	54.7	54.4	54.1	53.8	53.5	53.2	52.9	52.6	52.3	52.0	51.7	51.4	51.1	50.8	50.5	50.2	49.9	49.6	49.3	49.0
3803	91.7	91.4	91.1	90.8	90.5	90.2	89.9	89.6	89.3	89.0	88.7	88.4	88.1	87.8	87.5	87.2	86.9	86.6	86.3	86.0	85.7	85.4	85.1
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	92.5	92.2	91.9	91.6	91.3	91.0	90.7	90.4	90.1	89.8	89.5	89.2	88.9	88.6	88.3	88.0	87.7	87.4	87.1	86.8	86.5	86.2	85.9
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	91.7	91.4	91.1	90.8	90.5	90.2	89.9	89.6	89.3	89.0	88.7	88.4	88.1	87.8	87.5	87.2	86.9	86.6	86.3	86.0	85.7	85.4	85.1
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	92.5	92.2	91.9	91.6	91.3	91.0	90.7	90.4	90.1	89.8	89.5	89.2	88.9	88.6	88.3	88.0	87.7	87.4	87.1	86.8	86.5	86.2	85.9
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	91.7	91.4	91.1	90.8	90.5	90.2	89.9	89.6	89.3	89.0	88.7	88.4	88.1	87.8	87.5	87.2	86.9	86.6	86.3	86.0	85.7	85.4	85.1
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	92.5	92.2	91.9	91.6	91.3	91.0	90.7	90.4	90.1	89.8	89.5	89.2	88.9	88.6	88.3	88.0	87.7	87.4	87.1	86.8	86.5	86.2	85.9
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	91.7	91.4	91.1	90.8	90.5	90.2	89.9	89.6	89.3	89.0	88.7	88.4	88.1	87.8	87.5	87.2	86.9	86.6	86.3	86.0	85.7	85.4	85.1
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	92.5	92.2	91.9	91.6	91.3	91.0	90.7	90.4	90.1	89.8	89.5	89.2	88.9	88.6	88.3	88.0	87.7	87.4	87.1	86.8	86.5	86.2	85.9
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	91.7	91.4	91.1	90.8	90.5	90.2	89.9	89.6	89.3	89.0	88.7	88.4	88.1	87.8	87.5	87.2	86.9	86.6	86.3	86.0	85.7	85.4	85.1
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	92.5	92.2	91.9	91.6	91.3	91.0	90.7	90.4	90.1	89.8	89.5	89.2	88.9	88.6	88.3	88.0	87.7	87.4	87.1	86.8	86.5	86.2	85.9
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	91.7	91.4	91.1	90.8	90.5	90.2	89.9	89.6	89.3	89.0	88.7	88.4	88.1	87.8	87.5	87.2	86.9	86.6	86.3	86.0	85.7	85.4	85.1
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	92.5	92.2	91.9	91.6	91.3	91.0	90.7	90.4	90.1	89.8	89.5	89.2	88.9	88.6	88.3	88.0	87.7	87.4	87.1	86.8	86.5	86.2	85.9
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	57.6	57.3	57.0	56.7	56.4	56.1	55.8	55.5	55.2	54.9	54.6	54.3	54.0
2959	91.7	91.4	91.1	90.8	90.5	90.2	89.9	89.6	89.3	89.0	88.7	88.4	88.1	87.8	87.5	87.2	86.9	86.6	86.3	86.0	85.7	85.4	85.1
2959	60.6	60.3	60.0	59.7	59.4	59.1	58.8	58.5	58.2	57.9	5												

## STOP 6591 LAST FOR 1 UNITS AT HOUR 26 151 HOUR 15 6590901CD1

9476	94.86	9352	9340	9327	9310	9304	9252	9241	9229	9218	9209	9201	9155	9150	9146	9143	9139	9134	9130	9127	9121	9116	
2905	18.1	28.0	21.3	28.0	36.1	31.5	33.0	29.0	38.9	39.0	35.0	39.1	41.1	42.0	43.5	40.0	41.0	40.0	46.0	47.0	46.0	2905	
1595	7.	2.	2.	7.	5.	3.	1.	350	352.	350.	350.	353.	353.	353.	353.	353.	351.	351.	350.	350.	350.	2945	
2906	14.4	19.0	21.2	21.4	26.0	31.4	33.0	35.1	36.0	36.0	36.0	37.0	41.1	42.0	40.1	45.3	46.2	47.2	47.5	48.4	48.4	2946	
2906	1.	3.	6.	6.	3.	1.	359.	359.	359.	359.	359.	359.	359.	359.	359.	359.	351.	351.	350.	350.	350.	2946	
2936	16.5	22.7	27.0	28.3	29.1	31.2	32.9	35.0	36.9	35.0	35.0	37.4	39.0	40.9	42.3	49.0	50.9	53.0	53.0	53.0	53.0	2936	
2936	4.	16.	7.	9.	11.	36.6.	357.	356.	354.	346.	346.	347.	347.	347.	347.	347.	351.	351.	352.	352.	352.	352.	2936
2932	24.3	21.5	24.5	24.4	29.7	31.1	32.1	34.1	34.1	35.1	35.1	36.5	36.5	36.5	36.5	36.5	37.3	37.3	37.3	37.3	37.3	37.3	
2932	12.	19.	6.	6.	8.	354.	354.	354.	354.	354.	354.	349.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	2932
2927	24.1	25.4	26.7	26.9	29.4	30.9	32.6	34.1	36.6	36.9	36.9	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	2927	
2927	11.	1.	4.	2.	2.	356.	356.	356.	356.	356.	356.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	2927
2922	23.9	25.2	26.5	27.8	29.2	30.7	32.3	34.1	36.3	36.7	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	2922	
2922	5.	6.	3.	3.	3.	357.	357.	357.	357.	357.	357.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	2922
2918	23.7	25.9	26.9	27.6	29.1	30.5	32.0	34.1	36.6	36.9	36.9	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	37.0	2918	
2918	6.	5.	1.	1.	1.	355.	356.	355.	356.	356.	356.	346.	346.	346.	346.	346.	346.	346.	346.	346.	346.	346.	2918
2919	23.4	24.5	24.5	26.1	27.4	28.9	30.2	31.9	33.7	35.7	38.1	38.1	38.5	38.5	38.5	38.5	38.5	38.5	38.5	38.5	38.5	2919	
2919	7.	4.	4.	6.	5.	357.	357.	357.	357.	357.	357.	346.	346.	346.	346.	346.	346.	346.	346.	346.	346.	346.	2919
2909	23.4	24.6	25.9	27.2	28.6	29.6	31.0	33.5	35.2	37.7	40.1	42.2	44.6	45.5	46.9	48.0	48.7	49.9	50.9	51.3	52.3	2909	
2909	4.	4.	4.	5.	5.	353.	356.	356.	356.	356.	356.	346.	346.	346.	346.	346.	346.	346.	346.	346.	346.	346.	2909
2903	23.1	24.3	25.6	26.9	28.3	29.6	31.3	33.0	34.7	37.1	39.5	41.6	43.4	44.9	46.9	47.1	47.2	48.3	49.2	50.2	50.2	2903	
2903	4.	0.	357.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	2903
2956	22.7	23.9	25.3	26.4	27.6	29.1	30.5	32.0	33.5	36.9	38.9	39.0	40.0	40.0	40.0	40.0	40.0	40.0	40.0	40.0	40.0	2956	
2956	-2.	2.	355.	355.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	2956
2807	22.2	23.5	24.7	25.9	27.3	28.6	30.0	31.7	33.5	35.5	37.4	39.4	41.2	42.5	43.6	44.5	45.2	46.0	46.6	47.2	47.9	2807	
2807	-2.	0.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	356.	2807
2837	21.6	22.8	24.1	25.3	26.6	27.7	29.3	30.2	32.3	34.2	36.2	37.2	37.5	38.0	38.0	38.0	38.2	38.2	38.2	38.2	38.2	2837	
2837	356.	356.	354.	349.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	345.	2837
2627	20.4	21.1	22.9	24.5	25.9	27.1	28.9	29.9	31.3	33.1	34.1	35.1	36.1	36.9	37.6	38.0	38.9	39.6	40.1	40.7	41.2	2627	
2627	355.	351.	347.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	343.	2627
2816	20.9	21.3	22.6	23.7	25.0	26.1	27.1	28.9	30.2	31.6	32.3	32.3	32.3	32.3	32.3	32.3	32.3	32.3	32.3	32.3	32.3	2816	
2816	353.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	348.	2816	
2600	19.0	20.3	21.6	22.7	23.7	25.0	26.1	26.3	27.1	28.9	29.9	30.1	30.3	30.7	30.9	31.1	31.3	31.5	31.7	31.9	31.9	2600	
2600	350.	345.	341.	337.	332.	329.	325.	320.	316.	316.	316.	316.	316.	316.	316.	316.	316.	316.	316.	316.	316.	2600	
2750	16.1	19.2	20.4	21.5	22.4	23.7	24.9	26.2	27.1	28.0	28.9	29.8	30.0	30.4	30.4	30.4	30.5	30.5	30.5	30.5	30.5	2750	
2750	347.	342.	337.	333.	328.	325.	325.	325.	325.	325.	325.	325.	325.	325.	325.	325.	325.	325.	325.	325.	325.	325.	2750

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THE JOURNAL OF CLIMATE

	8923	8920	8918	8916	8912	8909	8905	8901	8816	8852	8847	8842	8837	8832	8826	8821	8815	8810	8805	8801	8757	8752	8747	8740			
2945	65.6	66.8	61.6	60.7	59.7	59.7	59.0	58.1	57.2	56.9	56.9	56.5	57.4	56.3	55.2	53.7	52.7	51.6	50.7	49.9	49.3	48.6	47.8	46.5	25.5		
2946	69.1	65.7	61.9	60.6	59.9	59.3	58.2	57.4	56.6	55.6	55.6	54.9	54.7	54.3	53.9	53.1	52.1	51.9	51.7	51.0	50.3	50.0	49.3	48.7	47.8	22.5	
2936	64.2	63.1	62.1	61.1	59.9	59.1	58.1	57.3	56.8	56.0	55.8	55.5	55.5	55.5	55.5	55.2	54.9	54.3	53.7	53.1	52.0	51.8	51.0	50.6	50.0	24.4	
2937	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	126.	
2932	63.7	62.9	62.7	60.9	57.9	57.5	57.0	56.8	56.6	56.4	56.4	56.2	56.2	56.2	56.2	56.2	56.1	56.1	56.1	56.1	56.1	56.1	56.1	56.1	56.1	23.6	
2933	132.	131.	130.	129.	130.	132.	125.	122.	122.	123.	123.	125.	127.	127.	127.	127.	127.	127.	127.	127.	127.	127.	127.	127.	127.	127.	24.0
2921	63.6	62.3	61.6	60.7	60.7	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	60.5	23.2	
2922	134.	137.	136.	137.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	136.	23.2
2922	60.4	59.5	57.5	56.5	54.5	53.0	52.0	51.0	50.0	49.0	49.0	48.0	47.0	46.0	45.0	44.0	43.0	42.0	41.0	40.0	39.0	38.0	37.0	36.0	35.0	23.2	
2922	144.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	143.	23.2
2910	59.1	58.7	57.5	57.5	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	23.2	
2910	145.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	23.2
2910	59.1	58.7	57.5	57.5	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	56.7	23.2	
2910	145.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	23.2
2909	62.0	56.1	55.7	55.7	55.3	54.6	54.2	53.8	53.4	53.0	52.6	52.2	51.8	51.4	51.0	50.6	50.2	50.0	49.8	49.6	49.4	49.2	49.0	48.8	48.6	48.4	23.2
2909	145.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	23.2
2909	62.0	56.1	55.7	55.7	55.3	54.6	54.2	53.8	53.4	53.0	52.6	52.2	51.8	51.4	51.0	50.6	50.2	50.0	49.8	49.6	49.4	49.2	49.0	48.8	48.6	48.4	23.2
2909	145.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	146.	23.2
2903	61.1	60.7	60.5	60.5	59.8	59.1	58.7	58.1	57.7	57.1	56.6	56.1	55.7	55.1	54.6	54.0	53.2	52.1	51.3	50.3	49.5	48.8	48.0	47.2	46.4	45.6	23.0
2903	169.	167.	165.	163.	161.	159.	157.	156.	154.	153.	152.	150.	148.	146.	144.	142.	140.	138.	136.	134.	132.	130.	128.	126.	124.	122.	120.
2856	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	58.6	
2856	176.	173.	171.	169.	167.	165.	163.	161.	159.	156.	154.	152.	150.	148.	146.	144.	142.	140.	138.	136.	134.	132.	130.	128.	126.	124.	122.
2807	56.2	56.9	56.2	55.9	55.7	55.6	55.3	54.7	54.2	53.6	53.0	52.3	51.7	51.1	50.5	50.0	49.5	49.3	48.4	47.7	47.0	46.3	45.7	45.0	44.3	43.6	23.2
2807	184.	181.	179.	176.	174.	172.	170.	168.	166.	164.	162.	160.	158.	156.	154.	152.	150.	148.	146.	144.	142.	140.	138.	136.	134.	132.	130.
2807	53.9	52.9	52.9	52.7	52.7	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	52.6	
2807	191.	186.	181.	179.	176.	174.	172.	170.	168.	166.	164.	162.	160.	158.	156.	154.	152.	150.	148.	146.	144.	142.	140.	138.	136.	134.	132.
2627	51.3	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	51.6	
2627	196.	193.	191.	188.	186.	184.	182.	180.	178.	176.	174.	172.	170.	168.	166.	164.	162.	160.	158.	156.	154.	152.	150.	148.	146.	144.	142.
2616	49.2	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	49.6	
2616	201.	198.	196.	193.	190.	188.	186.	184.	182.	180.	178.	176.	174.	172.	170.	168.	166.	164.	162.	160.	158.	156.	154.	152.	150.	148.	146.
2604	46.3	46.3	46.2	46.1	45.9	45.7	45.6	45.4	45.3	45.2	45.1	45.0	44.9	44.8	44.7	44.6	44.5	44.4	44.3	44.2	44.1	44.0	43.9	43.8	43.7	43.6	43.5
2604	201.	199.	197.	195.	193.	191.	189.	187.	185.	183.	181.	179.	177.	175.	173.	171.	169.	167.	165.	163.	161.	159.	157.	155.	153.	151.	149.
2758	43.9	42.9	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	42.8	
2758	204.	203.	202.	201.	200.	199.	198.	197.	196.	195.	194.	193.	192.	191.	190.	189.	188.	187.	186.	185.	184.	183.	182.	181.	180.	179.	178.